

## REVISION OF THE AGE OF CONSTRUCTION PHASES OF A MOUND DATED TO THE LATE COPPER–EARLY BRONZE AGE IN EASTERN HUNGARY RELYING ON <sup>14</sup>C-BASED CHRONOLOGIES

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**ABSTRACT.** Ecese Mound is a burial mound in the Hortobágy region of eastern Hungary. Built by prehistoric nomadic peoples from the east, it now stands on the border between two modern settlements. The construction of the mound was assumed to be related to representatives of the Pit Grave Culture populating the area between the Late Copper and Bronze Ages. This theory considered similarities in shape, orientation, and stratigraphy of this mound with other absolute-dated ones in the Hortobágy region alone. The mound comprises two construction layers as indicated by magnetic susceptibility and on-site stratigraphic observations. According to detailed sedimentological, geochemical analyses of samples taken from the bedrock, artificial stratigraphic horizons, and the overlying topsoil, there is a marked similarity between the soil forming the body of the mound in both artificial horizons and the underlying bedrock soil. In contrast the pedological, geological character of the modern topsoil is utterly different. According to our dating results, the uppermost stratigraphic horizon is coeval with the absolute-dated mounds in the region, assigning it to the period of the Pit Grave Culture. However, the lower anthropological horizon is older and dates to between the Early and Late Copper Ages.

**KEYWORDS:** chronology, kurgan, landscape history, radiocarbon dating, sedimentology.

### INTRODUCTION

Burial and dwelling mounds were the very first features subjected to incipient geoarchaeological studies (Jefferson 1783; Forchhammer et al. 1851; Vanuxem 1843). Mounds are among the first studied objects of Hungarian archeology too (Rómer 1868a, 1868b, 1868c, 1878). Since the early period of archaeological excavations determination of the date of origin has been the most important aspect of mound research, as not only the type but also the time of construction displayed large-scale variance. The Hungarian term “kunhalom”<sup>1</sup> (Győrffy 1821; Horváth 1825; Jerney 1851; Gyárfás 1870; Gárdonyi 1893, 1914) is a catch-all category into which all types of mounds have been thrown without consideration to their function or date of origin (Makkay 1964; Tóth and Tóth 2003; Tóth 2006; Barczy et al. 2009; Pető and Barczy 2011; Dani and Horváth 2012). This issue can only be tackled if the initial phase of each research includes stratigraphical and chronological analyses, so the very function and age of the mound are revealed (Sümegi et al. 2015a, 2015b). After setting up a stratigraphy using sedimentological, geochemical, and geophysical methods, a chronology is set up using absolute dates. A precise chronology can only be achieved by radiocarbon analyses (Gazdapusztai 1966–1967; Ecsedy 1979; Sümegi and Hertelendi 1998; Molnár et al. 2004; Gulyás et al. 2010; Molnár and Svingor 2011; Dani and Horváth 2012). Additional information can be obtained with the OSL analyses (Liritzis et al. 2013) of wattle-and-daub fragments or pottery remains conserved in the layers. It is important to note though that pottery and wattle-and-daub fragments piled up with the soil can originate from the very same culture that created the mound itself, but also from earlier cultures. Thus, OSL analysis of the pottery and wattle-and-daub fragments may result in the incorrect conclusion that some layers are older than their actual age (Gazdapusztai 1966–1967).

<sup>1</sup>Meaning “Cumanian mound.”

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In addition, uncertainty of OSL measurements is too high compared to  $^{14}\text{C}$  analysis rendering them unsuitable for the construction of some hundred or a couple of thousand year-resolution chronologies for archeological periods of the younger Holocene. These skewed results can be rectified by mass radiocarbon or AMS dating on organic materials (Molnár et al. 2004, 2013; Barczy et al. 2012; Dani and Horváth 2012) derived from the grave (wooden structure, bulrush shroud covering the corpse) or on the humus content of the piled-up soil.

In this paper results of a comprehensive stratigraphic work complemented by absolute chronology based on  $^{14}\text{C}$  AMS analyses on samples from the Ecse Mound (Sümegei 2012) are discussed along with age-depth models built via Bayesian analysis using models that seem to be best suited for capturing the deposition rates characterizing mound formation.

## LOCATION AND CHARACTERISTICS OF THE SITE

The Ecse Mound is in NE Hungary in the Hortobágy 12 km north-northeast to the town of Karcag (N47°25'31.11'', E20°57'47.71'') at a height of 93.5 m ASL. The mound is 5.5 m high with a length of 75.5 m and width of 67.5 m (Figure 1). The mound occupies a Pleistocene lag-surface wedging into the Holocene alluvia of the Hortobágy. It rises on the eastern end of a slightly elevated, elongated loess ridge that is clearly separable from its surroundings based on its vegetation and geomorphology. Traces of a ditch that was created when the earth was piled up on the mound are barely perceivable (Sümegei 2012).

Ecse Mound is mentioned (Gyárfás 1883; Benedek and Zádorné 1998) first in a charter describing village borders from 1521 (in the form “Echehalma”). In the Early Modern Era it was the border point between the villages of Asszonyszállás and Kápolnás. Today it lies on the administrative border between Karcag and Kunmadaras; the borderline breaks in an angle on the top of the mound.

Manuscript maps from the 18th–19th centuries and later printed maps consistently represent the whole area of the mound as pasture (Bede et al. 2016). In the beginning of the 20th century, however, its southern half was plowed due to the increased demand for arable land, and in 1943, this is the picture presented. Socialist large-scale agriculture and the consequent large-scale landscape transformations did not spare the Ecse Mound, either: in the 1950s rice parcels were established on its southern side, traces of which are still visible. In the 1960s the area served again as pasture and has been used the same way until today (Bede et al. 2016). In the wider vicinity of the mound farmsteads, dirt roads, ditches, embankments, grasslands and lower lying swamps can be found. The mound rises above its marshy, alkaline environment, thus most of its surface is covered by a loess grassland association (*Salvia nemorosae-Festucetum rupicolae*) and its derivatives. Arboreal vegetation is only sparsely present in the area. The mounds are characteristic refugia for the survival of such habitat types, having a significant conservation value, even the plant association itself (Joó 2003; Illyés and Bölöni 2007; Horváth et al. 2011).

## MATERIAL AND METHODS

### Field Survey, Sampling

The first step of our work included the collection of historical and high-resolution regional maps of the area to create a digital elevation model of the surrounding landscape (Figure 1). This was followed by a field survey during which mapping data points have been recorded on the mound itself to provide an up-to-date DEM of the mound for further geomorphological studies. This

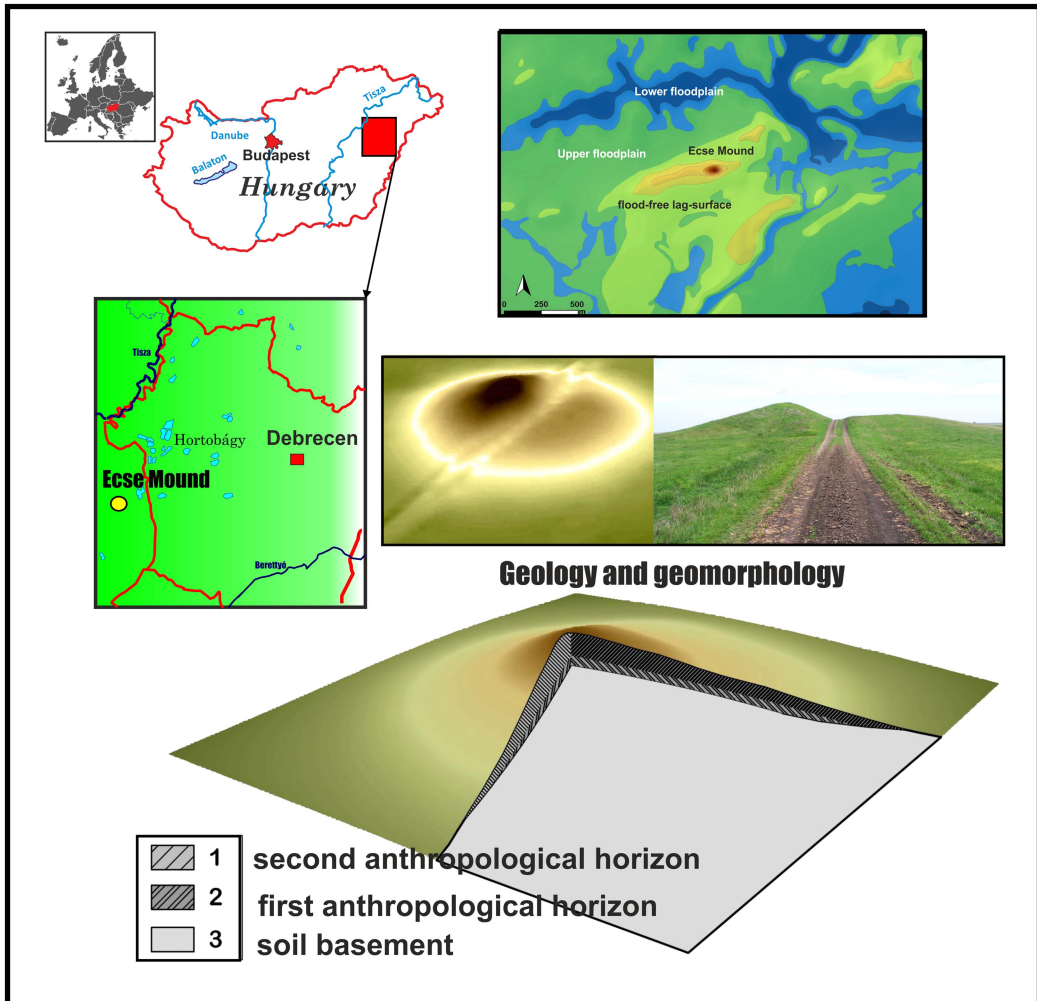


Figure 1 Location, geomorphology and observed stratigraphy of the Ecse Mound in the Hortobágy, NE Hungary.

was complemented by probe coring to reveal the spatial variability of the stratigraphy before actual sampling via undisturbed cores.

Sediment types were determined and described on the field using the Troels-Smith (1955) system internationally accredited for paleoecological works. Both wet and dry colors were determined (Munsell 2000).

**Magnetic Susceptibility**

Measuring magnetic susceptibility (MS) has proved to be one of the best methods to yield reliable stratigraphic data in case of studies of mounds (Sümegei 2012, 2013; Bede et al. 2014, 2015; Sümegei et al. 2015a). For this study samples were taken at 2–4-cm intervals. Prior to the start of the measurement, all samples were crushed in a glass mortar after weighing. Then samples were cased in plastic boxes and dried in air in an oven at 40°C for 24 hr. Afterwards, magnetic susceptibilities were measured at a frequency of 2 kHz using an MS2 Bartington

magnetic susceptibility meter with a MS2E high resolution sensor (Dearing 1994). All the samples were measured three times and the average values of magnetic susceptibility were computed and reported.

### Grain-Size Distribution and Loss on Ignition

The grain size composition of sedimentological samples was carried out using the laser-sedigraph method. First the samples were pretreated with 1 M HCl and H<sub>2</sub>O<sub>2</sub> to remove CaCO<sub>3</sub> and organic content respectively. A more detailed description of the pretreatment process is given by Konert and Vandenberghe (1997). All the samples were measured for 42 intervals between 0.0001 and 0.5 mm using an Easy Laser Particle Sizer 2.0 and Fritsch sieves in Szeged (Hungary). For LOI examination sub-samples were taken at every 2–4-cm intervals and the loss on ignition method was applied, commonly used for the analysis of the organic and carbonate content on calcareous sediments (Dean 1974).

### RADIOCARBON DATING

Six shell samples were submitted for radiocarbon dating taken from major stratigraphic units as depicted in Table 1 and Figure 1. AMS <sup>14</sup>C dating measurements were done in the internationally referenced AMS laboratory of Seattle, WA, USA (Table 1). The dead carbon effect was negligible in case of our chosen taxa because certain herbivorous gastropods are known to yield reliable ages for dating deposits of the past 40 ka with minimal measurement error on the scale of perhaps a couple of decades (Pigati et al. 2004, 2010, 2013; Sümegi and Újvári et al. 2014). This uncertainty is preserved even after calibration yielding us dates on the sub-centennial scale. Considering the presently available multicentennial resolution of prehistoric archeochronology this level of uncertainty suited our needs.

Preparation of the samples and the actual steps of the measurement followed the methods of Hertelendi et al. (1989, 1992) and Molnár et al. (2013). Shells were ultrasonically washed and dried at room temperature. Surficial contaminations and carbonate coatings were removed by pretreatment with weak acid etching (2% HCl) before graphitization. Conventional radiocarbon ages were converted to calendar ages using the software OxCal 4.2 (Bronk Ramsey 2009) and the most recent IntCal13 calibration curve (Reimer et al. 2013). Calibrated ages are reported as age ranges at the 2-sigma confidence level (95.4%). A U<sub>Sequence</sub> age-depth model was constructed for the upper part of the sequence representing the actual mound via Bayesian modeling using OxCal (Bronk Ramsey 2009). As these layers were artificially built up we may presume a relatively uniform deposition rate related to the events of mound formation. In our

Table 1 Conventional (year BP) <sup>14</sup>C ages for Ecse Mound.

Material	Conventional <sup>14</sup> C ages		Fraction of modern	
	(BP yr)	1σ	pMC	1 σ error
<i>Chondrula tridens</i>	531	29	93.60	0.34
<i>Chondrula tridens</i>	2926	25	69.47	0.11
<i>Unios crassus</i>	4281	27	58.69	0.20
<i>Chondrula tridens</i>	5475	30	50.58	0.19
<i>Chondrula tridens</i>	5804	24	48.55	0.17
<i>Anisus spirorbis</i>	10,266	37	28.00	0.13

models we used a U\_Sequence model assuming strictly uniform deposition for the anthropogenic stratigraphic horizons.

## RESULTS

### Lithology and Stratigraphy

The substrate sediment is comprised of fine sand and coarse-grained silt with substantial carbonate, low clay and organic content between 10 and 8 m. The sediment also contained tiny iron-manganese precipitation particles in the form of granules and coating. The overlying stratigraphic unit (8–7.8 m) is composed of very fine to fine sands with medium sand intercalations. This unit is overlain by yellowish brown calcareous clayey silt (7.8–5.8 m.) Between 5.8 and 4.15 m a meadow chernozem soil was encountered with well-developed A and B horizons. This is seen in a sudden increase in Corg values as well (Figure S1). This soil gave the base of the artificial mound. Start of the first man-made horizon was noted between the depths of 4.15 and 4.10 m also seen in elevated Fe content and magnetic susceptibility values (Figure S1). This horizon is overlain by another soil layer marking a different disturbance phase starting around 2.9–2.8 m. Start of this second soil layer is recorded in a drop of magnetic susceptibility and soluble Fe values and a peak in Corg (Figure S1).

The soil built up in the first two anthropological horizons is of polygonal structure with hydromorphic qualities that is very similar to the A and B layers of the underlying meadow chernozem soil. The organic content is significantly higher in this horizon (Figure S1), also showing an abrupt change of the soluble elements and insignificant level of carbonate content. However, the elemental and carbon concentrations must have changed considerably during the soil development process, and later due to ground water table fluctuations and precipitation percolating into the soil. Another disturbance layer was noted just below the modern topsoil (1.5 m) containing pottery as well as wattle and daub fragments. The last 1.5 m is the modern chernozem type topsoil. The sedimentological and geochemical properties of the topsoil overlying the artificial horizons is completely different from the artificially built (Figure S1). This is in line with earlier pedological and sedimentological observations made on other kurgans of the Hortobágy and Nagykunság (Sümegei 1992; Barczy et al. 2003, 2004, 2006; Joó et al. 2007; Sümegei and Szilágyi 2011; Szilágyi et al. 2013), that chernozem soil has developed on top of the artificial pile of kurgans. This type of top soil and related loess grassland association (*Salvia nemorosae-Festucetum rupicolae*) form the topmost layer of the kurgan.

### Chronology

The AMS-based chronology assessment of the Esce Mound was investigated from the bedrock to the first few artificially built-up horizons. According to the dates obtained, the bedrock of the base soil of the mound is dated to the Early Holocene (10,207–9877 cal BC). The upper part of the base soil of polyhedral structure was placed into the Late Neolithic as seen from radiocarbon dates received for a local steppe dwelling mollusk (4723–4558 cal BC) (Figure 2; Table 3; Phase 1). The next age corresponds to the first date at the boundary of the lower anthropological horizon (Figure 2; Tables 1–2). According to the modeled ages this boundary could be placed between 4446–4263 cal BC; i.e. Early Copper Age (Vaday 2004). Shell fragments of the aquatic bivalve *Unio crassus* found in the top boundary of the first artificially piled layer of terrestrial sediment suggest external, most probably human influence. <sup>14</sup>C studies on freshwater shells from Neolithic mounds in SE Hungary indicated no significant dead carbon

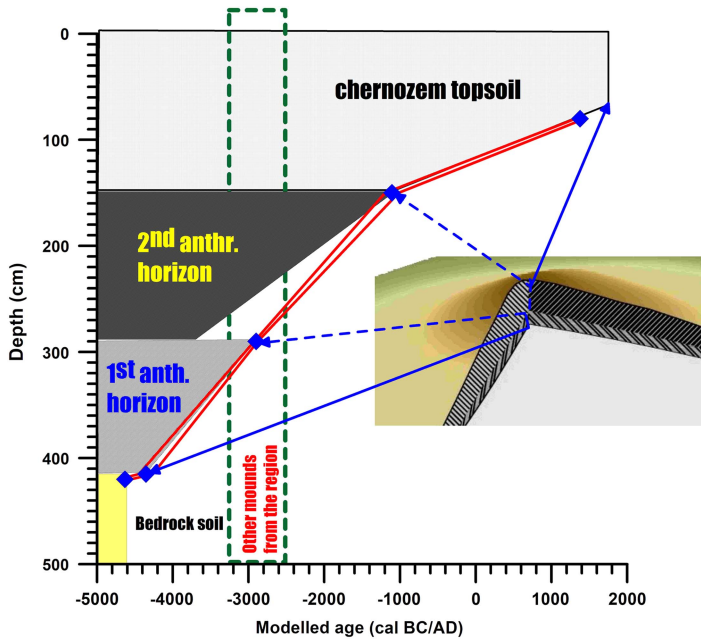


Figure 2 Constructed age-depth model for the mound.

Table 2 Modeled <sup>14</sup>C ages (cal BC/AD) for the individual stratigraphic horizons at the 2σ (95.4%) confidence level.

U_Sequence	Depth	Modeled ages (cal BC/AD)		Agreement (%)	Congruence (%)
		from (2σ)	to (2σ)		
Start		-10,207	-9877		95
			Phase 1		
D-AMS 6514	580	-10,207	-9877	89.1	95
D-AMS 6515	420	-4713	-4552	89.3	96
			Phase 2		
D-AMS 6517	415	-4446	-4263	100.7	97.3
D-AMS 6516	290	-2922	-2878	99.1	98.8
			Phase 3		
D-AMS 6518	150	-1195	-1017	94.5	97.9
D-AMS 6519	80	1316	1439	91.9	98.7
End		1316	1439		98.7

effect on obtained ages at the centennial scale (Gulyás et al. 2010). Modeled ages for this level representing the uppermost part of the first anthropological layer yielded ages 2922–2878 cal. BC (Figure 2; Table 2). As these two ages are bracketing the first anthropological horizon we may presume that the first anthropological horizon must have been built up during the Early Copper Age. The 2nd anthropological horizon clearly postdates the level marking the start of this horizon (Figure 2; Table 2). This this must have been built up during the latest phase of the Late Copper Age. Thus, based on our dates the mound must have been constructed by a

Table 3 Conventional (year BP) and calibrated (cal BC/AD)  $^{14}\text{C}$  ages from different kurgans around Esce Mound in the Great Hungarian Plain (Gazdapusztai 1966–1967; Ecsed 1973, 1979; Dani and Nepper 2006; Horváth et al. 2008, 2013; Dani and Horváth 2012).

Sites	Cultural affiliation	Conventional $^{14}\text{C}$ ages		Calibrated ages at the $2\sigma$ (95.4%) level (cal BC/AD yr)	
		yr BP	$1\sigma$	From	To
Tiszavasvári-Gyepáros	Pit Grave III. phase	4355	35	3087	2899
Tiszavasvári-Deákhalom	Pre Pit Grave II. phase	4350	40	3089	2894
Tiszavasvári-Deákhalom	Pre Pit Grave II. phase	4430	30	3326	2926
Hajdúnánás-Tedej-Lyukashalom	Pre Pit Grave II./III. phase	4270	40	3012	2705
Hajdúnánás-Tedej-Lyukashalom	Pit Grave III./IV. phase	4210	35	2901	2677
Hajdúszoboszló-Árkushalom	Early Pit Grave III. phase	4385	35	3095	2910
Balmazújváros-Hortobágy- Árkus-Ketőshalom	Early Pit Grave III. phase	4320	35	3020	2888
Hortobágy-Ohat-Dunahalom	Early Pit Grave III. phase	4030	35	2832	2470

local community of the Pit grave (Yamna) Culture (Gazdapusztai 1966/1967; Ecsedy 1979; Vaday 2004; Horváth et al. 2013) inhabiting the area between the late Copper Age–Bronze Age period.

## DISCUSSION

Mounds were constructed in the Hortobágy and its wider region as early as 3300 BC (first appearance of the peoples of the Pit-Grave Culture or Pre-Pit Grave Culture) until as late as the 15th century, which is a 4900–5000-yr timespan.

There have been assumptions (Bede et al. 2014, 2015) before the current chronological study that the Esce Mound had been built by communities of the Pit-grave (Yamna) culture (Merpert 1974; Gimbutas 1980; Rassamakin 1994) considering the similarities in shape, orientation and stratigraphy of this earth-pyramid, and comparative geomorphological research of other mounds in the Hortobágy region (Sümegei 1992; Barczy et al. 2003, 2004, 2006; Joó et al. 2007; Sümegei and Szilágyi 2011; Szilágyi et al. 2013). According to our investigations the first phase of mound construction could be placed to the Early Copper Age, while the second part dates to the the Late Copper Age, i.e. the time of first infiltration of the Yamna (Pit Grave or Ochre Grave) Culture. It is interesting to note that  $^{14}\text{C}$  ages from surrounding mounds covering a period of the Pre-Pit, Early Pit, and Classical Pit Cultures, clearly overlap with our ages obtained for the base of the second anthropogenic unit of Esce Mound (Table 3; Figure 2).

## CONCLUSION

This paper can be considered as a pilot research to a much larger scale scientific program with the objective of studying kurgans by the means of undisturbed core sampling. By publishing these preliminary data, we also wanted to draw attention to the need of concentrated and focused research efforts, and using standardized methodology in kurgan research, so the results

from different researches done by different research groups are consistent and comparable. As of today, comparative studies are virtually impossible not only due to the different drilling and sampling techniques, but also for the lack of standardized methodology in fine stratigraphy and common understanding of geology, paleoecology, and geoarchaeology.

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## SUPPLEMENTARY MATERIAL

To view supplementary material for this article, please visit <https://doi.org/10.1017/RDC.2018.107>

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