

Second-order Lagrangian description of tri-dimensional gravity wave interactions

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We revisit and supplement the description of gravity waves based on perturbation expansions in Lagrangian coordinates. A general analytical framework is developed to derive a second-order Lagrangian solution to the motion of arbitrary surface gravity wave fields in a compact and vectorial form. The result is shown to be consistent with the classical second-order Eulerian expansion by Longuet-Higgins (*J. Fluid Mech.*, vol. 17, 1963, pp. 459–480) and is used to improve the original derivation by Pierson (1961 Models of random seas based on the Lagrangian equations of motion. *Tech. Rep.* New York University) for long-crested waves. As demonstrated, the Lagrangian perturbation expansion captures nonlinearities to a higher degree than does the corresponding Eulerian expansion of the same order. At the second order, it can account for complex nonlinear phenomena such as wave-front deformation that we can relate to the initial stage of horseshoe-pattern formation and the Benjamin–Feir modulational instability to shed new light on the origins of these mechanisms.

Key words: surface gravity waves, waves/free-surface flows

1. Introduction

The Lagrangian description of interactions between multiple surface gravity waves was pioneered by Pierson (1961) half a century ago. Pierson explicitly derived a first-order solution for two-dimensional surfaces and pushed the calculation to the second order for long-crested surfaces. He showed that first-order results of a Lagrangian analysis included more realistic features than did using its Eulerian counterpart, such as sharp crests and flat troughs. In the present work, we revisit and correct this classical analysis to provide a general analytical framework, and to derive a compact and vectorial form of a second-order Lagrangian description of arbitrary tri-dimensional gravity wave fields. The analysis of tri-dimensional multiple wave systems is much richer than the analysis of long-crested surfaces or monochromatic waves as some geometrical and dynamical characteristics of the wave field can only be accounted for by considering interactions between different, non-aligned free wave vectors.

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To date, exploration of the numerical and analytical possibilities offered by the Lagrangian formalism somehow have been overlooked. A renewed interest in Lagrangian approaches and their mathematical (Yakubovich & Zenkovich 2001; Buldakov, Taylor & Taylor 2006; Clamond 2007) or practical implications (Gjosund 2003; Fouques, Krogstad & Myrhaug 2006; Fouques & Stansberg 2009) has arisen and has provided the means to better evaluate the statistical and geometrical description of free surface and mass transport (Socquet-Juglard *et al.* 2005; Lindgren 2006; Aberg 2007; Aberg & Lindgren 2008; Noguier, Guérin & Chapron 2009; Hsu, Chen & Wang 2010; Hsu, Ng & Hwung 2012). The underlying reason behind these remarkable properties is that the Lagrangian representation is clearly well-suited to the description of steep waves and is a very useful mathematical tool for the correct evaluation of statistical quantities (such as height, slope and curvature distribution) of random gravity wave fields at limited costs in terms of analytical complexity.

Our main finding is given by (4.47) which summarizes the Lagrangian expressions of second-order displacements of water particles and pressure in the whole fluid domain. The analysis is restricted to infinite depth but there is no conceptual difficulty in relaxing this assumption. Full consistency with the second-order Eulerian expansion of Longuet-Higgins (1963) is demonstrated. Pierson's (1961) original second-order Lagrangian solution for long-crested waves is discussed and adjusted to agree with both Longuet-Higgins (1963) and our own derivations. We further discuss two remarkable phenomena which are not captured by second-order Eulerian expansions. First, the formation of horseshoe patterns is identified as being the result of a non-isotropic drift current. Second, Benjamin–Feir modulational instability is also revealed to be inherently present in the second-order Lagrangian framework as a simple beat effect between two neighbouring harmonics instead of an energy exchange between carrier and sideband waves.

2. Eulerian versus Lagrangian expansions

We shall consider an incompressible fluid of constant density ρ and of infinite depth, subject only to the restoring force of gravity (surface tension and viscosity are ignored). The pressure is set to a constant: p_a at the free surface of the fluid. A fixed system of axes $(\hat{x}, \hat{y}, \hat{z})$ with upwards-directed vertical vector \hat{z} is chosen.

2.1. Eulerian description

In the Eulerian description, any position in space is identified by its coordinates (x, y, z) , which can be decomposed into its horizontal projection $\mathbf{r} = (x, y)$ and vertical elevation z . Under potential assumption, the evolving field of gravity waves is described by its elevations $\eta(\mathbf{r}, t)$ and velocity potential $\Phi(x, y, z, t)$, with t being the time variable. Under Eulerian coordinates the potential solves Laplace's equation inside the volume together with dynamic and kinematic conditions at the borders:

$$\left. \begin{aligned} \Delta \Phi &= 0, & z < \eta(x, y, t) \\ \lim_{z \rightarrow -\infty} \nabla \Phi &= 0, \\ \Phi_t + \frac{1}{2} \nabla \Phi \cdot \nabla \Phi &= -g\eta, & z = \eta(x, y, t) \\ \Phi_z &= \eta_t + \nabla \eta \cdot \nabla \Phi, & z = \eta(x, y, t), \end{aligned} \right\} \quad (2.1)$$

where g is the acceleration due to gravity. System (2.1) describes potential waves. In the classical perturbative approach (Hasselmann 1962; Longuet-Higgins 1963; Weber

& Barrick 1977) the field of elevation η and the velocity potential $\underline{\Phi}$ at the position and time (\mathbf{r}, t) are sought in the form

$$\left. \begin{aligned} \eta &= \eta_0 + \eta_1 + \eta_2 + \dots \\ \underline{\Phi} &= \underline{\Phi}_0 + \underline{\Phi}_1 + \underline{\Phi}_2 + \dots \end{aligned} \right\} \quad (2.2)$$

The naught terms, η_0 and $\underline{\Phi}_0$, are the reference solutions corresponding to a flat fluid interface and the next terms, η_1 and $\underline{\Phi}_1$, are the solutions provided by the linearized equations. The successive terms, η_n and $\underline{\Phi}_n$, are n th-order corrections with respect to one small parameter. In the general case of multiple waves, this small parameter is not well identified but in the case of a monochromatic wave it can be linked to the wave steepness.

2.2. Lagrangian description

In the Lagrangian approach (Lamb 1932), fluid evolution is described by the motion of fluid particles. The spatial coordinates $\mathbf{R} = (x, y, z)$ of the particles now depend on their independent reference labels $\boldsymbol{\zeta} = (\alpha, \beta, \delta)$ and time t , that is explicitly $x = x(\alpha, \beta, \delta, t)$, $y = y(\alpha, \beta, \delta, t)$ and $z = z(\alpha, \beta, \delta, t)$. Hereafter $\boldsymbol{\zeta}$ is chosen to be the locus of particles at rest. For ease of reading we shall introduce dedicated notations for the horizontal component of particle labels and positions, $\boldsymbol{\xi} = (\alpha, \beta)$ and $\mathbf{r} = (x, y)$, respectively.

The evolution of particle coordinates is driven by Newton’s law of dynamics:

$$\mathbf{R}_{tt} + g\hat{\mathbf{z}} = -\frac{1}{\rho}\nabla_{\mathbf{R}}p \quad (2.3)$$

where $p = p(\mathbf{R})$ is the local pressure. This dynamical equation is coupled with the continuity equation:

$$|\mathbb{J}| = 1; \quad \frac{\partial}{\partial t}|\mathbb{J}| = 0 \quad \text{with } \mathbb{J} = \begin{bmatrix} x_\alpha & y_\alpha & z_\alpha \\ x_\beta & y_\beta & z_\beta \\ x_\delta & y_\delta & z_\delta \end{bmatrix}. \quad (2.4a,b)$$

Multiplying (2.3) by \mathbb{J} gives

$$\mathbb{J}\mathbf{R}_{tt} + g\nabla(\mathbf{R} \cdot \hat{\mathbf{z}}) + \frac{1}{\rho}\nabla p = 0 \quad (2.5)$$

which is the basic equation given by Lamb (1932). From now on the spatial gradient relative to the independent Lagrangian variables (α, β, δ) will be denoted by ∇ .

Solutions to these equations need not be irrotational. However, if a function $F(\boldsymbol{\zeta}, t)$ can be found such that

$$dF = (\mathbb{J}\mathbf{R}_t) \cdot d\boldsymbol{\zeta} \quad (2.6)$$

is a perfect differential then there is no vorticity (see § A.1). Here $d\boldsymbol{\zeta} = (d\alpha, d\beta, d\delta)$ denotes an infinitesimal label variation. Following the methodology described by Stoker (1957) we may seek at solution in the form of a simultaneous perturbation expansion for position, pressure and the vorticity function:

$$\left. \begin{aligned} \mathbf{R} &= \mathbf{R}_0 + \mathbf{R}_1 + \mathbf{R}_2 + \dots \\ p &= p_a - \rho g\delta + p_1 + p_2 + \dots \\ F &= F_0 + F_1 + F_2 + \dots \end{aligned} \right\} \quad (2.7)$$

where the naught variables refer to particles at rest.

3. First-order solution: the Gerstner wave

Let us map the fluid domain onto the half-space $\delta \leq 0$. From now on, $\delta = 0$ corresponds to the free surface η under pressure p_a . The zeroth-order solution to expansion (2.7) is related to particles at rest and can be written

$$\left. \begin{aligned} \mathbf{R}_0 &= \boldsymbol{\zeta}, \\ p_0 &= p_a - \rho g \delta \\ F_0 &= 0 \\ |\mathbb{J}| &= 1. \end{aligned} \right\} \tag{3.1}$$

First-order quantities are solutions to linearized Lagrangian equations. When taken at the first order, equation (2.5) can be written

$$\mathbf{R}_{1tt} + g \nabla(\mathbf{R}_1 \cdot \hat{\mathbf{z}}) + \frac{1}{\rho} \nabla p_1 = 0 \tag{3.2}$$

and the continuity equation is expressed by

$$x_{1\alpha} + y_{1\beta} + z_{1\delta} = \nabla \cdot \mathbf{R}_1 = 0. \tag{3.3}$$

In order to simplify the calculations presented in the next section and to ensure an irrotational solution at the first order (see (3.9) below), we shall investigate solutions of the form $\mathbf{R}_1 = \nabla w$ in an effort to see whether there exists a function for $w(\boldsymbol{\zeta}, t)$. This last quantity must satisfy the following equation:

$$\nabla(w_{tt} + gw_{\delta} + p_1/\rho) = 0. \tag{3.4}$$

Setting p_1 to 0 at $\delta = 0$ gives

$$w = \cos(\mathbf{k} \cdot \boldsymbol{\xi} - \omega t) e^{k\delta}; \quad \omega^2 = gk; \quad p_1 = 0 \tag{3.5a-c}$$

where $\mathbf{k} = (k_\alpha, k_\beta)$ is an independent bi-dimensional vector in the (α, β) domain and k a constant parameter. At the first order in ϵ , the relation of continuity (3.3) can be written

$$\Delta w = (-k_\alpha^2 - k_\beta^2 + k^2)w = 0 \tag{3.6}$$

leading to $\|\mathbf{k}\| = k$. As \mathbf{R}_1 is a spatial displacement, a suitable solution is $\mathbf{R}_1 = \nabla(ak^{-1}w)$ which leads to the first-order solution:

$$\left. \begin{aligned} \mathbf{r} &= \boldsymbol{\xi} - a\hat{\mathbf{k}} \sin(\mathbf{k} \cdot \boldsymbol{\xi} - \omega t) e^{k\delta} \\ z &= \delta + a \cos(\mathbf{k} \cdot \boldsymbol{\xi} - \omega t) e^{k\delta} \\ p &= p_0 - \rho g \delta. \end{aligned} \right\} \tag{3.7}$$

From now on, we shall use the notation $\hat{\mathbf{k}} = \mathbf{k}/k$ for the direction of a vector \mathbf{k} and k for its norm. This solution describes the trajectories of water particles as circles whose radii decrease exponentially with water depth. The spatial profile of such waves is a trochoid moving in the direction \mathbf{k} with a crest to trough wave amplitude defined by a , being the circle radius of the trajectories of particles at the free surface.

Two centuries ago, Gerstner (1809) derived an exact solution to the equation of motion (2.3) and obtained the same solution (3.7) for water particle trajectories (\mathbf{r}, z) with, however, a slightly different pressure term:

$$p = p_0 - \rho g \delta + \frac{1}{2} k^{-2} \rho \omega^2 e^{2k\delta}. \tag{3.8}$$

The Gerstner wave has been described in classical textbooks (e.g. Lamb 1932; Kinsman 1965) even though its stability was investigated only recently (Naciri & Mei 1992; Leblanc 2004). It has always been criticized in view of its non-vanishing vorticity. This calls for some discussion on the presence of vorticity. Wind waves do in general exhibit vorticity, although it is in fact low. The main reason for which most of the studies have been devoted to irrotational waves is the considerable simplification offered by potential theory in the analytical derivations. It turns out that the predictions of potential theory agree reasonably well with observations, which does not mean that real waves are irrotational but rather that vorticity has only secondary effects. However, discrepancies are bound to become visible as the quality and accuracy of observations improve and it will soon become necessary to account for vorticity. The main shortcoming of the Gerstner solution is that it does not address a wide class of solutions with low vorticity. Its vorticity has in fact a very special distribution and there is no rationale for it to be more relevant than any other distribution of the same order. In the present analytical framework, the construction of a weakly nonlinear solution to the exact inviscid equations is more general and it is possible to examine arbitrary distributions with low vorticity and evaluate, at least coarsely, the importance of this effect.

As already derived by Pierson (1961), equation (2.6) at first order in ϵ can be written

$$dF = \mathbf{R}_1 \cdot d\boldsymbol{\zeta} \tag{3.9}$$

which is a perfect differential of F_1 since $dF = \nabla F_1 \cdot d\boldsymbol{\zeta}$ with

$$F_1 = ak^{-1}w_t = \frac{a\omega}{k} \sin(\mathbf{k} \cdot \boldsymbol{\xi} - \omega t) e^{k\delta}. \tag{3.10}$$

Therefore, there is no vorticity at the first order and the Gerstner wave (with the corresponding pressure given by (3.7)) is an irrotational solution at the considered order of the expansion.

Owing to the linearity of (2.5), we can write an extended solution to the first-order equations as a continuous superposition of independent harmonics defined by their wavenumber \mathbf{k} in the form:

$$\mathbf{R}_1 = \nabla \Phi_1 \quad \text{with } \Phi_1 = \phi_1 + \text{c.c.}, \tag{3.11}$$

where ‘c.c.’ designates the complex conjugate of a given quantity and

$$\phi_1 = \frac{1}{2} \int_{\mathbb{R}^2} \frac{A(\mathbf{k})}{k} e^{i(\mathbf{k} \cdot \boldsymbol{\xi} - \omega t)} e^{k\delta} d\mathbf{k}. \tag{3.12}$$

Here $A(\mathbf{k})$ is the orbital amplitude and the factor 1/2 accounts for the complex plus conjugate formulation of Φ_1 . Such an orbital spectrum has been already introduced in the statistical studies of Lagrangian wave fields (Pierson 1961; Daemrich & Woltering 2008; Lindgren & Lindgren 2011) and describes the spectral content of particle motion. It is sometimes termed the ‘undressed’ spectrum (Elfouhaily *et al.* 1999) when it refers to a nonlinear transformation of an underlying linear surface (Creamer *et al.* 1989).

In the present state of knowledge, establishing the relationship between the orbital (Lagrangian) and the surface (Eulerian) spectrum is still an issue. When the amplitude $A(\mathbf{k})$ is taken to be a complex random variable with independent uniformly distributed random phases, the resulting function ϕ_1 is a complex random Gaussian process by

virtue of the law of large numbers. However, the random surface η defined by the locus of particles at the free surface is no longer Gaussian. This implies that the corresponding distribution of elevation, slopes and curvatures distributions deviate from the normal distribution. Statistical properties of such random wave fields have been studied in detail (e.g. Pierson 1961; Gjosund 2003; Aberg & Lindgren 2008; Lindgren & Aberg 2009; Nouguier *et al.* 2009; Lindgren & Lindgren 2011) and have been found to be more consistent with ocean wave field measurements. *A contrario*, it should be noted that a first-order expansion in the Eulerian framework, which expresses the surface and its derivatives as a linear superposition of free harmonics, is bound to the Gaussian statistics.

4. Second-order Lagrangian solution

This section is devoted to the second-order Lagrangian expansion. We recall the corresponding equations and detail the calculations to derive the second-order displacements and the pressure terms as functions of the Lagrangian variables. To simplify notation we shall omit the integration elements ($d\mathbf{k}, d\mathbf{k}'$) and domains (\mathbb{R}^2 and $\mathbb{R}^2 \times \mathbb{R}^2$) in the following single and double integrals.

4.1. Second-order equations

Retaining the second-order terms in (2.5) we obtain:

$$\mathbf{R}_{2tt} + g\nabla z_2 + \nabla p_2/\rho = -\mathbf{H}(\Phi_1)\nabla\Phi_{1tt}, \tag{4.1}$$

where \mathbf{H} is the Hessian operator, that is the square matrix built with the second-order partial derivatives relative to the (α, β, δ) variables:

$$\mathbf{H}(\Phi_1) = \begin{bmatrix} \partial_{\alpha\alpha}^2 & \partial_{\alpha\beta}^2 & \partial_{\alpha\delta}^2 \\ \partial_{\beta\alpha}^2 & \partial_{\beta\beta}^2 & \partial_{\beta\delta}^2 \\ \partial_{\delta\alpha}^2 & \partial_{\delta\beta}^2 & \partial_{\delta\delta}^2 \end{bmatrix} \Phi_1. \tag{4.2}$$

For practical purposes we rewrite the right-hand side of (4.1) as

$$-\mathbf{H}(\Phi_1)\nabla\Phi_{1tt} = \mathbf{S} + \mathbf{T} \tag{4.3}$$

with

$$\mathbf{S} = (S^\alpha, S^\beta, S^\delta) = -\mathbf{H}(\phi_1)\nabla\phi_{1tt} + \text{c.c.} \tag{4.4}$$

$$\mathbf{T} = (T^\alpha, T^\beta, T^\delta) = -\mathbf{H}(\phi_1)\nabla\phi_{1tt}^* + \text{c.c.} \tag{4.5}$$

where the superscript ‘*’ refers to the complex conjugate. Straightforward derivations given in § A.2 lead to

$$\left. \begin{aligned} (S^\alpha, S^\beta) &= \iint \mathcal{N} g k k' \frac{i}{2} (\widehat{\mathbf{k}} + \widehat{\mathbf{k}}') + \text{c.c.} \\ S^\delta &= \iint \mathcal{N} g k k' + \text{c.c.} \end{aligned} \right\} \tag{4.6}$$

and

$$\left. \begin{aligned} (T^\alpha, T^\beta) &= \iint \underline{\mathcal{N}} g k k' \frac{i}{2} (\widehat{\mathbf{k}} - \widehat{\mathbf{k}}') + \text{c.c.} \\ T^\delta &= \iint \underline{\mathcal{N}} g k k' + \text{c.c.} \end{aligned} \right\} \tag{4.7}$$

where the kernels \mathcal{N} and $\underline{\mathcal{N}}$ depend on the variables $\mathbf{k}, \mathbf{k}', \boldsymbol{\xi}, \delta$ and t and are defined as follows:

$$\mathcal{N} = \mathcal{B}e^{-i(\omega+\omega')t}e^{i(\mathbf{k}+\mathbf{k}')\cdot\boldsymbol{\xi}} \quad \text{and} \quad \underline{\mathcal{N}} = \underline{\mathcal{B}}e^{-i(\omega-\omega')t}e^{i(\mathbf{k}+\mathbf{k}')\cdot\boldsymbol{\xi}} \quad (4.8a,b)$$

with

$$\mathcal{B}(\mathbf{k}, \mathbf{k}', \boldsymbol{\xi}) = \frac{1}{4}(1 - \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}')A(\mathbf{k})A(\mathbf{k}')e^{i(\mathbf{k}+\mathbf{k}')\cdot\boldsymbol{\xi}} \quad (4.9)$$

$$\underline{\mathcal{B}}(\mathbf{k}, \mathbf{k}', \boldsymbol{\xi}) = \frac{1}{4}(1 + \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}')A(\mathbf{k})A^*(\mathbf{k}')e^{i(\mathbf{k}-\mathbf{k}')\cdot\boldsymbol{\xi}}. \quad (4.10)$$

Analogously, the continuity equation (2.4) at the second order can be written

$$x_{2\alpha} + y_{2\beta} + z_{2\delta} + \Phi_{1\alpha\alpha}\Phi_{1\beta\beta} + \Phi_{1\alpha\alpha}\Phi_{1\delta\delta} + \Phi_{1\beta\beta}\Phi_{1\delta\delta} - \Phi_{1\alpha\beta}^2 - \Phi_{1\alpha\delta}^2 - \Phi_{1\beta\delta}^2 = 0 \quad (4.11)$$

and can be rewritten in the form (see §A.3)

$$\nabla \cdot \mathbf{R}_2 = V + W \quad (4.12)$$

with

$$V = \iint \frac{1}{2}(kk' - \mathbf{k} \cdot \mathbf{k}')\mathcal{N} + \text{c.c.} \quad \text{and} \quad W = \iint \frac{1}{2}(kk' + \mathbf{k} \cdot \mathbf{k}')\underline{\mathcal{N}} + \text{c.c.} \quad (4.13a,b)$$

4.2. Second-order expressions

Due to the linearity of (4.1) and (4.12), we shall first consider the solution to (4.1) with the sole \mathbf{S} term on the right-hand side, that is,

$$\mathbf{R}_{2tt} + g\nabla z_2 + \nabla p_2/\rho = \mathbf{S}. \quad (4.14)$$

We furthermore assume that \mathbf{r}_2, z_2 and p_2 can be written as the following integrals:

$$\mathbf{r}_2 = \iint \mathcal{N}i\mathcal{R} + \text{c.c.} \quad (4.15)$$

$$z_2 = \iint \mathcal{N}\mathcal{Z} + \text{c.c.} \quad (4.16)$$

$$p_2 = \rho g \iint \mathcal{N}\mathcal{P} + \text{c.c.}, \quad (4.17)$$

where \mathcal{R}, \mathcal{Z} and \mathcal{P} are unknown kernels depending on \mathbf{k} and \mathbf{k}' . Inserting these expressions in (4.14) leads to a set of equations for the kernels:

$$\left. \begin{aligned} -\mathcal{Z}\Omega^+ + (\mathcal{Z} + \mathcal{P})(k + k') &= kk' \\ -\Omega^+\mathcal{R} + (\mathbf{k} + \mathbf{k}')(\mathcal{Z} + \mathcal{P}) &= \frac{1}{2}kk'(\widehat{\mathbf{k}} + \widehat{\mathbf{k}}'), \end{aligned} \right\} \quad (4.18)$$

where we have defined

$$\Omega^\pm = \left(\sqrt{k} \pm \sqrt{k'}\right)^2. \quad (4.19)$$

Inserting again (4.15)–(4.17) into (4.12) and keeping only the terms involving the kernel \mathcal{N} leads to a third equation:

$$-\mathcal{R} \cdot (\mathbf{k} + \mathbf{k}') + \mathcal{Z}(k + k') = \frac{1}{2}(kk' - \mathbf{k} \cdot \mathbf{k}'). \quad (4.20)$$

Equations (4.18) and (4.20) can easily be solved leading to

$$\left. \begin{aligned} \mathcal{R} &= \frac{\omega k + \omega' k'}{2(\omega + \omega')} \\ \mathcal{Z} &= \frac{1}{4}(k + k' + \Omega^-) \\ \mathcal{P} &= \sqrt{kk'}. \end{aligned} \right\} \tag{4.21}$$

Analogously, we have to solve (4.1) with the sole term T on the right-hand side, which gives

$$\mathbf{R}_{2it} + g \nabla z_2 + \nabla p_2 / \rho = T. \tag{4.22}$$

Again, we assume that there exist r_2, z_2 and p_2 in the form given in (4.15)–(4.17) with some other kernels $\underline{\mathcal{N}}, \underline{\mathcal{R}}, \underline{\mathcal{Z}}$ and $\underline{\mathcal{P}}$. Since (4.12) involving the kernel \mathcal{N} has already been solved, the only remaining terms are those involving $\underline{\mathcal{N}}$ in (4.12). A set of three equations is thus obtained for the unknown kernels:

$$\left. \begin{aligned} -\underline{\mathcal{Z}}\Omega^- + (\underline{\mathcal{Z}} + \underline{\mathcal{P}})(k + k') &= kk' \\ -\Omega^- \underline{\mathcal{R}} + (\mathbf{k} - \mathbf{k}')(\underline{\mathcal{Z}} + \underline{\mathcal{P}}) &= \frac{1}{2}kk'(\widehat{\mathbf{k}} - \widehat{\mathbf{k}}') \\ -\underline{\mathcal{R}} \cdot (\mathbf{k} - \mathbf{k}') + \underline{\mathcal{Z}}(k + k') &= \frac{1}{2}(kk' + \mathbf{k} \cdot \mathbf{k}'). \end{aligned} \right\} \tag{4.23}$$

Again, this system can easily be solved, leading to

$$\left. \begin{aligned} \underline{\mathcal{R}} &= \frac{\omega k + \omega' k'}{2(\omega - \omega')} \\ \underline{\mathcal{Z}} &= \frac{1}{4}(k + k' + \Omega^+) \\ \underline{\mathcal{P}} &= -\sqrt{kk'} \end{aligned} \right\} \text{ if } \omega \neq \omega'. \tag{4.24}$$

The case $\omega = \omega'$ will be discussed in detail in §4.3. At this point we have found a solution to the second-order Lagrangian expansion (4.1) in the form

$$\mathbf{r}_2 = \iint i(\mathcal{N}\mathcal{R} + \underline{\mathcal{N}}\underline{\mathcal{R}}) + \text{c.c.} \tag{4.25}$$

$$z_2 = \iint (\mathcal{N}\mathcal{Z} + \underline{\mathcal{N}}\underline{\mathcal{Z}}) + \text{c.c.} \tag{4.26}$$

$$p_2 = \rho g \iint (\mathcal{N}\mathcal{P} + \underline{\mathcal{N}}\underline{\mathcal{P}}) + \text{c.c.} \tag{4.27}$$

However, the expression of p_2 does not satisfy the boundary condition $p_2 = 0$ at $\delta = 0$ and needs to be corrected. Noting that $\mathcal{N} = \mathcal{B}e^{-i(\omega+\omega')t} e^{(k+k')\delta}$ and $\underline{\mathcal{N}} = \underline{\mathcal{B}}e^{-i(\omega-\omega')t} e^{(k+k')\delta}$, a very simple way to satisfy to the boundary condition is to complete p_2 in the form:

$$p_2 = \rho g \iint \mathcal{P}\mathcal{B}e^{-i(\omega+\omega')t} (e^{(k+k')\delta} - e^{K^+\delta}) + \underline{\mathcal{P}}\underline{\mathcal{B}}e^{-i(\omega-\omega')t} (e^{(k+k')\delta} - e^{K^-\delta}) + \text{c.c.} \tag{4.28}$$

where the additional kernels K^+ and K^- must be determined. The pressure at the second order can thus be written as

$$p_2 = \iint ((\mathcal{N} - \mathcal{N}')\mathcal{P} + (\underline{\mathcal{N}} - \underline{\mathcal{N}}')\underline{\mathcal{P}}) + \text{c.c.} \tag{4.29}$$

where we have introduced the two kernels \mathcal{N}' and $\underline{\mathcal{N}}'$, which only differ from \mathcal{N} and $\underline{\mathcal{N}}$, respectively, by the real exponential term:

$$\mathcal{N}' = \mathcal{B}e^{-i(\omega+\omega')t}e^{K^+\delta} \quad \text{and} \quad \underline{\mathcal{N}}' = \underline{\mathcal{B}}e^{-i(\omega-\omega')t}e^{K^-\delta}. \tag{4.30a,b}$$

It is therefore natural to assume a complete expression of r_2 and z_2 in the form

$$r_2 = \iint i(\mathcal{N}'\mathcal{R} - \mathcal{N}'\mathcal{R}' + \underline{\mathcal{N}}'\underline{\mathcal{R}} - \underline{\mathcal{N}}'\underline{\mathcal{R}}') + \text{c.c.} \tag{4.31}$$

$$z_2 = \iint (\mathcal{N}'\mathcal{Z} - \mathcal{N}'\mathcal{Z}' + \underline{\mathcal{N}}'\underline{\mathcal{Z}} - \underline{\mathcal{N}}'\underline{\mathcal{Z}}') + \text{c.c.}, \tag{4.32}$$

where the primed kernels need to be found. To achieve this, we recall these expressions in (4.1) and identify the terms pertaining to the \mathcal{N}' kernel only. This leads to the following equations:

$$-\Omega^+\mathcal{Z}' + K^+(\mathcal{Z}' + P) = 0 \tag{4.33}$$

$$-\Omega^+\mathcal{R}' + (\mathbf{k} + \mathbf{k}')(\mathcal{Z}' + P) = 0, \tag{4.34}$$

as well as

$$-\mathcal{R}' \cdot (\mathbf{k} + \mathbf{k}') + K^+\mathcal{Z}' = 0 \tag{4.35}$$

from the continuity equation. Inserting (4.33) into (4.34) and multiplying by $(\mathbf{k} + \mathbf{k}')/\Omega^+$ leads to

$$-\mathcal{R}' \cdot (\mathbf{k} + \mathbf{k}') + \frac{\|\mathbf{k} + \mathbf{k}'\|^2}{K^+}\mathcal{Z}' = 0 \tag{4.36}$$

which is consistent with (4.35) if and only if

$$K^+ = \|\mathbf{k} + \mathbf{k}'\| \tag{4.37}$$

(we discard the mathematical solution $K^+ = -\|\mathbf{k} + \mathbf{k}'\|$ which is non-physical because of the asymptotic constraint $p_2 \rightarrow 0$ when $\delta \rightarrow -\infty$). We can now solve (4.33) and (4.34) to obtain

$$\left. \begin{aligned} \mathcal{R}' &= \frac{\sqrt{kk'}(\mathbf{k} + \mathbf{k}')}{\Omega^+ - \|\mathbf{k} + \mathbf{k}'\|} \\ \mathcal{Z}' &= \frac{\sqrt{kk'}\|\mathbf{k} + \mathbf{k}'\|}{\Omega^+ - \|\mathbf{k} + \mathbf{k}'\|} \end{aligned} \right\} \tag{4.38}$$

Repeating the same procedure with the kernel $\underline{\mathcal{N}}'$ leads to another set of equations:

$$\left. \begin{aligned} -\Omega^-\underline{\mathcal{Z}}' + K^-(\underline{\mathcal{Z}}' + P) &= 0 \\ -\Omega^-\underline{\mathcal{R}}' + (\mathbf{k} - \mathbf{k}')(\underline{\mathcal{Z}}' + P) &= 0 \\ -\underline{\mathcal{R}}' \cdot (\mathbf{k} - \mathbf{k}') + K^-\underline{\mathcal{Z}}' &= 0 \end{aligned} \right\} \tag{4.39}$$

which admit the solution $K^- = \|\mathbf{k} - \mathbf{k}'\|$ and

$$\left. \begin{aligned} \underline{\mathcal{R}}' &= \frac{-\sqrt{kk'}(\mathbf{k} - \mathbf{k}')}{\Omega^- - \|\mathbf{k} - \mathbf{k}'\|} \\ \underline{\mathcal{Z}}' &= \frac{-\sqrt{kk'}\|\mathbf{k} - \mathbf{k}'\|}{\Omega^- - \|\mathbf{k} - \mathbf{k}'\|} \end{aligned} \right\} \tag{4.40}$$

4.3. Interaction of harmonics of equal frequencies

The complete kernels involved in the integral representation of \mathbf{R}_2 and p_2 have now been found. However, in order to complete the solution to the second-order Lagrangian equations we need to discuss the case $\omega = \omega'$ which was initially discarded in (4.24).

A generalized expression of the horizontal second-order term corresponding to kernel solutions (4.24) for the case $\omega = \omega'$ would be written as the limit

$$\mathbf{r}_2 = \lim_{\gamma \rightarrow 0} \iint_{\mathbb{R}^4 - \mathcal{E}} i \mathcal{RN} + \text{c.c.} \tag{4.41}$$

where \mathcal{E} is the \mathbb{R}^4 subdomain so that $|\omega - \omega'| < \gamma$ and where \mathcal{R} , defined at (4.24), contains a singularity at $\omega = \omega'$.

If this integral were to admit a finite value, it would have to be defined in the sense of Cauchy principal value (PV):

$$\mathbf{r}_2 = \text{PV} \iint_{\mathbb{R}^4} i \frac{\omega \mathbf{k} + \omega' \mathbf{k}'}{2(\omega - \omega')} \mathcal{B} e^{-i(\omega - \omega')t} e^{(k+k')\delta} + \text{c.c.} \tag{4.42}$$

The existence of the finite limit (4.41) is shown in appendix B ensuring that (4.42) is the correct expression of \mathbf{r}_2 .

4.4. Second-order vorticity

A complete Lagrangian second-order solution has now been found. We can verify *a posteriori* that it is indeed irrotational. For this, we have to investigate the second-order expression of the function dF , that is,

$$dF_2 = [\mathbf{R}_{2t} + \mathbf{H}(\Phi_1) \nabla(\Phi_{1t})] \cdot d\boldsymbol{\zeta}, \tag{4.43}$$

which we chose to rewrite in the form

$$dF_2 = \int_t dt [\mathbf{R}_{2tt} + \mathbf{H}(\Phi_1) \nabla(\Phi_{1tt}) + \mathbf{H}(\Phi_{1t}) \nabla(\Phi_{1t})] \cdot d\boldsymbol{\zeta} \tag{4.44}$$

where the symbol $\int_t dt$ refers to temporal integration. Inserting (4.1) into this last expression leads to

$$dF_2 = \nabla \left[\int_t dt \left(-(gz_2 + p_2/\rho) + \frac{1}{2}(\Phi_{1t})^2 \right) \right] \cdot d\boldsymbol{\zeta}. \tag{4.45}$$

This provides F_2 in the form:

$$F_2 = \int_t dt \left(\frac{1}{2}(\Phi_{1t})^2 - gz_2 - p_2/\rho \right). \tag{4.46}$$

The existence of such a function F_2 warrants the absence of vorticity at the second order.

4.5. Second-order solution

To summarize all of the expressions established previously, the general solution to the second-order terms of (2.7) can be written as follows:

$$\left. \begin{aligned}
 \mathbf{r}_2 &= \iint \mathbf{i} \left(\frac{\omega \mathbf{k} + \omega' \mathbf{k}'}{2(\omega + \omega')} e^{(k+k')\delta} - \frac{\sqrt{kk'}(\mathbf{k} + \mathbf{k}')}{\Omega^+ - \|\mathbf{k} + \mathbf{k}'\|} e^{\|\mathbf{k} + \mathbf{k}'\|\delta} \right) \mathcal{B} e^{-i(\omega + \omega')t} + \text{c.c.} \\
 &+ \text{PV} \iint \mathbf{i} \left(\frac{\omega \mathbf{k} + \omega' \mathbf{k}'}{2(\omega - \omega')} e^{(k+k')\delta} + \frac{\sqrt{kk'}(\mathbf{k} - \mathbf{k}')}{\Omega^- - \|\mathbf{k} - \mathbf{k}'\|} e^{\|\mathbf{k} - \mathbf{k}'\|\delta} \right) \underline{\mathcal{B}} e^{-i(\omega - \omega')t} + \text{c.c.} \\
 z_2 &= \iint \left(\frac{k + k' + \Omega^-}{4} e^{(k+k')\delta} - \frac{\sqrt{kk'}\|\mathbf{k} + \mathbf{k}'\|}{\Omega^+ - \|\mathbf{k} + \mathbf{k}'\|} e^{\|\mathbf{k} + \mathbf{k}'\|\delta} \right) \mathcal{B} e^{-i(\omega + \omega')t} + \text{c.c.} \\
 &+ \iint \left(\frac{k + k' + \Omega^+}{4} e^{(k+k')\delta} + \frac{\sqrt{kk'}\|\mathbf{k} - \mathbf{k}'\|}{\Omega^- - \|\mathbf{k} - \mathbf{k}'\|} e^{\|\mathbf{k} - \mathbf{k}'\|\delta} \right) \underline{\mathcal{B}} e^{-i(\omega - \omega')t} + \text{c.c.} \\
 p_2 &= \rho g \iint \sqrt{kk'} \left(e^{(k+k')\delta} - e^{\|\mathbf{k} + \mathbf{k}'\|\delta} \right) \mathcal{B} e^{-i(\omega + \omega')t} + \text{c.c.} \\
 &- \rho g \iint \sqrt{kk'} \left(e^{(k+k')\delta} - e^{\|\mathbf{k} - \mathbf{k}'\|\delta} \right) \underline{\mathcal{B}} e^{-i(\omega - \omega')t} + \text{c.c.}
 \end{aligned} \right\} \tag{4.47}$$

where \mathcal{B} and $\underline{\mathcal{B}}$ are defined in (4.9) and (4.10) and Ω^\pm in (4.19).

5. Comparison with classical models

5.1. Consistency with the Eulerian approach of M. S. Longuet-Higgins

Before investigating the consistency of this model with classical Eulerian models, it is instructive to establish the correspondence between Eulerian and Lagrangian expansions. Let us consider the surface $\eta(\mathbf{r}, t)$ implicitly defined by the locus of particle trajectories $(\mathbf{r}(t), z(t))$ and denote $\eta = \eta_0 + \eta_1 + \eta_2 + \dots$ as its Eulerian expansion in order of steepness above a reference plane. Applying successive Taylor expansions and making use of the correspondence between the (α, β, δ) Lagrangian labels and the (x, y, z) coordinate system of the Eulerian description, it can be easily shown that

$$\left. \begin{aligned}
 \eta_0 &= z_0 \\
 \eta_1 &= z_1 - \mathbf{r}_1 \cdot \nabla_\xi \eta_0 \\
 \eta_2 &= z_2 - \mathbf{r}_1 \cdot \nabla_\xi \eta_1 - \mathbf{r}_2 \cdot \nabla_\xi \eta_0 - \frac{1}{2} \mathbf{r}_1 \nabla_\xi \nabla_\xi \eta_0 \mathbf{r}_1 \\
 \eta_n &= z_n - \dots
 \end{aligned} \right\} \tag{5.1}$$

where ∇_ξ is the horizontal bi-dimensional gradient and $\nabla_\xi \nabla_\xi$ the corresponding Hessian. The expansion can in principle be pursued at an arbitrary order even though it becomes algebraically more complex. From this it can be seen that any n th-order term in the surface elevation (η_n) can be obtained from the combination of an n th-order term in the vertical particle position (z_n) and lower-order terms (\mathbf{r}_p, z_p) , $p \leq n - 1$. Hence, any given order of the Lagrangian expansion provides the complete corresponding Eulerian order and is moreover involved in higher-order Eulerian terms.

The classical Eulerian approach (Hasselmann 1962; Longuet-Higgins 1963) to the nonlinear theory of gravity waves consist in seeking both the elevation η and the velocity potential $\underline{\Phi}$ at the free surface in a perturbation series (2.2). The expansion

is usually performed about the mean horizontal plane of the leading order η_0 so that no zeroth-order term is present:

$$\eta(\boldsymbol{\xi}, t) = \eta_1(\boldsymbol{\xi}, t) + \eta_2(\boldsymbol{\xi}, t) + \dots \tag{5.2}$$

$$\Phi(\boldsymbol{\xi}, t) = \Phi_1(\boldsymbol{\xi}, t) + \Phi_2(\boldsymbol{\xi}, t) + \dots \tag{5.3}$$

In these two equations and the rest of this section, the fixed Eulerian coordinate system (x, y) has been simply replaced by the (α, β) system. The first-order terms are given by the classical spectral representation,

$$\eta_1(\boldsymbol{\xi}, t) = \sum_{j=1}^N a_j \cos \psi_j, \quad \psi_j = \mathbf{k}_j \cdot \boldsymbol{\xi} - \omega_j t + \varphi_j \tag{5.4}$$

$$\Phi_1(\boldsymbol{\xi}, t) = \sum_{j=1}^N b_j \cos \psi_j, \tag{5.5}$$

where φ_j is the phase associated to the \mathbf{k}_j component. The higher-order terms in the expansion involve n th-order multiplicative combinations of these spectral components. The perturbation expansions of elevation and velocity potential are identified simultaneously by injecting the successive Fourier expansions in Navier–Stokes equations. The leading, quadratic, nonlinear term for elevation was provided by Longuet-Higgins (1963) in the form

$$\eta_2(\boldsymbol{\xi}, t) = \frac{1}{2} \sum_{i,j=1}^N a_i a_j [K_{ij} \cos \psi_i \cos \psi_j + K'_{ij} \sin \psi_i \sin \psi_j], \tag{5.6}$$

$$K_{ij} = (k_i k_j)^{-1/2} [B_{ij}^- + B_{ij}^+ - \mathbf{k}_i \cdot \mathbf{k}_j] + k_i + k_j \tag{5.7}$$

$$K'_{ij} = (k_i k_j)^{-1/2} [B_{ij}^- - B_{ij}^+ - k_i k_j] \tag{5.8}$$

$$B_{ij}^\pm = \frac{\Omega_{ij}^\pm (\mathbf{k}_i \cdot \mathbf{k}_j \mp k_i k_j)}{\Omega_{ij}^\pm - \|\mathbf{k}_i \pm \mathbf{k}_j\|} \tag{5.9}$$

$$\Omega_{ij}^\pm = (\sqrt{k_i} \pm \sqrt{k_j})^2 \tag{5.10}$$

where as usual $k = \|\mathbf{k}\|$. (Note that the factor 1/2 is missing in the original paper by Longuet-Higgins, as was later acknowledged by the author himself.) The first-order Lagrangian expansion was shown to be close but not perfectly consistent with the second-order Eulerian perturbation expansion of Longuet-Higgins (see Nougier *et al.* 2009). We will now show that full consistency is achieved with Longuet-Higgins approach with the second-order Lagrangian expansion at the surface, that is,

$$\left. \begin{aligned} \eta_1 &= z_1 \\ \eta_2 &= z_2 - \mathbf{r}_1 \cdot \nabla_{\boldsymbol{\xi}} z_1 \end{aligned} \right\} \tag{5.11}$$

where, again, $\nabla_{\boldsymbol{\xi}}$ is the horizontal bi-dimensional gradient. It should be noted from (5.11) that \mathbf{r}_2 is absent emphasizing that the second-order Eulerian formalism misses all effects related to \mathbf{r}_2 contribution.

From (4.47), we have at the free surface ($\delta = 0$)

$$\begin{aligned} z_2 &= \frac{1}{2} \iint \left\{ \left(\frac{k + k' + \Omega^-}{4} - \frac{\sqrt{kk'} \|\mathbf{k} + \mathbf{k}'\|}{\Omega^+ - \|\mathbf{k} + \mathbf{k}'\|} \right) (1 - \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}') a_k a_{k'} \cos(\psi + \psi') \right. \\ &\quad \left. + \left(\frac{k + k' + \Omega^+}{4} + \frac{\sqrt{kk'} \|\mathbf{k} - \mathbf{k}'\|}{\Omega^- - \|\mathbf{k} - \mathbf{k}'\|} \right) (1 + \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}') a_k a_{k'} \cos(\psi - \psi') \right\} \tag{5.12} \end{aligned}$$

with $a_k = \|A(\mathbf{k})\|$ and $\psi = \mathbf{k} \cdot \boldsymbol{\xi} - \omega t + \varphi_k$ where φ_k is the phase of $A(\mathbf{k})$. Following some basic algebra z_2 can be rewritten in the form

$$z_2 = \frac{1}{2} \iint a_k a_{k'} \left[K \cos(\psi) \cos(\psi') + \left(K' + \mathbf{k}' \cdot \widehat{\mathbf{k}} + \mathbf{k} \cdot \widehat{\mathbf{k}} \right) \sin(\psi) \sin(\psi') \right] \quad (5.13)$$

where kernels K and K' are the continuous version of kernels K_{ij} and K'_{ij} of (5.7) wherein the subscripts i and j are related to non-primed and primed variables.

To complete the expression (5.11) we observe that

$$\begin{aligned} -\mathbf{r}_1 \cdot \nabla_{\boldsymbol{\xi}} \eta_1|_{\delta=0} &= -\nabla_{\boldsymbol{\xi}}(\Phi_1) \cdot \nabla_{\boldsymbol{\xi}}(\Phi_{1\delta})|_{\delta=0} \\ &= -\iint (\mathbf{k}' \cdot \widehat{\mathbf{k}}) a_k a_{k'} \sin(\psi) \sin(\psi') \\ &= -\iint \frac{1}{2} (\mathbf{k}' \cdot \widehat{\mathbf{k}} + \mathbf{k} \cdot \widehat{\mathbf{k}}) a_k a_{k'} \sin(\psi) \sin(\psi'). \end{aligned} \quad (5.14)$$

The combination of (5.13) and (5.14) yields

$$\eta_2 = \frac{1}{2} \iint a_k a_{k'} [K \cos(\psi) \cos(\psi') + K' \sin(\psi) \sin(\psi')], \quad (5.15)$$

which is the continuous version of (5.6) derived by Longuet-Higgins (1963).

5.2. Consistency with the Lagrangian derivation of W. J. Pierson

In 1961 W. J. Pierson derived a Lagrangian second-order solution to the discrete long-crested problem. He considered waves travelling in the positive α direction only and found the solutions in the forms (equations (27) and (28) in Pierson (1961)):

$$\begin{aligned} x(\alpha, \delta, t) &= \alpha - \sum_i a_i e^{k_i \delta} \sin(\psi_i) - \sum_{j>i} \sum_i \frac{a_i a_j}{g} \left(\frac{\omega_i^3 + \omega_j^3}{\omega_j - \omega_i} \right) e^{(k_j+k_i)\delta} \sin(\psi_j - \psi_i) \\ &\quad + \sum_{j>i} \sum_i \frac{a_i a_j}{g} (\omega_j + \omega_i) \omega_j e^{(k_j-k_i)\delta} \sin(\psi_j - \psi_i) + \sum_i a_i^2 \omega_i k_i e^{2k_i \delta} t \end{aligned} \quad (5.16)$$

$$\begin{aligned} z(\alpha, \delta, t) &= \delta + \sum_i a_i e^{k_i \delta} \cos(\psi_i) + \sum_{j>i} \sum_i \frac{a_i a_j}{g} (\omega_i^2 + \omega_i \omega_j + \omega_j^2) e^{(k_j+k_i)\delta} \\ &\quad \times \cos(\psi_j - \psi_i) - \sum_{j>i} \sum_i \frac{a_i a_j}{g} (\omega_j + \omega_i) \omega_j e^{(k_j-k_i)\delta} \cos(\psi_j - \psi_i) \end{aligned} \quad (5.17)$$

$$\begin{aligned} p(\alpha, \delta, t) &= p_a - \rho g \delta + \rho g \sum_i \frac{a_i^2 k_i}{2} (e^{2k_i \delta} - 1) - 2\rho \sum_{j>i} \sum_i a_i a_j \omega_i \omega_j e^{(k_j+k_i)\delta} \\ &\quad \times \cos(\psi_j - \psi_i) + 2\rho \sum_{j>i} \sum_i a_i a_j \omega_i \omega_j e^{(k_j-k_i)\delta} \cos(\psi_j - \psi_i) \end{aligned} \quad (5.18)$$

with $\psi_i = k_i \alpha - \omega_i t + \varphi_i$.

The comparison of our continuous solution with the discrete formulation of Pierson is not straightforward due to the PV formulation of one of the terms. However, it

should be noted that within a small subspace \mathcal{D} of \mathbb{R}^4 around the singularity domain ($|\omega - \omega'| < \varepsilon$) we have

$$PV \iint_{\mathcal{D}} i \left(\frac{\omega \mathbf{k} + \omega' \mathbf{k}'}{2(\omega - \omega')} e^{(k+k')\delta} e^{-i(\omega - \omega')t} \right) \underline{\mathcal{B}} + \text{c.c.} \simeq \iint_{\mathcal{D}} \frac{1}{2} \omega k (\widehat{\mathbf{k}} + \widehat{\mathbf{k}'}) e^{(k+k')\delta} \underline{\mathcal{B}} t. \tag{5.19}$$

This result corresponds to a temporal secular term. More detailed comments on this term can be found in § 6.1.

Moreover, we can note that (see (4.40)):

$$\underline{\mathcal{R}'} \xrightarrow[k \rightarrow k']{} k (\widehat{\mathbf{k}} - \widehat{\mathbf{k}'}). \tag{5.20}$$

Within a small subspace \mathcal{D}' of \mathbb{R}^4 defined by $\|\mathbf{k} - \mathbf{k}'\| < \varepsilon$ and due to the symmetry of the previous limit we have

$$\iint_{\mathcal{D}'} i \frac{\sqrt{kk'}(\mathbf{k} - \mathbf{k}')}{\Omega^- - \|\mathbf{k} - \mathbf{k}'\|} e^{\|\mathbf{k} - \mathbf{k}'\|\delta} \underline{\mathcal{B}} e^{-i(\omega - \omega')t} + \text{c.c.} \xrightarrow[\varepsilon \rightarrow 0]{} 0 \tag{5.21}$$

since integration is realized over all \mathbf{k} and \mathbf{k}' .

Restricting solution (4.47) to the discrete case of long-crested waves travelling in the same positive α direction ($\widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}'} = 1$, $\|\mathbf{k} - \mathbf{k}'\| = s(k - k')$ where s is the sign of $k - k'$) we obtain the following expressions for the second-order displacements and pressure:

$$x_2 = - \sum_{\substack{i,j \\ i \neq j}} \left[\frac{\omega_i k_i + \omega_j k_j}{2(\omega_i - \omega_j)} e^{(k_i+k_j)\delta} + \frac{\sqrt{k_i k_j} (k_i - k_j)}{\Omega_{ij}^- - s(k_i - k_j)} e^{s(k_i - k_j)\delta} \right] a_i a_j \sin(\psi_i - \psi_j) + \sum_i a_i^2 \omega_i k_i e^{2k_i \delta} t \tag{5.22}$$

$$z_2 = \sum_{i,j} \left[\frac{k_i + k_j + \Omega_{ij}^+}{4} e^{(k_i+k_j)\delta} + \frac{\sqrt{k_i k_j} s (k_i - k_j)}{\Omega_{ij}^- - s(k_i - k_j)} e^{s(k_i - k_j)\delta} \right] a_i a_j \cos(\psi_i - \psi_j) \tag{5.23}$$

$$p_2 = -\rho g \sum_{i,j} \sqrt{k_i k_j} (e^{(k_i+k_j)\delta} - e^{s(k_i - k_j)\delta}) a_i a_j \cos(\psi_i - \psi_j) \tag{5.24}$$

where non-primed and primed variables of (4.47) are related to the subscripts i and j , respectively. Making use of the dispersion relationship $\omega^2 = gk$ we can rewrite after straightforward manipulations

$$x_2 = - \sum_{j>i} \sum_i \frac{a_i a_j}{g} \left(\frac{\omega_i^3 + \omega_j^3}{\omega_j - \omega_i} \right) e^{(k_j+k_i)\delta} \sin(\psi_j - \psi_i) + \sum_{j>i} \sum_i \frac{a_i a_j}{g} (\omega_j + \omega_i) \omega_j e^{(k_j-k_i)\delta} \sin(\psi_j - \psi_i) + \sum_i a_i^2 \omega_i k_i e^{2k_i \delta} t \tag{5.25}$$

$$z_2 = \sum_{j>i} \sum_i \frac{a_i a_j}{g} (\omega_i^2 + \omega_i \omega_j + \omega_j^2) e^{(k_j+k_i)\delta} \cos(\psi_j - \psi_i) - \sum_{j>i} \sum_i \frac{a_i a_j}{g} (\omega_j + \omega_i) \omega_j e^{(k_j-k_i)\delta} \cos(\psi_j - \psi_i) + \sum_i \frac{1}{2} a_i^2 k_i e^{2k_i \delta} \tag{5.26}$$

$$\begin{aligned}
 p_2 = & \rho g \sum_i \frac{a_i^2 k_i}{2} (e^{2k_i \delta} - 1) - 2\rho \sum_{j>i} \sum_i a_i a_j \omega_i \omega_j e^{(k_j+k_i)\delta} \cos(\psi_j - \psi_i) \\
 & + 2\rho \sum_{j>i} \sum_i a_i a_j \omega_i \omega_j e^{(k_j-k_i)\delta} \cos(\psi_j - \psi_i)
 \end{aligned} \tag{5.27}$$

which differs from the original derivation of Pierson (1961) (5.16)–(5.18) by the constant term $\sum_i (a_i^2 k_i e^{2k_i \delta})/2$ in the vertical displacement corresponding to the mean of z_2 . A closer inspection of Pierson’s original derivation shows that he used $\partial|\mathbb{J}|/\partial t=0$ as its basic continuity equation. However, this does not necessarily imply that $|\mathbb{J}|=1$ and can lead to erroneous solutions. Using equation $\partial|\mathbb{J}|/\partial t=0$ instead of $|\mathbb{J}|=1$ allows the cancellation of all of the time-independent terms in the solutions. This is the reason why the mean level of z_2 is absent in the derivations of Pierson (1961) which must be rectified as (5.25)–(5.27).

6. Analysis of the second-order solution

6.1. Stokes drift

We will now investigate some remarkable properties of the second-order Lagrangian solution. The first is the customary Stokes drift, first introduced in the celebrated work by Stokes (1847) and extended to the tri-dimensional case by Kenyon (1969) and Phillips (1977). The Stokes drift manifests itself in a net horizontal displacement after one wave period or, more generally, after time averaging. The net mass transport can be evaluated using the horizontal velocity \mathbf{r}_{2t} estimation. As shown in the derivation below, only the third integral term \mathbf{r}_2 (see (4.42)) in the expression of \mathbf{r}_2 has a non-vanishing temporal mean. Equation (6.1) gives the horizontal velocity for this term only (note that the apparent singularity disappears after differentiation):

$$\mathbf{r}_{2t} = \iint \frac{1}{2} (\omega \mathbf{k} + \omega' \mathbf{k}') e^{(k+k')\delta} \underline{\mathcal{B}} e^{-i(\omega-\omega')t} + \text{c.c.} \tag{6.1}$$

We now consider the time average of this quantity:

$$\langle \mathbf{r}_{2t} \rangle_t = \lim_{T \rightarrow \infty} \frac{1}{T} \int_{-T/2}^{T/2} dt \iint_{\mathbb{R}^4} \frac{1}{2} (\omega \mathbf{k} + \omega' \mathbf{k}') \underline{\mathcal{B}} e^{-i(\omega-\omega')t} e^{(k+k')\delta} + \text{c.c.} \tag{6.2}$$

Inverting time and space integrals and using

$$\lim_{T \rightarrow \infty} \frac{1}{T} \int_{-T/2}^{T/2} \cos[(\omega - \omega')t] dt = \delta(\omega - \omega') \tag{6.3}$$

where δ is the Dirac distribution, we obtain

$$\langle \mathbf{r}_{2t} \rangle_t = \iint_{\mathbb{R}^4} \delta(\omega - \omega') \frac{1}{2} \omega k (\hat{\mathbf{k}} + \hat{\mathbf{k}}') \underline{\mathcal{B}} e^{2k\delta} + \text{c.c.} \tag{6.4}$$

All other terms in \mathbf{r}_{2t} have a vanishing temporal mean due to their $\omega - \omega'$ dependency which appears after temporal differentiation and due to the Dirac function. This is why $\underline{\mathbf{r}}_{2t}$ is replaced by \mathbf{r}_{2t} in (6.4). Equation (6.4) is thus the total mean average of

the horizontal displacement of particles. Using again (6.3) in the space domain, we derive the spatial mean of (6.4) which can be written

$$\langle \mathbf{r}_{2t} \rangle_{\xi t} = \iint_{\mathbb{R}^4} \bar{\delta}(\omega - \omega') \bar{\delta}(\hat{\mathbf{k}} - \hat{\mathbf{k}}') \frac{1}{2} \omega k (\hat{\mathbf{k}} + \hat{\mathbf{k}}') \frac{1}{4} (1 + \hat{\mathbf{k}} \cdot \hat{\mathbf{k}}') A(\mathbf{k}) A^*(\mathbf{k}') e^{2k\delta} + \text{c.c.} \quad (6.5)$$

Simplified as

$$\langle \mathbf{r}_{2t} \rangle_{\xi t} = \int \omega k \|A(\mathbf{k})\|^2 e^{2k\delta} \quad (6.6)$$

the classical Stokes drift velocity is easily identified. The mean Stokes drift $\langle \mathbf{r}_2 \rangle_{\xi t}$ is thus already included as a part of the \mathbf{r}_2 expression (4.47) and is the results of the self-interaction of the different harmonics. Clamond (2007) derived this result for a monochromatic wave and noted that after subtraction of this mass transport component, the orbits of water particles remained closed and symmetric even for steep waves (see also Longuet-Higgins 1987). As noted before in § 5.1, equation (5.11), the contribution of \mathbf{r}_2 is absent in the Eulerian expansion, leading to the absence of the Stokes drift in the second-order Eulerian expansion.

6.2. Distortion of wave fronts

In the case of tri-dimensional multiple wave interactions, a residual spatial Stokes drift pattern, namely $\langle \mathbf{r}_{2t} \rangle_t - \langle \mathbf{r}_{2t} \rangle_{\xi t}$, remains. It results from the interaction of harmonics of equal time frequency but having different propagation directions. This phenomena which cannot exist in the bi-dimensional case (because $\omega = \omega'$ implies $k = k'$) is responsible for the increase of the wave shape asymmetry over time. An example is shown in figure 1. Two harmonics of equal frequency but propagating in different directions create a spatially varying shear over the sea surface (figure 1*b*). This shear tends to slow down the troughs relatively to the crests leading to an asymmetric wave shape that can be related to the first stage of the formation of the well-known horseshoe patterns (see figure 1*c*). The front-back symmetry of the waves and the absence of slope skewness are nonetheless preserved.

Shrira, Badulin & Kharif (1996) and later Annenkov & Shrira (1999) proposed a mathematical solution to explain the apparition and the persistence of the horseshoe pattern by quintet resonant interactions coupled with wind and dissipation and noted that waves developed front-back asymmetries. The two main characteristics of the horseshoe patterns are (a) a life time largely exceeding the associated wave period and (b) a persistent shape with front-back asymmetry.

For clarity, in the explanations below, the term ‘harmonic’ is used for a Lagrangian wave vector component and the term ‘wave’ is used for an Eulerian (surface) wave vector component. Even though cross-comparison of harmonic interactions and wave interactions is not an easy task, we can try to estimate which harmonics are involved in the development of such wave front deformations. Bi-harmonic interaction terms are present in both horizontal (x_2, y_2) and vertical (z_2) second-order Lagrangian displacements. Therefore, simply using bi-harmonic Lagrangian interactions should *a priori* make it possible to obtain, at least partially, the interactions of a wave quartet in the Eulerian framework. Horseshoe patterns observed by Collard & Caulliez (1999) present peculiar features that can be compared with the model presented. Their experiment starts from an almost monochromatic wave with wavenumber \mathbf{k}_0 . The wave field later degenerates and gives rise to crescent-shape patterns. The spatiotemporal

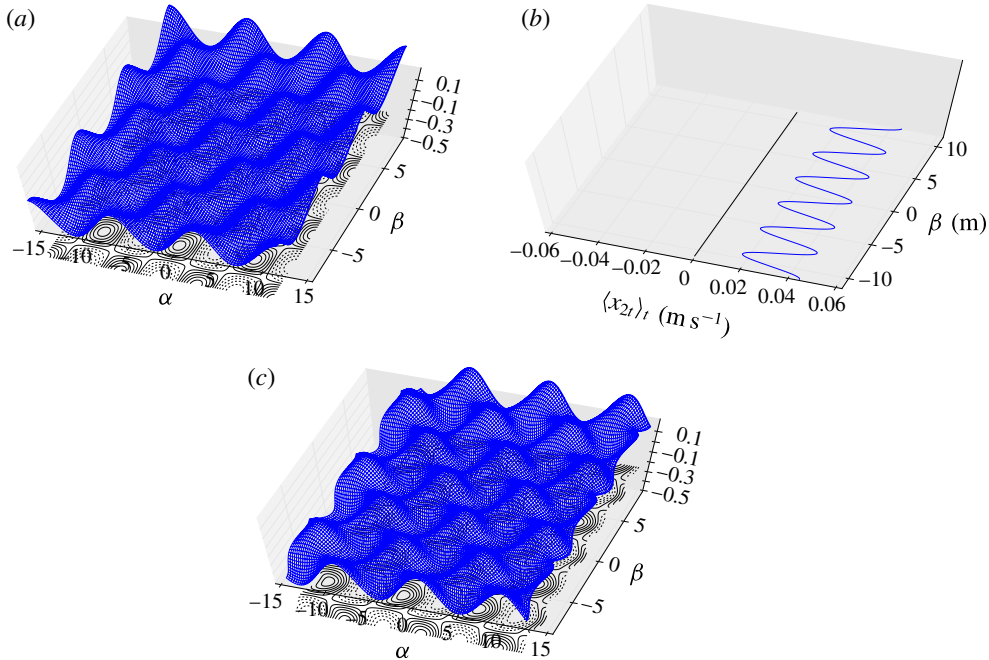


FIGURE 1. (Colour online) Interaction of two harmonics of equal amplitudes ($A = 0.08$ m) and wavenumbers ($k = 1.104$ rad m^{-1}) but different directions of propagation: $+$ and -48.2° relative to direction α . (a) Sea surface elevation at $t = 0$. (b) α component of the Stokes drift velocity: $\langle r_{2t} \rangle_t$ (β component vanishes). Such component is independent of α . We can see a mean Stokes drift of about 3 cm s^{-1} and its spatial variations leading to increasing crescent shape patterns. (c) Sea surface elevation after a long time period of 30 s. The surface has developed distorted wave fronts.

analysis of this experiment shows that a pair of harmonics ($\mathbf{k}_1, \mathbf{k}_2$) are created so that:

$$\mathbf{k}_1 + \mathbf{k}_2 = 3\mathbf{k}_0 \quad \text{and} \quad \omega_1 + \omega_2 = 3\omega_0. \quad (6.7a,b)$$

The case called ‘steady pattern’ in Collard & Caulliez (1999) is defined with $k_1 = k_2$ and consequently $\omega_1 = \omega_2$. In this specific case, the results presented in figure 1 suggest that no \mathbf{k}_0 component is necessary to obtain wave front deformation since the secular term in $\langle r_{2t} \rangle_t - \langle r_{2t} \rangle_{\xi t}$ is generated by the \mathbf{k}_1 and \mathbf{k}_2 component interaction only. However, since the \mathbf{k}_0 component is obviously present in the experiment of Collard & Caulliez (1999), it was not experimentally possible to check whether such component is indispensable to the wave front deformation. Yet, it could be said that \mathbf{k}_0 is indirectly necessary to the emergence of the perfectly symmetric pair of wave vectors ($\mathbf{k}_1, \mathbf{k}_2$) through the resonant interaction defined by (6.7). It should therefore be interesting to experimentally show that a unique bi-harmonic structure such as that presented in figure 1 is sufficient to create horseshoe patterns. However, in order to obtain such wave shapes, our methodology only performs satisfactorily when very low steepness and very long time periods are considered. Indeed, first- and second-order steady nonlinearities should remain low until the secular second-order term becomes observable. Experiments made by Kimmoun, Branger & Kharif (1999) show that beyond a certain level of steepness, no wave front asymmetry is observed and that

steady first- and second-order nonlinearities, as described in §6.3, become the main contributors to the wave shape deformation.

However, in the simulation presented in figure 1, the requirement of low steepness of the \mathbf{k}_1 and \mathbf{k}_2 component is fulfilled and, for comparison purposes with Collard & Caulliez (1999), we added a \mathbf{k}_0 component in the orbital spectrum which was chosen so that $\mathbf{k}_1 + \mathbf{k}_2 = 3\mathbf{k}_0$, $\varphi_0 = -\pi/2$ and $\varphi_1 = \varphi_2 = 0$ (but obviously where $\omega_1 + \omega_2 \neq 3\omega_0$). This reproduced the ‘steady’ horseshoe patterns presented in Collard & Caulliez (1999). Figure 2(a) shows the wave field obtained after 30 s and figure 2(b,c) show the temporal record of the water surface elevation and its Fourier transform. As pointed out by Collard & Caulliez (1999), such wave field contains a $(3/2)\omega_0$ harmonic (0.524 Hz) emanating from the simple first-order contribution of \mathbf{k}_1 and \mathbf{k}_2 . We believe that the front–back asymmetry of the horseshoe pattern observed in real conditions comes from higher-order interactions (Lagrangian cubic order) that would create a \mathbf{k}_0 component with an angular frequency of ω_0 but with a slightly different phase.

In any case, even if the spatial drift we found tended to slowly twist the wave shape and made it tend towards the horseshoe pattern, this drift led to a constant increase of the surface deformation over time giving rise to unrealistic shapes after a long time period. Moreover, as already noted by Shrira *et al.* (1996), steady wave solutions of inviscid equations do not present front–back asymmetries. Hence, the secular term can only belong to a transitory state of the surface and cannot be used for long time periods as suggested by the domain of validity of the series expansion.

As already mentioned, the Stokes drift manifests itself through a secular term which is undesirable in a perturbation expansion. Indeed, as commented by Buldakov *et al.* (2006), third-order solutions will make the secular term interact with the leading order creating unrealistic diverging secular terms in both horizontal and vertical particles expansion. As a result, the second-order solution cannot be valid at arbitrary long time periods. Furthermore, any attempt to pursue the Lagrangian expansion beyond the second order should be accompanied with a particles relabelling as suggested by Clamond (2007) who claimed that a steady solution with Stokes drift cannot be found without adapting the Lagrangian references. The apparition of the mean secular term in a Lagrangian expansion comes from a misrepresentation of steady waves and can be avoided, at least in the case of a monochromatic wave, by a correct time-and-space-dependent water particles relabelling leading to a valid solution at all time and orders. However, as this paper deals with tri-dimensional multiple wave system and is restricted to the second-order expansion, we shall not enter into such details and leave these considerations for further studies.

6.3. Sharp crests, mean elevation and skewness coefficient

For the temporal mean of the second-order vertical displacement at the surface we write

$$\langle z_2 \rangle_{t,\delta=0} = \iint_{\mathbb{R}^4} \bar{\partial}(\omega - \omega') \frac{1}{2} k \underline{\mathcal{B}} + \text{c.c.} \quad (6.8)$$

This shows that the second-order vertical displacement has a non-vanishing mean due to the interaction of waves of equal frequency. The $(\hat{\mathbf{k}} - \hat{\mathbf{k}}')$ phase term in $\underline{\mathcal{B}}$ describes a spatial oscillating pattern perpendicular to the mean direction of the waves and is the main contributor to the vertical second-order displacement. To illustrate this statement, we use the same bi-harmonic system as described above. Figure 3 displays first- and second-order surface slices along an equi- α contour corresponding to a crest

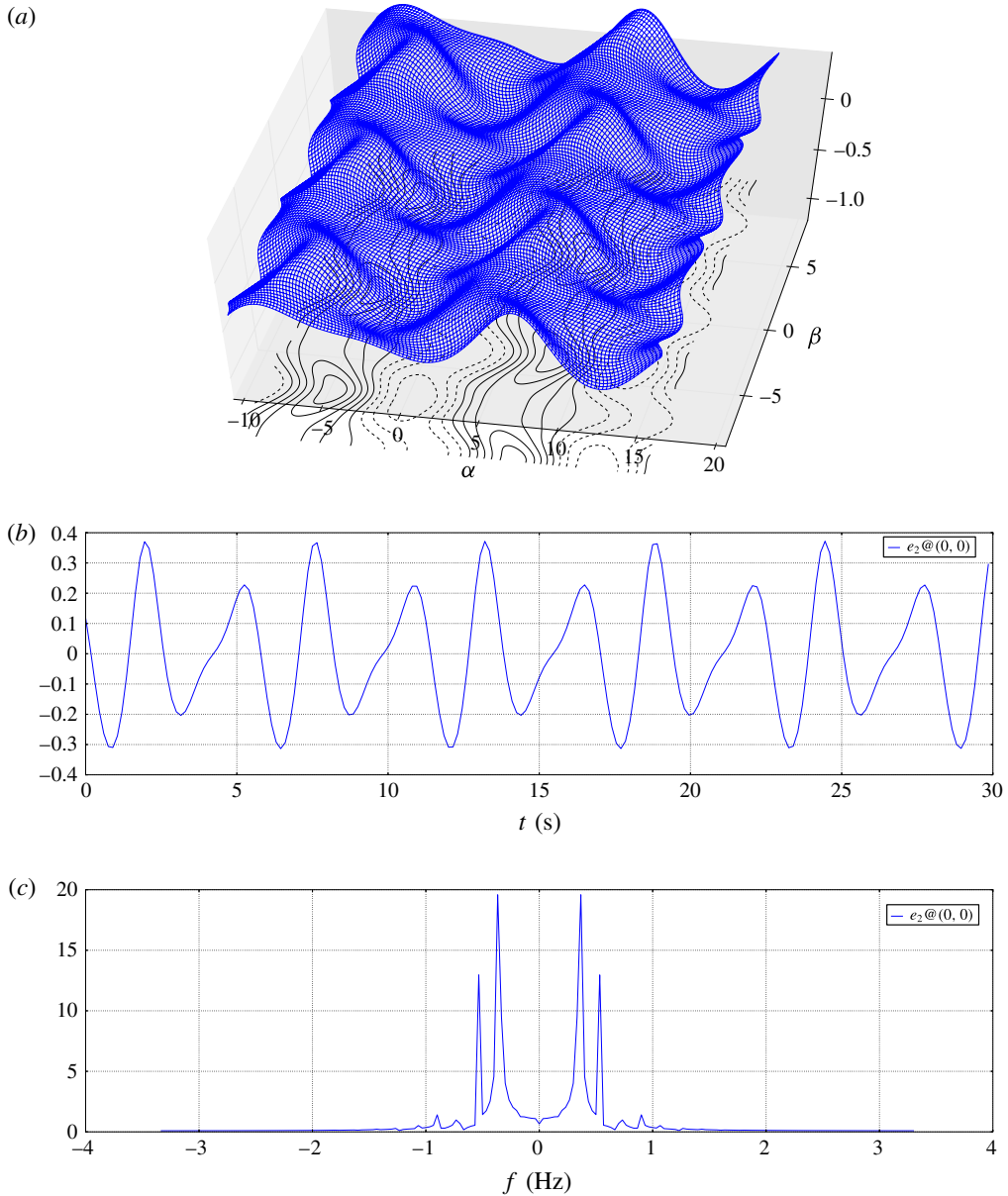


FIGURE 2. (Colour online) Surface elevation obtained with three harmonics, $k_0 = 0.491 \text{ rad m}^{-1}$, $k_1 = k_2 = 1.104 \text{ rad m}^{-1}$, respective amplitudes, $a_0 = 0.24 \text{ m}$ and $a_1 = a_2 = 0.06 \text{ m}$, directions of propagation relative to α , $\theta_0 = 0$, $\theta_1 = -\theta_2 = \arccos(2/3)$, and phases $\varphi_0 = -\pi/2$, $\varphi_1 = \varphi_2 = 0$. (a) Sea surface elevation at $t = 30 \text{ s}$. (b) Temporal record of the water surface elevation at fixed point $(\alpha, \beta) = (0, 0)$. (c) Fourier transform of the temporal record (figure 2b). We can identify frequency $f_0 = (2\pi)^{-1}\sqrt{gk_0}$ and harmonic $(3/2)f_0$ coming for k_1 and k_2 contribution.

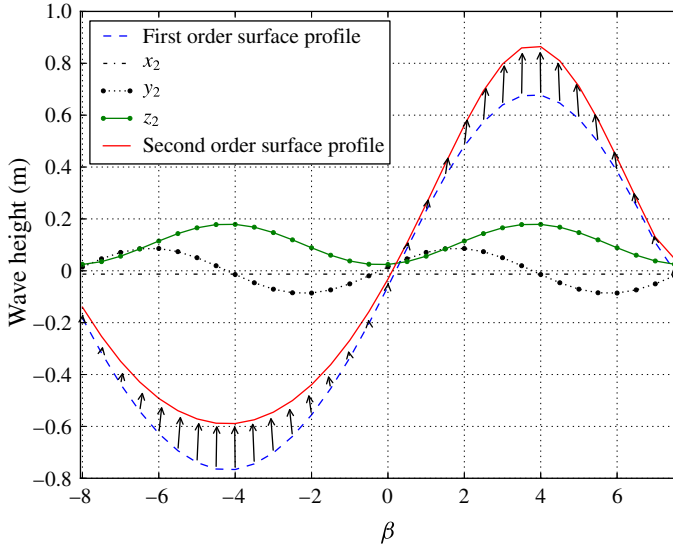


FIGURE 3. (Colour online) Slices along an equi- α contour (perpendicular to the mean direction of wave) of the first- and second-order Lagrangian surface. The slices pass through a wave crest. The different second-order contributions are superimposed as well as their combined effects (arrows).

position. As can be plainly seen, the second-order vertical term tends to permanently sharpen the crests and flatten the troughs by a positive vertical shift relative to the z_2 mean level. Contrarily to the Lagrangian first-order terms, the sharpening and flattening effects apply in a direction perpendicular to the wave direction leading to a more ‘short-crested’ wave pattern. The horizontal term y_2 has the same effect even though it is in quadrature with the vertical motion. The combination of both effects is represented using arrows on figure 3.

Kimmoun *et al.* (1999) used an advanced methodology to derive the surface topography from wave tank experiments. In their sixth experiments, two waves with equal wavelength and different directions interact and develop a short-crested wave field that is analysed and compared with their theoretical calculations. The second wave is obtained by reflexion of the first on a vertical wall. They pushed their theoretical Eulerian calculations up to the third order in the wave steepness parameter to compare with their observations. The sixth experiment (see figure 4a) clearly shows ‘rhombic form of the crest and the elliptic form of the troughs’ that also appear on the second-order Lagrangian simulation plotted on figure 4(b). First-order Eulerian and Lagrangian simulations are respectively plotted in figure 4(c,d).

It was shown by Pierson (1961, equation (45)) that the first-order Lagrangian surface has a relative mean level:

$$\bar{\eta}_1 = - \int k \|A(\mathbf{k})\|^2. \tag{6.9}$$

The mean sea level is affected by the non-vanishing mean of the second-order elevation term (see mean level of z_2 on figure 3). At the surface,

$$\langle z_2 \rangle_{\xi,t} = \int \frac{1}{2} k \|A(\mathbf{k})\|^2 \tag{6.10}$$

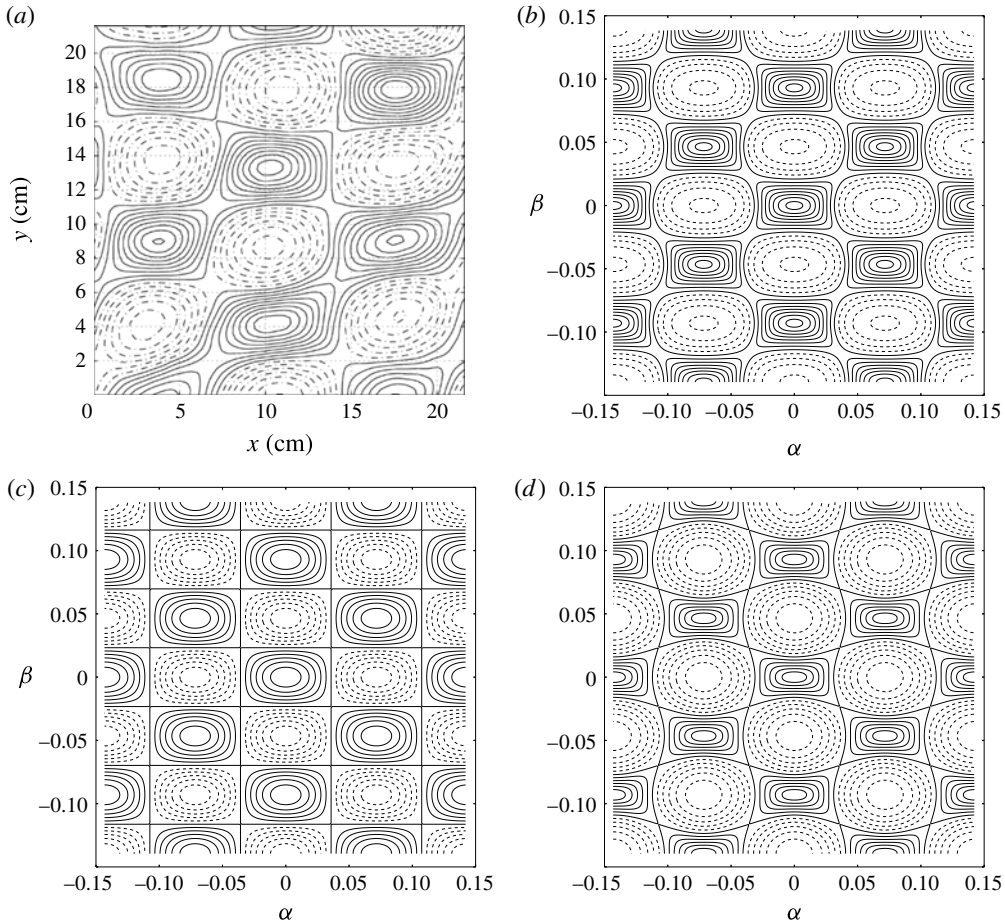


FIGURE 4. Interaction of two harmonics of equal amplitudes ($A = 0.43$ cm) and wavelength ($\lambda = 7.8$ cm) but different directions of propagation: $+$ and -33° relative to direction β : (a) figure 14 from Kimmoun *et al.* (1999); (b) second-order Lagrangian profile; (c) first-order Eulerian profile; (d) first-order Lagrangian profile.

is the unique second-order contributor to mean surface elevation at the leading order giving the overall sea level:

$$\bar{\eta}_2 = - \int \frac{1}{2} k \|A(\mathbf{k})\|^2. \quad (6.11)$$

When this mean level is naturally tared by a Lagrangian sensor (free-floating buoys, etc.), it results in a mean sea level greater by an amount of $|\bar{\eta}_2|$ than an Eulerian measuring system (fixed probes, etc.) does. This conclusion has already been reached by Longuet-Higgins (1986) (see equation (3.7)) via a different route. The author emphasized the importance of this effect in particular with respect to ocean surface remote sensing applications. However, as already stated by Longuet-Higgins (1987), Lagrangian orbits are highly symmetrical at the second-order leading to a vanishing skewness. In random ocean wave fields, such second-order dynamical effects have strong impacts on waves height, slope and curvature distributions and are responsible for their deviation from the Gaussian law. These statistical properties are of great

interest in the ocean remote-sensing community but a systematic study goes beyond the scope of this paper and is left for further developments.

6.4. Modulational Benjamin–Feir instability, a simple beat effect

It is now well known that Benjamin–Feir (BF) instability results from a nonlinear quartet-wave resonant phenomenon (Benjamin & Feir 1967). An initial uniform monochromatic Stokes wave of moderate amplitude develops side-band harmonics with an exponential rate of growth and degenerates into a sequence of wave packets.

In this section we do not wish to enter into a complex analytical analysis but would like to show, on the basis of theoretical and numerical considerations, that the periodic regime of BF instability is already present (at least partly) and symmetric (i.e. with no frequency downshift) in the Lagrangian second-order solution.

We shall not study the growth period since we only consider periodic solutions but we shall show that, at the second order of the nonlinear Lagrangian parameter, a periodic modulation process exists between the carrier and two existing sideband harmonics and can be interpreted as a Lagrangian Benjamin–Feir modulation. We shall show that, surprisingly, what is considered to be a periodical exchange of energy between waves from a Eulerian point of view is in fact a simple beat effect which appears naturally when a two-wave system has close frequencies in the Lagrangian framework. The same initial sea state is used in the Eulerian framework and shows that this phenomenon is clearly absent up to the second order.

6.4.1. System of three aligned harmonics

In order to illustrate this statement, let us consider a bi-dimensional and uni-directional case defined up to the second order by (5.16) and (5.17). We focus on a bi-harmonics system defined by its wavenumbers k_0 and k_2 ($0 < k_0 < k_2$) where k_0 is the carrier wavenumber, k_2 the satellite wavenumber and φ_0 and φ_2 their respective phases. The carrier wave is chosen with wavenumber $k_0 = \pi/2$ rad m⁻¹ propagating in the α direction and corresponding to a 4 m wavelength and a 1.6 s time period. Its orbital amplitude $a_0 = 0.2228$ m is chosen so that $s = k_0 a_0 = 0.35$. It must however be emphasized that a_0 is the orbital spectral amplitude and that the real amplitude of the carrier never exceed 0.2 m leading to a maximum steepness of 0.3. The satellite wavenumber is $k_2 = k_0 \times (1 + p)$ with $p = 0.1$. Its orbital amplitude is $a_2 = a_0 \times c$ with $c = 0.08$.

We generate a surface of 90 m length with a 12.5 cm label spatial sampling over a thousand periods of the carrier wave and evaluate, at each step in time, the spectral amplitude of the surface $\eta(\alpha, t)$ (a numerical interpolation of the surface profile on a regular grid was realized prior to its Fourier transform), the horizontal $x(\alpha, t)$ and vertical $z(\alpha, t)$ particle displacement processes at the surface $\delta = 0$. We therefore evaluate the Fourier transforms $\hat{\eta}(k, t)$, $\hat{x}(k, t)$ and $\hat{z}(k, t)$ defined by

$$\hat{\Psi}(k, t) = \int (\Psi(\alpha, t) - \bar{\Psi}) e^{ik\alpha} d\alpha, \quad (6.12)$$

where Ψ stands for any of the three quantities η , x or z and where the upper line $\bar{\Psi}$ refers to the spatial average. These quantities are constant in time and are given in table 1 together with the orbital spectral amplitudes $|A(k)|$. We have selected the wavenumbers associated to non-vanishing amplitudes. All harmonics of the orbital spectrum are aligned and produce a unique temporal secular term corresponding to a

k	$k_2 - k_0$	$2k_0 - k_2$	k_0	k_2	$2k_2 - k_0$
$ A(k) $	0	0	0.2228	0.0178	0
$ \hat{x}(k) $	0.2785	0	0.2228	0.0178	0
$ \hat{z}(k) $	0.0062	0	0.2228	0.0178	0
$ \hat{\eta}(k) $	0.0018	0.04627	0.1979	0.0598	0.0093

TABLE 1. Orbital $|A(k)|$, horizontal $|\hat{x}|$, vertical $|\hat{z}|$ motions and surface $|\hat{\eta}|$ spectral amplitudes obtained from a two-wave orbital system: $(k_0, k_2) = \pi/2 \times (1, 1 + p)$ and $(a_0, a_2) = s/k_0 \times (1, c)$ with $s = 0.35$, $p = 0.1$ and $c = 0.08$.

global horizontal translation of the sea surface profile. This main constant drift can easily be removed by adapting the frame of reference ensuring the validity of the second-order expansion, in this case only, even for long time periods.

Even though neither the vertical $|\hat{z}|$ nor horizontal $|\hat{x}|$ displacement spectra contain a $2k_0 - k_2$ component, this component is present in the surface spectrum. The $k_2 - k_0$ component is an important contributor to the second-order horizontal displacement:

$$-\frac{a_0 a_2}{g} \left(\frac{\omega_0^3 + \omega_2^3}{\omega_2 - \omega_0} \right) \sin [(k_2 - k_0)x - (\omega_2 - \omega_0)t + \varphi_2 - \varphi_0]. \quad (6.13)$$

Table 1 clearly shows the $k_0 - (k_2 - k_0) = 2k_0 - k_2$ harmonic in the surface spectrum due to the combination of the horizontal $k_2 - k_0$ term and the k_0 term. We can therefore easily deduce that the angular frequency and phase of this term become $2\omega_0 - \omega_2$ and $2\varphi_0 - \varphi_2$. The other interaction term with wavenumber $k_0 + (k_2 - k_0) = k_2$ has the same frequency ω_2 and phase φ_2 as the orbital first-order k_2 component of the orbital spectrum and simply affects its amplitude.

The case presented in table 1 shows that the k_2 component of the surface spectrum benefits from a constructive interaction of the k_0 and $(k_2 - k_0)$ terms since its amplitude (0.0598) is increased relatively to the specified orbital amplitude (0.0178). Conversely, the amplitude of the k_0 component is decreased relatively to orbital amplitude due to a destructive interaction. As expected, the surface spectrum also contains a very small $2k_2 - k_0$ term arising from the combination term $k_2 + (k_2 - k_0)$.

Now, let us suppose that an extra k_1 component is added to the orbital spectrum in such a way that $k_1 = 2k_0 - k_2$ as shown in figure 5 and denote ω_1 and φ_1 as the associated angular frequency and phase. This component will therefore have the same spatial wavenumber as the $k_0 - (k_2 - k_0)$ term presented above but with a slightly different temporal frequency. These two terms will therefore generate a temporal beat effect with angular frequency $\Delta\omega$ such as

$$\Delta\omega = \omega_1 - (2\omega_0 - \omega_2) \quad (6.14)$$

and thus the phase of k_1 amplitude temporal evolution will only depend on the global phase:

$$\theta = \varphi_1 - (2\varphi_0 - \varphi_2). \quad (6.15)$$

Inverting k_1 and k_2 in the previous considerations we obtain the same behaviour for the k_2 component. Now, letting the triple-harmonic structure system evolve in time leads to a periodic evolution of the two side-band harmonic amplitudes which share the same time period:

$$T = \frac{2\pi}{\Delta\omega} \quad (6.16)$$

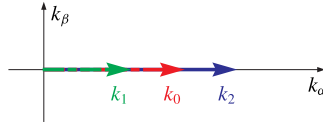


FIGURE 5. (Colour online) Spectral repartition of the orbital spectrum components with $k_1 + k_2 = 2k_0$.

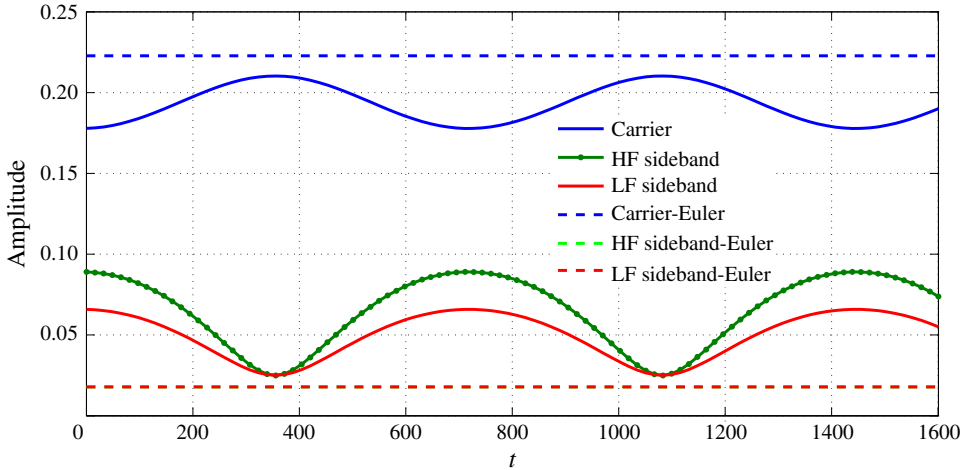


FIGURE 6. (Colour online) Time evolution of the spectral amplitudes of the carrier wave ($\omega_0 = 3.93 \text{ rad s}^{-1}$), the low- and the high-frequency sidebands ($\omega_1 = 3.74$ and $\omega_2 = 4.11 \text{ rad s}^{-1}$). Solid lines represent the Lagrangian expansion and dashed lines the Eulerian expansion. The period defined in (6.16) is 726 s.

and the same evolution phase depending on the unique value θ . The time evolution of the carrier, high-frequency (HF) and low-frequency (LF) side-band amplitudes is presented in figure 6. The corresponding Eulerian case is presented for comparison purposes and clearly shows that the BF modulation is absent up to the second order. Shemer (2010) had already derived these two results using a different technique in the Eulerian framework by pushing the nonlinearity at the third order and considering the evolution of a wave quartet.

It should be noted that the mean level and the variations of each satellite amplitude are not fully controlled by the ratio $c = a_2/a_0$ and depend on the carrier characteristics and on the other satellite. This makes the quantitative comparison between the two approaches complicated. However, it does not change the conclusion that a strong modulation–demodulation of the carrier wave and of the two satellites is present at the second order in the Lagrangian framework while the Eulerian point of view does not show any interaction even if the sea surface spectrum presents new harmonics relative to the first order.

6.4.2. Carrier harmonic with two lateral side-band harmonics

Let us now consider a symmetric triple-harmonic structure such as $2k_0 = k_1 + k_2$ with $\omega_1 - \omega_0 > 0$ and $\omega_2 - \omega_0 > 0$. This configuration is possible in the tri-dimensional case only and is represented in figure 7. Figure 8 show surface profiles derived using

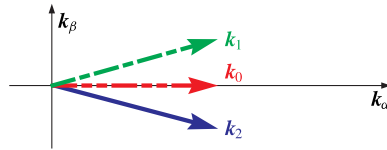


FIGURE 7. (Colour online) Spectral repartition of the orbital spectrum components with $k_1 + k_2 = 2k_0$.

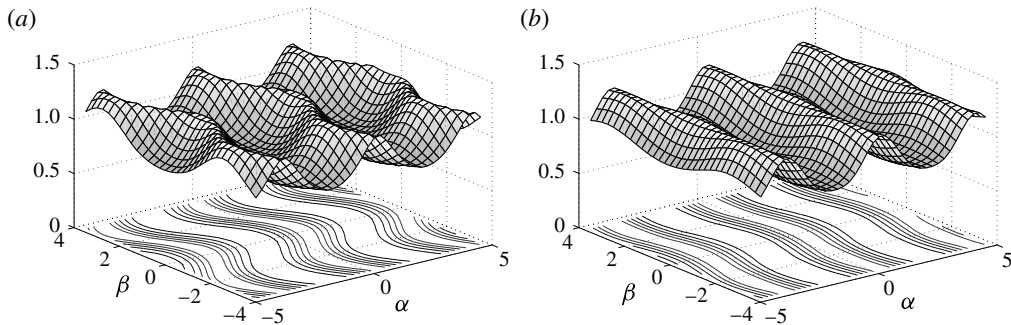


FIGURE 8. Sea surface profiles derived from second-order solutions of Lagrangian (a) and Eulerian (b) expansion. Here $t = 17$ s.

second-order solutions of the Eulerian (Longuet-Higgins 1963) expansion and the Lagrangian (2.7) expansion with (3.7) and (4.47) using the same triple-harmonic structure. The carrier wave is chosen with amplitude $a_0 = 0.2$ m and propagates in the α direction with wavenumber $k_0 = 1.58$ rad m^{-1} corresponding to a 3.97 m wavelength and a 1.58 s time period. Two satellites of equal amplitude $a_1 = a_2 = 0.04$ m with wavenumbers $k_1 = k_2 = k_0 / [\cos(37.08^\circ)]$ rad m^{-1} propagate with angles $+37.08^\circ$ and -37.08° relative to the α direction. The phases of each of the three harmonics are set to zero. Spatial sampling is 25 cm.

We generate an $8 \text{ m} \times 8 \text{ m}$ surface with a 25 cm spatial sampling over 10 periods of the carrier wave and a time evolution process is realized by increasing the time variable. A bi-dimensional spectral analysis of the surfaces is realized at each step in time by fast Fourier transform. Again, a numerical interpolation of the Lagrangian surface on a regular grid is realized prior to the Fourier transform. Figure 9 shows the surfaces spectra obtained at $t = 17$ s showing the three-wave pattern. As expected, the surface in the Lagrangian framework contains more harmonics than in the Eulerian framework due to the multiple possible combinations between horizontal and vertical particle harmonics. Figure 10 shows the time evolution of the three harmonic amplitudes. Again, the Eulerian case is presented for comparison purposes showing that the BF modulation is absent.

In the presented tri-dimensional structure, we can see that sideband harmonics modulations are synchronous leading to strong opposite modulation between carrier and harmonics amplitudes. We can also see that the mean amplitude of the carrier in the Lagrangian framework is always notably smaller than the prescribed value (0.2) which is the consequence of constructive harmonics interactions. This decrease of the carrier amplitude is amplified by the fact that the prescribed amplitude is the orbital spectrum amplitude and not the sea surface spectrum amplitude. On the

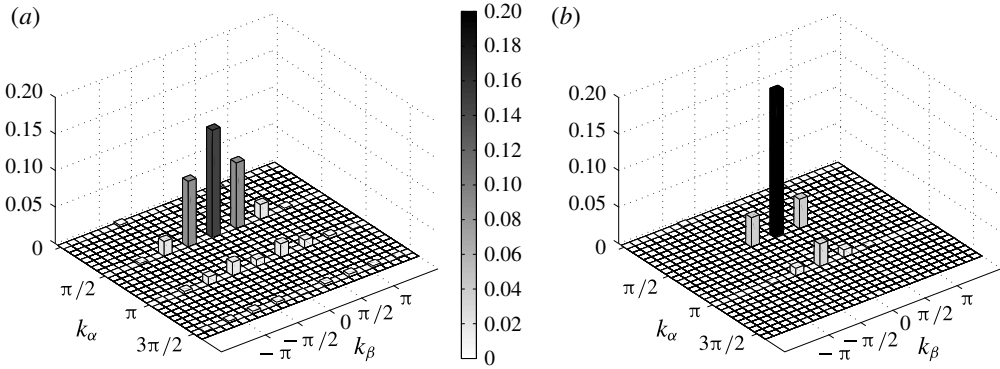


FIGURE 9. Eulerian (a) and Lagrangian (b) surface spectra. Here $t = 17$ s.

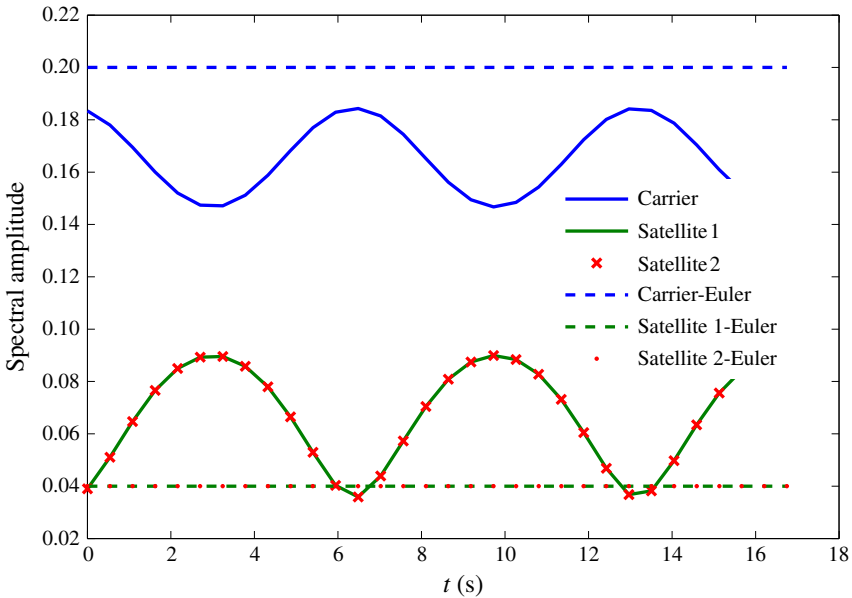


FIGURE 10. (Colour online) Evolution of the carrier and satellites amplitudes as a function of time.

contrary, the two sideband harmonics take advantage of the positive interaction permanently increasing their mean amplitudes. In any case, the observed modulation depth is related to the amplitude of the second-order horizontal term relatively to the first-order component having the same wavenumber. Moreover, we note that the horizontal second-order term is inversely proportional to the difference of the carrier and sideband harmonic frequencies (see (6.13)). Increasing this difference rapidly leads to a strong reduction of the modulation process. Additional numerical simulations with greater frequency differences between the carrier and the sideband harmonics were made and confirm this statement, which is consistent with the proximity of the frequencies observed in BF instability experiments.

Here, we focus on the modulation–demodulation resonance that can be related to a BF modulational instability. We have shown that, from a Lagrangian point of

view, no energy is exchanged between the involved orbital harmonics. Conversely, the Eulerian interpretation of this phenomenon, based on the surface spectrum analysis instead of the orbital analysis, is a permanent and periodical energy exchange between the carrier wave and its two sideband harmonics. This shows that the Lagrangian formulation is in a certain way a more natural and easier point of view. Moreover, in the Lagrangian framework, the time invariability of the harmonics amplitudes suggests that the side-band generation process (the instability itself) can be clearly separated from the modulational part (beating phenomena). However, it is known that asymmetric evolution of the sideband harmonics, responsible for the frequency down-shift effect, is obtained when the modulation increases and when stronger nonlinear effects or dissipation are taken into account. These phenomena are clearly absent at the Lagrangian second-order and will be considered in future studies together with the derivation of the instability domain of an initial monochromatic Stokes wave.

7. Conclusion

In this paper, the second-order perturbation expansion in Lagrangian coordinates has been derived to study the interactions between deep-water surface gravity waves. In its compact and vectorial form, the proposed solution extends initial investigations (Pierson 1961), fully recovers the classical second-order Eulerian expansion (Longuet-Higgins 1963) and naturally includes the well-known Stokes drift velocity. As further illustrated in the case of tri-dimensional wave interactions, a residual spatial Stokes drift will result from harmonics of equal frequency but having different propagation directions. This phenomenon leads to an increase of the wave shape asymmetry along the propagation, and can be related to the development of short-crested wave patterns, as a possible initial stage of formation of horseshoe patterns. Indeed, Lagrangian second-order terms will contribute to sharpening and flattening effects, but, contrary to the first-order correction, these effects are applied in the perpendicular direction to the wave's direction.

The modulation aspect of the BF instability is further shown to be captured as a beat effect in the Lagrangian framework. A periodic modulation emerges between the carrier and two sideband harmonics. As demonstrated, the orbital spectrum remains unchanged as the waves evolve in time, while the corresponding surface Eulerian spectrum exhibits periodical variations for the carrier and sideband harmonic amplitudes. It should be noted that the asymmetric evolution of the sideband harmonics, and the associated frequency downshift, are not recovered at this second Lagrangian order.

The extension of the proposed expansion to the case of varying depth and surface current could also follow the same formalism, and its further investigation should be considered in the future.

Acknowledgements

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Appendix A

A.1. Perfect differential and vorticity

There is no vorticity if the velocity field \mathbf{R}_t can be written in the form:

$$\mathbf{R}_t = \nabla F \quad (\text{A } 1)$$

where F is any scalar function. Noting that $dF = \nabla F \cdot d\mathbf{R}$, we thus have

$$dF = x_t dx + y_t dy + z_t dz. \tag{A 2}$$

Replacing terms dx, dy and dz by their respective particle-label-dependent expressions:

$$dx = x_\alpha d\alpha + x_\beta d\beta + x_\delta d\delta \tag{A 3}$$

$$dy = y_\alpha d\alpha + y_\beta d\beta + y_\delta d\delta \tag{A 4}$$

$$dz = z_\alpha d\alpha + z_\beta d\beta + z_\delta d\delta \tag{A 5}$$

where $d\boldsymbol{\zeta} = (d\alpha, d\beta, d\delta)$ denotes an infinitesimal label variation, we can rewrite

$$dF = (\mathbb{J}\mathbf{R}_t) \cdot d\boldsymbol{\zeta} \tag{A 6}$$

where \mathbb{J} is defined in (2.4). Thereby, if a function $F(\boldsymbol{\zeta}, t)$ can be found such that dF is a perfect differential, there is no vorticity.

A.2. Combination of first-order terms in Newton’s law

Consider the right-hand side of (4.1), $-\mathbf{H}(\Phi_1)\nabla\Phi_{1tt}$, where \mathbf{H} and ∇ are respectively the Hessian and the gradient operator:

$$\mathbf{H}(\Phi_1) = \begin{bmatrix} \phi_{1\alpha\alpha} & \phi_{1\alpha\beta} & \phi_{1\alpha\delta} \\ \phi_{1\alpha\beta} & \phi_{1\beta\beta} & \phi_{1\beta\delta} \\ \phi_{1\alpha\delta} & \phi_{1\beta\delta} & \phi_{1\delta\delta} \end{bmatrix} + \text{c.c.} \quad \text{and} \quad \nabla\Phi_{1tt} = \begin{bmatrix} \phi_{1\alpha tt} \\ \phi_{1\beta tt} \\ \phi_{1\delta tt} \end{bmatrix} + \text{c.c.} \tag{A 7a,b}$$

We can write $-\mathbf{H}(\Phi_1)\nabla\Phi_{1tt} = \mathbf{S} + \mathbf{T}$ where $\mathbf{S} = (S^\alpha, S^\beta, S^\delta)$ and $\mathbf{T} = (T^\alpha, T^\beta, T^\delta)$ are tri-dimensional vectors:

$$\mathbf{S} = -\mathbf{H}(\phi_1)\nabla\phi_{1tt} + \text{c.c.} \tag{A 8}$$

$$\mathbf{T} = -\mathbf{H}(\phi_1)\nabla\phi_{1tt}^* + \text{c.c.} \tag{A 9}$$

where the star superscript ‘*’ means the complex conjugate. We will investigate successively the explicit form of \mathbf{S} and \mathbf{T} . Introducing the two kernels

$$\mathcal{K} = \frac{1}{4} \frac{A(\mathbf{k})A(\mathbf{k}')}{kk'} e^{i(\mathbf{k}+\mathbf{k}')\cdot\boldsymbol{\xi} - i(\omega+\omega')t} e^{(\mathbf{k}+\mathbf{k}')\delta} \tag{A 10}$$

$$\underline{\mathcal{K}} = \frac{1}{4} \frac{A(\mathbf{k})A^*(\mathbf{k}')}{kk'} e^{i(\mathbf{k}-\mathbf{k}')\cdot\boldsymbol{\xi} - i(\omega-\omega')t} e^{(\mathbf{k}+\mathbf{k}')\delta} \tag{A 11}$$

and using ϕ_1 expression, the α component of \mathbf{S} can be written

$$\begin{aligned} S^\alpha &= - \iint [(ik_\alpha)^2(ik'_\alpha)(-i\omega')^2 + (ik_\alpha)(ik_\beta)(ik'_\beta)(-i\omega')^2 + (ik_\alpha)kk'(-i\omega')^2] \mathcal{K} + \text{c.c.} \\ &= - \iint ik_\alpha gk' [k_\alpha k'_\alpha + k_\beta k'_\beta - kk'] \mathcal{K} + \text{c.c.} \\ &= \iint ik_\alpha gk' (kk' - \mathbf{k} \cdot \mathbf{k}') \mathcal{K} + \text{c.c.} \end{aligned} \tag{A 12}$$

Making use of the symmetric integration over k and k' in the second term we can rewrite

$$S^\alpha = \iint g \frac{ikk'}{2} \left(\frac{k_\alpha}{k} + \frac{k'_\alpha}{k'} \right) (kk' - \mathbf{k} \cdot \mathbf{k}') \mathcal{K} + \text{c.c.} \tag{A 13}$$

The same procedure can be applied to the β component, leading to

$$(S^\alpha, S^\beta) = \iint \mathcal{N} gkk' \frac{i(\widehat{\mathbf{k}} + \widehat{\mathbf{k}'})}{2} + \text{c.c.} \tag{A 14}$$

where \mathcal{N} is defined is (4.8). As to the δ component of S , it is found to be

$$\begin{aligned} S^\delta &= - \iint [(ik_\alpha)k(ik'_\alpha)(-i\omega')^2 + (ik_\beta)k(ik'_\beta)(-i\omega')^2 + k^2k'(-i\omega')^2] \mathcal{K} + \text{c.c.} \\ &= \iint \mathcal{N} gkk' + \text{c.c.} \end{aligned} \tag{A 15}$$

For the α component of T ,

$$T^\alpha = \iint ik_\alpha gk' (\mathbf{k} \cdot \mathbf{k}' + kk') \underline{\mathcal{K}} + \text{c.c.}, \tag{A 16}$$

we invert k and k' in the c.c. expression to obtain

$$T^\alpha = \iint g \frac{ikk'}{2} \left(\frac{k_\alpha}{k} - \frac{k'_\alpha}{k'} \right) (kk' + \mathbf{k} \cdot \mathbf{k}') \underline{\mathcal{K}} + \text{c.c.} \tag{A 17}$$

Applying the same technique to the β component we come up with (4.6) which can be written

$$(T^\alpha, T^\beta) = \iint \frac{i(\widehat{\mathbf{k}} - \widehat{\mathbf{k}'})}{2} gkk' (\mathbf{k} \cdot \mathbf{k}' + kk') \underline{\mathcal{K}} + \text{c.c.} \tag{A 18}$$

$$= \iint \frac{i(\widehat{\mathbf{k}} - \widehat{\mathbf{k}'})}{2} gkk' \underline{\mathcal{N}} + \text{c.c.} \tag{A 19}$$

Finally, the δ component of T can easily be derived as

$$T^\delta = \iint gkk' (kk' + \mathbf{k} \cdot \mathbf{k}') \underline{\mathcal{K}} + \text{c.c.} \tag{A 20}$$

$$= \iint gkk' \underline{\mathcal{N}} + \text{c.c.} \tag{A 21}$$

A.3. Combination of first-order terms in the conservation law

The combination of first-order terms in the conservation law (4.12) can be written

$$\Phi_{1\alpha\alpha} \Phi_{1\beta\beta} + \Phi_{1\alpha\alpha} \Phi_{1\delta\delta} + \Phi_{1\beta\beta} \Phi_{1\delta\delta} - \Phi_{1\alpha\beta}^2 - \Phi_{1\alpha\delta}^2 - \Phi_{1\beta\delta}^2 \tag{A 22}$$

where $\Phi_1 = \phi_1 + \phi_1^*$. After the combination of all of the terms in the form $\phi_{1mn}\phi_{1pq}$ and $\phi_{1mn}^*\phi_{1pq}^*$ where m, n, p, q can be any of the variables α, β or δ we obtain

$$\begin{aligned} & \iint [(ik_\alpha)^2(ik'_\beta)^2 + (ik_\alpha)^2k'^2 + (ik_\beta)^2k'^2 - (ik_\alpha)(ik_\beta)(ik'_\alpha)(ik'_\beta) \\ & \quad - (ik_\alpha)k(ik'_\alpha)k' - (ik_\beta)k(ik'_\beta)k'] \mathcal{K} + \text{c.c.} \\ & = \iint \left[\frac{1}{2}(k_\alpha k'_\beta - k_\beta k'_\alpha)^2 + kk'(\mathbf{k} \cdot \mathbf{k}' - kk') \right] \mathcal{K} + \text{c.c.} \\ & = - \iint \left[kk'(1 - \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}') \frac{kk' - \mathbf{k} \cdot \mathbf{k}'}{2} \right] \mathcal{K} + \text{c.c.} \\ & = - \iint \frac{kk' - \mathbf{k} \cdot \mathbf{k}'}{2} \mathcal{N} + \text{c.c.} \end{aligned} \tag{A 23}$$

In the same manner, the combination of the terms $\phi_{1mn}\phi_{1pq}^*$ gives

$$- \iint \frac{kk' + \mathbf{k} \cdot \mathbf{k}'}{2} \mathcal{N} + \text{c.c.} \tag{A 24}$$

Appendix B

In this appendix we prove the existence of the following integral in the sense of the Cauchy PV:

$$\mathbf{r}_2 = \text{PV} \iint_{\mathbb{R}^4} i \frac{\omega \mathbf{k} + \omega' \mathbf{k}'}{2(\omega - \omega')} \mathcal{B} e^{-i(\omega - \omega')t} e^{(k+k')\delta} + \text{c.c.} \tag{B 1}$$

By definition, this can be rewritten

$$\mathbf{r}_2 = \lim_{\gamma \rightarrow 0} \iint_{\mathbb{R}^4 - \mathcal{E}} - \frac{\mathbf{H}(\mathbf{k}, \mathbf{k}')}{4(\omega - \omega')} e^{(k+k')\delta} \tag{B 2}$$

where \mathcal{E} is the domain such as $|\omega - \omega'| < \gamma$ and

$$\mathbf{H}(\mathbf{k}, \mathbf{k}') = (\omega \mathbf{k} + \omega' \mathbf{k}') (1 + \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}') |A(\mathbf{k})| |A(\mathbf{k}')| \sin[(\mathbf{k} - \mathbf{k}') \cdot \boldsymbol{\xi} - (\omega - \omega')t + \varphi_k - \varphi_{k'}] \tag{B 3}$$

where φ_k and $\varphi_{k'}$ are respectively the phases of $A(\mathbf{k})$ and $A(\mathbf{k}')$. We denote $\mathbf{H}^0(k, \widehat{\mathbf{k}}, \widehat{\mathbf{k}}')$ the value of (B 3) when $k = k'$ ($\omega = \omega'$):

$$\begin{aligned} \mathbf{H}^0(k, \widehat{\mathbf{k}}, \widehat{\mathbf{k}}') & = \mathbf{H}(\mathbf{k}, k\widehat{\mathbf{k}}') \\ & = \omega k (\widehat{\mathbf{k}} + \widehat{\mathbf{k}}') (1 + \widehat{\mathbf{k}} \cdot \widehat{\mathbf{k}}') |A(\mathbf{k})| |A(k\widehat{\mathbf{k}}')| \sin[k(\widehat{\mathbf{k}} - \widehat{\mathbf{k}}') \cdot \boldsymbol{\xi} + \varphi_k - \varphi_{k\widehat{\mathbf{k}}'}]. \end{aligned} \tag{B 4}$$

Focusing now on the integral,

$$\mathbf{r}_2^0 = \lim_{\gamma \rightarrow 0} \iint_{\mathbb{R}^4 - \mathcal{E}} - \frac{\mathbf{H}^0(k, \widehat{\mathbf{k}}, \widehat{\mathbf{k}}')}{4(\omega - \omega')} e^{(k+k')\delta} \tag{B 5}$$

and noting that $\mathbf{H}^0(k, \widehat{\mathbf{k}}, \widehat{\mathbf{k}}) = -\mathbf{H}^0(k, \widehat{\mathbf{k}}, \widehat{\mathbf{k}}')$ leads to a vanishing value of \mathbf{r}_2^0 since integration is realized over all directions of \mathbf{k} and \mathbf{k}' . We can thus rewrite (B 2) in the form

$$\mathbf{r}_2 = \lim_{\gamma \rightarrow 0} \iint_{\mathbb{R}^4 - \mathcal{E}} - \frac{\mathbf{H}(\mathbf{k}, \mathbf{k}') - \mathbf{H}^0(k, \widehat{\mathbf{k}}, \widehat{\mathbf{k}}')}{4(\omega - \omega')} e^{(k+k')\delta} \tag{B 6}$$

or, more explicitly,

$$\begin{aligned} \mathbf{r}_2 = & \lim_{\gamma \rightarrow 0} \iint_{\mathbb{R}^4 - \mathcal{E}} \frac{i}{4} (1 + \widehat{\mathbf{k}} \cdot \widetilde{\mathbf{k}}) e^{(k+k')\delta} \\ & \times \left[\frac{(\omega \mathbf{k} + \omega' \mathbf{k}') A(\mathbf{k}) A^*(\mathbf{k}') e^{i(k-k') \cdot \xi - i(\omega - \omega')t} - k\omega (\widehat{\mathbf{k}} + \widetilde{\mathbf{k}}) A(\mathbf{k}) A^*(\mathbf{k}\widetilde{\mathbf{k}}) e^{ik(\widehat{\mathbf{k}} - \widetilde{\mathbf{k}}) \cdot \xi}}{\omega - \omega'} \right] \\ & + \text{c.c.} \end{aligned} \quad (\text{B } 7)$$

The limit when $\omega \rightarrow \omega'$ of the term between brackets can be written

$$-ik\omega (\widehat{\mathbf{k}} + \widetilde{\mathbf{k}}) A(\mathbf{k}) A^*(\mathbf{k}\widetilde{\mathbf{k}}) e^{ik(\widehat{\mathbf{k}} - \widetilde{\mathbf{k}}) \cdot \xi} \quad (\text{B } 8)$$

ensuring (B 2) to be a finite limit and that \mathbf{r}_2 is integrable in the Cauchy PV sense.

REFERENCES

- ABERG, S. 2007 Wave intensities and slopes in Lagrangian seas. *Adv. Appl. Probab.* **39** (4), 1020–1035.
- ABERG, S. & LINDGREN, G. 2008 Height distribution of stochastic Lagrange ocean waves. *Prob. Engng Mech.* **23** (4), 359–363.
- ANNENKOV, S. YU. & SHRIRA, V. I. 1999 Sporadic wind wave horse-shoe patterns. *Nonlinear Process. Geophys.* **6** (1), 27–50.
- BENJAMIN, T. B. & FEIR, J. E. 1967 The disintegration of wave trains on deep water. Part 1. Theory. *J. Fluid Mech.* **27**, 417–430.
- BULDAKOV, E. V., TAYLOR, P. H. & TAYLOR, R. E. 2006 New asymptotic description of nonlinear water waves in Lagrangian coordinates. *J. Fluid Mech.* **562** (1), 431–444.
- CLAMOND, D. 2007 On the Lagrangian description of steady surface gravity waves. *J. Fluid Mech.* **589**, 433–454.
- COLLARD, F. & CAULLIEZ, G. 1999 Oscillating crescent-shaped water wave patterns. *Phys. Fluids* **11** (11), 3195–3197.
- CREAMER, D. B., HENYEV, F., SCHULT, R. & WRIGHT, J. 1989 Improved linear representation of ocean surface waves. *J. Fluid Mech.* **205**, 135–161.
- DAEMRICH, K.-F. & WOLTERING, S. 2008 How nonlinear are linear waves? In *German Joint Symposium on Hydraulic and Ocean Engineering (JOINT 2008)*.
- ELFOUHAILY, T., THOMPSON, D., VANDEMARK, D. & CHAPRON, B. 1999 Weakly nonlinear theory and sea state bias estimation. *J. Geophys. Res.* **104** (C4), 7641–7647.
- FOUQUES, S., KROGSTAD, H. E. & MYRHAUG, D. 2006 A second order Lagrangian model for irregular ocean waves. *J. Offshore Mech. Arctic Engng* **128** (3), 177–183.
- FOUQUES, S. & STANSBERG, C. T. 2009 A modified linear Lagrangian model for irregular long-crested waves. In *ASME 2009 28th International Conference on Ocean, Offshore and Arctic Engineering*, pp. 495–502. American Society of Mechanical Engineers.
- GERSTNER, F. J. 1809 Theorie der wellen. *Ann. Phys.* **32**, 412–445.
- GJOSUND, S. H. 2003 A Lagrangian model for irregular waves and wave kinematics. *Trans. ASME* **125**, 94–102.
- HASSELMANN, K. 1962 On the nonlinear energy transfer in a gravity-wave spectrum. Part 1. General theory. *J. Fluid Mech.* **12**, 481–500.
- HSU, H.-C., CHEN, Y.-Y. & WANG, C.-F. 2010 Perturbation analysis of short-crested waves in Lagrangian coordinates. *Nonlinear Anal.* **11** (3), 1522–1536.
- HSU, H.-C., NG, C.-O. & HWUNG, H.-H. 2012 A new Lagrangian asymptotic solution for gravity-capillary waves in water of finite depth. *J. Fluid Mech.* **14** (1), 79–94.
- KENYON, K. E. 1969 Stokes drift for random gravity waves. *J. Geophys. Res.* **74** (28), 6991–6994.

- KIMMOUN, O., BRANGER, H. & KHARIF, C. 1999 On short-crested waves: experimental and analytical investigations. *Eur. J. Mech. (B/Fluids)* **18** (5), 889–930.
- KINSMAN, B. 1965 *Wind Waves: Their Generation and Propagation on the Ocean Surface*. Prentice-Hall.
- LAMB, H. 1932 *Hydrodynamics*. Cambridge University Press/Dover.
- LEBLANC, S. 2004 Local stability of Gerstner's waves. *J. Fluid Mech.* **506**, 245–254.
- LINDGREN, G. 2006 Slepian models for the stochastic shape of individual Lagrange sea waves. *Adv. Appl. Probab.* **38** (2), 430–450.
- LINDGREN, G. & ABERG, S. 2009 First order Lagrange model for asymmetric ocean waves. *J. Offshore Mech. Arctic Engng* **131** (3), 031602.
- LINDGREN, G. & LINDGREN, F. 2011 Stochastic asymmetry properties of 3D Gauss–Lagrange ocean waves with directional spreading. *Stoch. Models* **27** (3), 490–520.
- LONGUET-HIGGINS, M. S. 1963 The effects of nonlinearities on statistical distributions in the theory of sea waves. *J. Fluid Mech.* **17**, 459–480.
- LONGUET-HIGGINS, M. S. 1986 Eulerian and Lagrangian aspects of surface waves. *J. Fluid Mech.* **173**, 683–707.
- LONGUET-HIGGINS, M. S. 1987 Lagrangian moments and mass transport in Stokes waves. *J. Fluid Mech.* **179**, 547–555.
- NACIRI, M. & MEI, C. C. 1992 Evolution of a short surface wave on a very long surface wave of finite amplitude. *J. Fluid Mech.* **235**, 415–452.
- NOUGUIER, F., GUÉRIN, C. A. & CHAPRON, B. 2009 'Choppy wave' model for nonlinear gravity waves. *J. Geophys. Res.* **114** (C13), 09012.
- PHILLIPS, O. M. 1977 *The Dynamics of the Upper Ocean*. Cambridge University Press.
- PIERSON, W. J. 1961 Models of random seas based on the Lagrangian equations of motion. *Tech. Rep.*. New York Univ., Coll. of Eng. Res. Div., Dept. of Meteorology and Oceanography, prepared for the Office of Naval Research under contract NONR-285(03).
- SHEMER, L. 2010 On Benjamin–Feir instability and evolution of a nonlinear wave with finite-amplitude sidebands. *Nat. Hazards Earth Syst. Sci.* **10** (11), 2421–2427.
- SHRIRA, V. I., BADULIN, S. I. & KHARIF, C. 1996 A model of water wave 'horse-shoe' patterns. *J. Fluid Mech.* **318**, 375–405.
- SOCQUET-JUGLARD, H., DYSTHE, K., TRULSEN, K., KROGSTAD, H. E. & LIU, J. 2005 Probability distributions of surface gravity waves during spectral changes. *J. Fluid Mech.* **542**, 195–216.
- STOKER, J. J. 1957 *Water Waves: The Mathematical Theory with Applications*. Interscience.
- STOKES, G. G. 1847 On the theory of oscillatory waves. *Trans. Camb. Phil. Soc.* **8** (441), 197–229.
- WEBER, B. L. & BARRICK, D. E. 1977 On the nonlinear theory for gravity waves on the ocean's surface. *J. Phys. Oceanogr.* **7** (1), 11–21.
- YAKUBOVICH, E. I. & ZENKOVICH, D. A. 2001 Matrix approach to Lagrangian fluid dynamics. *J. Fluid Mech.* **443**, 167–196.