

Kalana Lagerstätte crinoids: Early Silurian (Llandovery) of central Estonia

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Abstract.—The Kalana Lagerstätte of early Aeronian (Llandovery, Silurian) age in central Estonia preserves a diverse shallow marine biota dominated by non-calcified algae. This soft-tissue flora and decalcified and calcified crinoids are preserved *in situ*, in a lens of microlaminated, dolomitized micrite interbedded in a sequence of dolomitized packstones and wackestones. Although the Lagerstätte is dominated by non-calcified algae, crinoids (together with brachiopods and gastropods) are among the most common organisms that were originally comprised of a carbonate skeleton. Two new crinoids are described from this unit, *Kalanacrinus mastikae* n. gen. n. sp. (large camerate) and *Tartucrinus kalanaensis* n. gen. n. sp. (small disparid). Interestingly, these two crinoids display contrasting preservation, with the more common large camerate preserved primarily as a decalcified organic residue, whereas the smaller disparid is preserved primarily in calcite. Preservation was assessed using elemental mapping of C, Ca, S, and Si. Columns have the highest portion of Ca, once living soft tissue is indicated by C, S was dispersed as pyrite or associated with organics, and Si is probably associated with clay minerals in the matrix. This new fauna increases our understanding of the crinoid radiation on Baltica following Late Ordovician extinctions.

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Introduction

Crinoids suffered a mass extinction at the onset of Late Ordovician glaciations (Eckert, 1988; Donovan, 1989, 1994; Peters and Ausich, 2008), which resulted in the demise of the early Paleozoic Crinoid Evolutionary Fauna (Ausich et al., 1994; Baumiller, 1994). Recovery and establishment of the middle Paleozoic Crinoid Evolutionary Fauna occurred through the Llandovery (Peters and Ausich, 2008; Ausich and Deline, 2012). Historically, our understanding of this recovery was hampered by a general lack of Llandovery crinoid faunas. Beginning during the 1980s, Llandovery crinoids have been described primarily from Laurentia and Avalonia (e.g., Eckert, 1984; Ausich, 1984a, b, 1985, 1986a, b, c, 1987; Donovan, 1993; Fearnhead and Donovan, 2007; Ausich and Copper, 2010). The Baltica paleocontinent is well known for its rich Wenlock and Ludlow (Silurian) crinoid faunas, primarily from Gotland, Sweden (e.g., Angelin, 1878; Bather, 1893; Franzén, 1982, 1983). Only recently have Llandovery crinoids been described from Estonia (Ausich et al., 2015; Ausich and Wilson, 2016).

Besides Avalonia and Baltica faunas, relatively little is known about Silurian crinoids (crown-based taxa) from the remainder of Europe. Recently, one Wenlock taxon was described from the southern Urals (Bogolepova et al., 2018). Crinoids are also known from the Ludlow of the Czech Republic, and scyphocrinitidae are known from several sites in the Silurian-Devonian boundary strata (Webster and Webster, 2013).

Here we report a small crinoid fauna dominated by a single species from the Kalana Konservat Lagerstätte of central Estonia. The first reports of the Kalana Lagerstätte biota have been published recently (Tinn et al., 2009, 2015; Mastik and Tinn, 2015, 2017; Tinn and Märss, 2018). Because of the unique kind of preservation, only two crinoid species are well enough understood to be fully described and named: a new diplobathrid camerate, *Kalanacrinus mastikae* n. gen. n. sp., and a new disparid, *Tartucrinus kalanaensis* n. gen. n. sp. Whereas the unusual preservation, described below, is not ideal for crinoid preservation, this Lagerstätte offers a unique glimpse of the non-calcified fauna and flora that may have thrived alongside crinoids in other Silurian subtidal benthic communities.

Geographic and stratigraphic occurrences

Specimens described herein were collected from the Kalana Lagerstätte exposed in the Kalana (Otisaare) Quarry located near the village of Kalana, in central Estonia (58.722118°N, 26.038226°E) (Fig. 1). The operating quarry opens the upper part of the Jõgeva Beds, Nurmekund Formation, Raikküla Regional Stage (Llandovery, Silurian). The Raikküla Stage consists of a variety of carbonate rocks. The Kalana section is dominated by wackestones and packstones, but the strata also contain numerous 1–20 mm thick lenses and irregular interbeds of light- to dark-brown, organic-rich, microlaminated, argillaceous limestones.

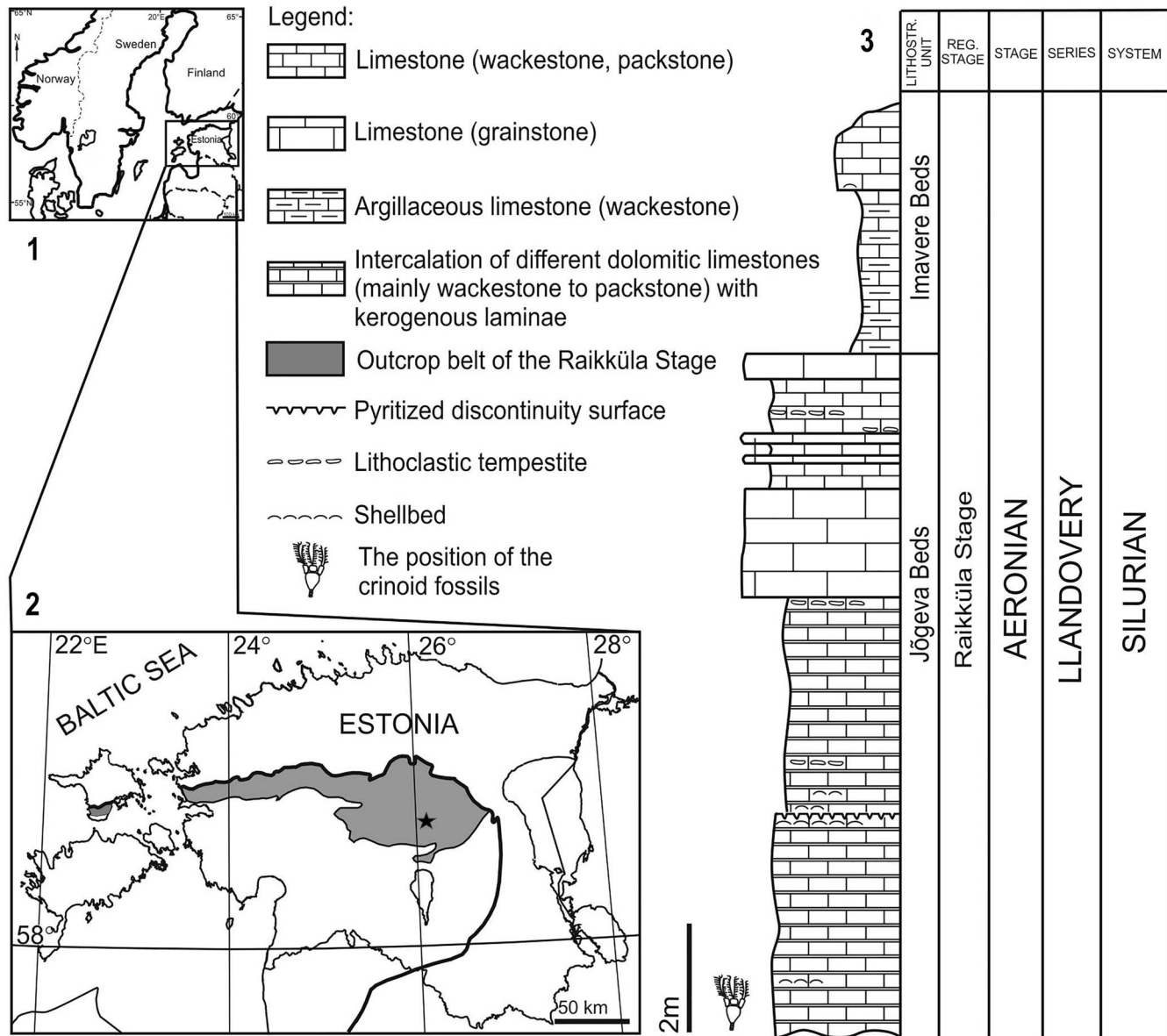


Figure 1. Locality and stratigraphic column for the Kalana Lagerstätte. (1) General map of northwestern Europe; (2) map of Estonia, with the outcrop belt of the Raikküla Stage indicated by shading, and the Kalana Quarry is indicated by a star; (3) stratigraphic column of the Kalana Quarry with the position of the Kalana Lagerstätte indicated by the profile of a crinoid.

Strata containing the exceptionally preserved fossils were present only at the lowermost part of the section, near the bottom of the quarry (Fig. 1). Strata containing the Kalana exceptionally preserved fossils are no older than the middle Aeronian *Pribylograptus leptotheca* graptolite Zone and are in the middle of the *Pranognathus tenuis* conodont Zone (Ainsaar et al., 2014; Männik et al., 2016), which places these strata in the middle Aeronian.

Preservation of the Kalana Lagerstätte

The Kalana Lagerstätte is within a shallow-water carbonate sequence. The Jõgeva Beds are dominated by dolomitized packstones and wackestones that contain numerous intervals of 1–10 mm thick lenses of light- to dark-brown, organic-rich,

microlaminated, commonly dolomitized micrite. As described in Tinn et al. (2009, 2015), Mastik and Tinn (2015, 2017), Männik et al. (2016), and Tinn and Märss (2018), the Kalana Lagerstätte preserves numerous taxa of non-calcified thallophytic algae, the most diverse among them dasyclades (Chlorophyta), but the most abundant and common algal species is *Leveilleites hartnageli* Foerste, 1923, described as a putative rhodophyte. Common faunal elements include graptolites, conodonts, scolecodonts, bryozoans, sponges, and crinoids. The majority of gastropod and rhynchonelliformean brachiopod specimens occur in storm-deposited coquina lenses. Although relatively rare, the remainder of the fauna includes tabulate and rugose corals, orthocone and coiled cephalopods, trilobites, and rare vertebrate fossils (Tinn et al., 2009; Ainsaar et al., 2014; Mastik and Tinn, 2015; Männik et al., 2016; Tinn and Märss, 2018). The

preservation varies largely from soft-bodied fossils with exquisitely preserved details to decalcified molds of shelly fauna.

Element mapping diagrams (Fig. 2) present the distribution of elements in the crinoid specimens. The column has the largest portion of preserved calcite (marked with element calcium, Ca, in the diagrams) in the fossils (Fig. 2.1–2.6). Calcium distribution is correlated with the height and width of the columnals. Nodals are the widest and highest columnals, and priminternodals, secundinternodals, and tertinternodals decrease progressively in width and height. The largest amount of calcium is retained in the position of the nodals with decreasing amounts in the priminternodal and secundinternodal positions (Fig. 2.4). Otherwise, calcium is more randomly distributed with a concentration along one side of the column. The soft tissues of once living animals have left a fine carbonaceous (C) film on and around the specimens (Fig. 2.3).

The arms have a different pattern of preservation (Fig. 2.7–2.12). In the arms calcium is mostly randomly distributed, but again, a concentration is present along one side of the arm (Fig. 2.10). Some of the brachials have a small amount of calcite (Ca in the diagrams), but large parts of them, as well as most of the attached pinnules, are preserved as carbonaceous films only (Fig. 2.9). Silica (Si) in the matrix is probably associated with mostly clay minerals, and sulfur (S) occurs either as dispersed pyrite or is associated with organics.

Whereas the paleoenvironment of the Kalana Lagerstätte and the mechanism of fossilization of these exceptionally preserved specimens are not completely clear yet, some features indicate microbial activity. The role of microbes in the preservation of soft tissues has been demonstrated by a number of experiments and actuopaleontologic studies (for further references see Briggs and McMahon, 2016). At the same time, microbes can also play an important role in the dissolution of minerals, forming chemically complex and reactive microenvironments (Uroz et al., 2009; Dong, 2010; Ahmed and Holmström, 2015). Thus, we currently hypothesize that the partial preservation and/or dissolution of the Kalana crinoid specimens should be attributed to microbial activity on the specimens.

In the case of crinoids, the vast majority of crinoid material is partly decalcified and flattened (Fig. 3.1, 3.2); however, curiously, a few specimens have typical preservation with crinoid ossicles preserved as single crystals of calcite. Rarely (e.g., TUG 1736-8), a single specimen grades from a calcified proximal calyx to a decalcified distal calyx and arms (Fig. 3.3). A decalcified specimen is preserved with shallow relief and a faint brown to black coloration, which is in part carbonaceous and similar to the preservation of non-calcified alga (Fig. 3.1, 3.2). Carbon films are preserved in primarily what would have been the inside of a calyx (Figs. 4.1, 5), and some carbonaceous material is preserved on what is best interpreted as along the outside of a specimen. The two crinoids from the Lagerstätte beds have contrasting preservation. *Kalanacrinus mastikae* n. gen. n. sp. is nearly always decalcified and flattened, whereas all known specimens of *Tartuacrinus kalanaensis* n. gen. n. sp. are preserved in calcite. Further, some isolated pluricolumnals in the Lagerstätte beds are calcified, and these belong to *T. kalanaensis* n. gen. n. sp. One would assume that the smaller crinoid *T. kalanaensis* n. gen. n. sp. would have been more susceptible to decalcification, but it was not the case. The much larger

K. mastikae n. gen. n. sp. would have contained much more soft tissue than *T. kalanaensis* n. gen. n. sp. It is possible that the larger volume of organics contributed to conditions that favored decalcification in diagenetic microenvironments of the Kalana Lagerstätte.

Unfortunately, the flattened, decalcified specimens of *Kalanacrinus mastikae* n. gen. n. sp. tend to split along a single crown surface that is typically one calyx wall and through the arms. Parts and counterparts of this material are two views of a single side of the body wall, as well as a slice through the arms. Either the inside or outside of poorly defined calyx plates is preserved and the corresponding mold of that surface (either internal or exterior) will preserve the reverse relief. This preservation is especially challenging for character discrimination and definition of taxa. Regardless, the full morphology of *K. mastikae* n. gen. n. sp. can be assembled from numerous specimens.

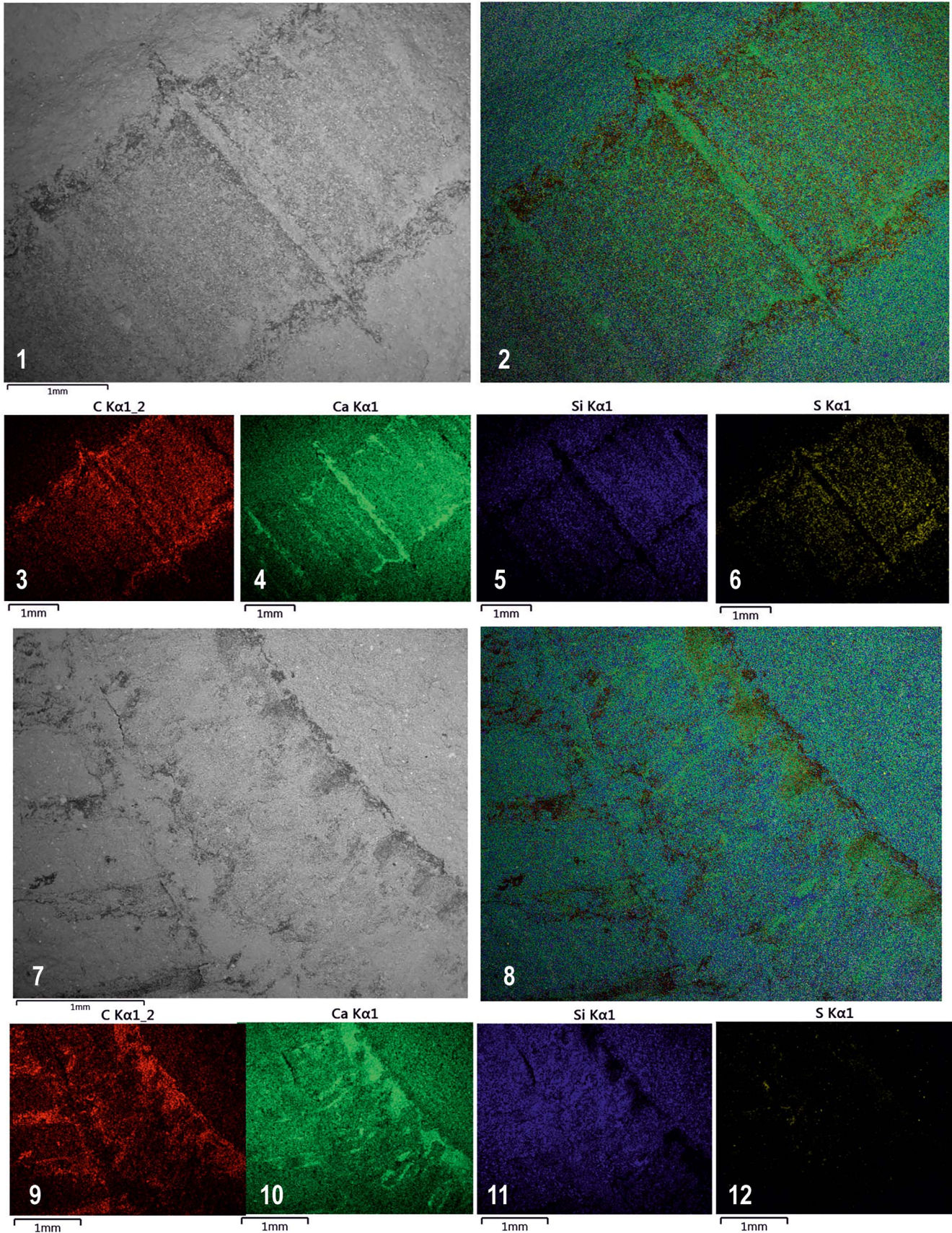
Crinoids are typically well preserved with calcite in the Silurian of Baltica (Bather, 1893; Franzén, 1982, 1983; Ausich et al., 2012, 2015; Ausich and Wilson, 2016), and the mode of preservation in the Kalana Lagerstätte is unique in Baltica. However, Silurian Lagerstätte are known from several locations in Ludlow strata of Laurentia (e.g., Erdtmann and Prezbindowski, 1974; Lo Duca, 1990; Klussendorf, 1994; Lo Duca and Brett, 1997; Saunders et al., 2009). Several aspects of the Kalana Lagerstätte preservation and biota are similar to these Laurentian occurrences. However, in North America, only the Mississinewa Shale Member of the Wabash Formation contains well-preserved crinoids (Lane and Ausich, 1995). Like Kalana crinoids, Mississinewa crinoids are preserved as molds and casts with minor amounts of calcite preservation. However, organic residues are absent in the molds and casts of Mississinewa crinoids.

Materials and methods

Materials used for the descriptions were collected during the field seasons of 2013 and 2014. Most of the crinoids were found in a lens with a diameter of ~6–7 m, at the lowermost level (bottom) of the eastern part of the quarry. The position of the specimens lying in close proximity to each other and the exquisite preservation of the finest details suggest that the specimens were not transported and were most probably buried in situ. Among other fossils associated with these crinoids, the most notable are algae, such as *Leveilleites hartmageli* Foerste, 1923; *Palaeocymopolia silurica* Mastik and Tinn, 2015; and *Kalania pusilla* Tinn et al., 2015. Considering their nearly complete preservation, the algae presumably lived close to the crinoids.

Specimens were photographed with the stereomicroscope system Leica S9i at the Department of Geology, University of Tartu. Specimens were also studied with the Zeiss EVO MA15 scanning electron microscope (SEM) backscattered electrons detector (BSE) in low vacuum regime. Elemental analyses were performed with Oxford X-MAX 80 energy dispersive detector system and Aztec Energy software.

Specimens of *Kalanacrinus mastikae* n. gen. n. sp. are largely preserved as flattened molds with organic residue. Thus, the measured width of a fossil specimen does not represent the width of the specimen when it was alive. To estimate the true width, the diameter of a flattened specimen was doubled to yield the circumference, and the circumference was divided by π to



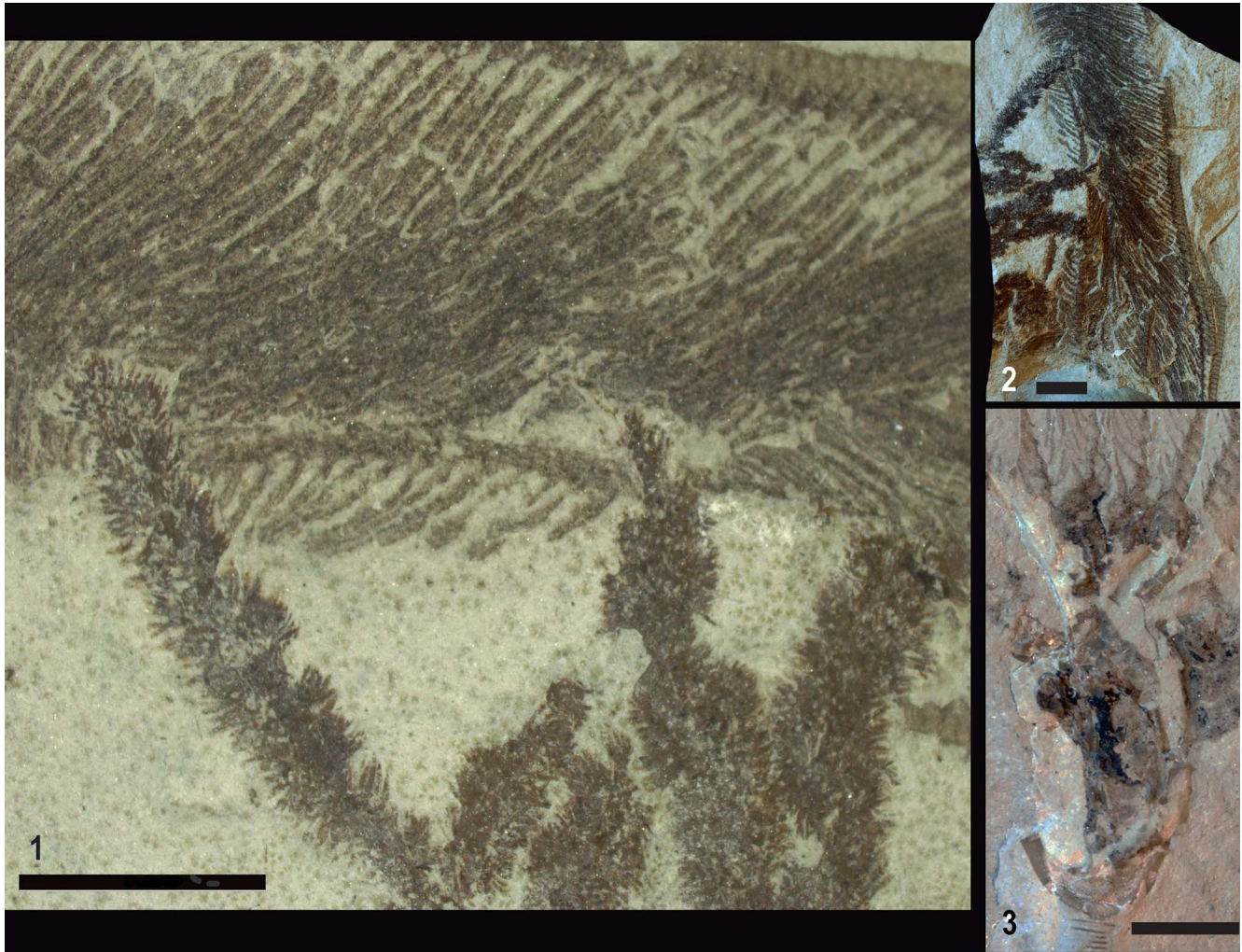


Figure 3. Preservation details of *Kalanacrinus mastikae* n. gen. n. sp.: (1, 2) preservation of adjacent decalcified-calcite arms and non-calcareous algae; TUG 1736-16; (1) enlargement of portion of (2); (3) specimen with most of calyx and proximal arms preserved as carbon film, but infrabasal and basal plate preserved in calcite; paratype, TUG 1736-8; all scale bars 5.0 mm.

estimate the uncompacted diameter when the animal was alive. The nature of this material is such that most specimens described here were collected as molds. The letters A and B following a specimen number refer to the mold and cast of a single specimen. Where it can be determined, the half in positive relief is designated as A, and the half in negative relief is designated B.

Repositories and institutional abbreviations.—Specimens are deposited in the following institutions: TUG, University of Tartu; GIT, Institute of Geology, Geological Institute of Tallinn.

Systematic paleontology

The superfamilial classification used here follows Cole (2017), Wright (2017), and Wright et al. (2017); and family-level classifications follow Moore and Teichert (1978). Morphologic

terminology follows Ubaghs (1978) and Ausich et al. (1999), with modifications as noted in Zamora et al. (2015). The plating of interrays is given by the number of plates in each range from proximal-most plate to the last range before the tegmen. In the posterior interray, the primanal is indicated by “P,” and the first interradial plate in regular interrays is indicated by “1.” A “?” indicates that more distal plating is unknown. Abbreviations used in designating measurements include CrH, crown height; CaH, calyx height; CaW, calyx width; CoH, column height. An * indicates a measurement was incomplete. Following Ausich’s (2018) work on the distribution of simple and compound radial plates in disparid crinoids, the radial circling configuration is given as five numbers, so that 11212 signifies that the A-ray radial plate is simple, the B-ray radial plate is simple, the C-ray radial plate is compound, the D-ray radial plate is simple, and the E-ray radial plate is compound.

Figure 2. *Kalanacrinus mastikae* n. gen. n. sp. (1) BSE image of the specimen (TUG 1736-21) showing variable z-contrast between carbon, calcite, and sedimentary matrix; (2) montage of four charge contrast images of the specimen; (3–6) elemental maps for C, Ca, Si, and S of the same specimen; (7) BSE image of the specimen (TUG 1736-22), showing variable z-contrast between carbon, calcite, and sedimentary matrix; (8) montage of four charge contrast images of the specimen; (9–12) elemental maps for C, Ca, Si, and S of the same specimen.

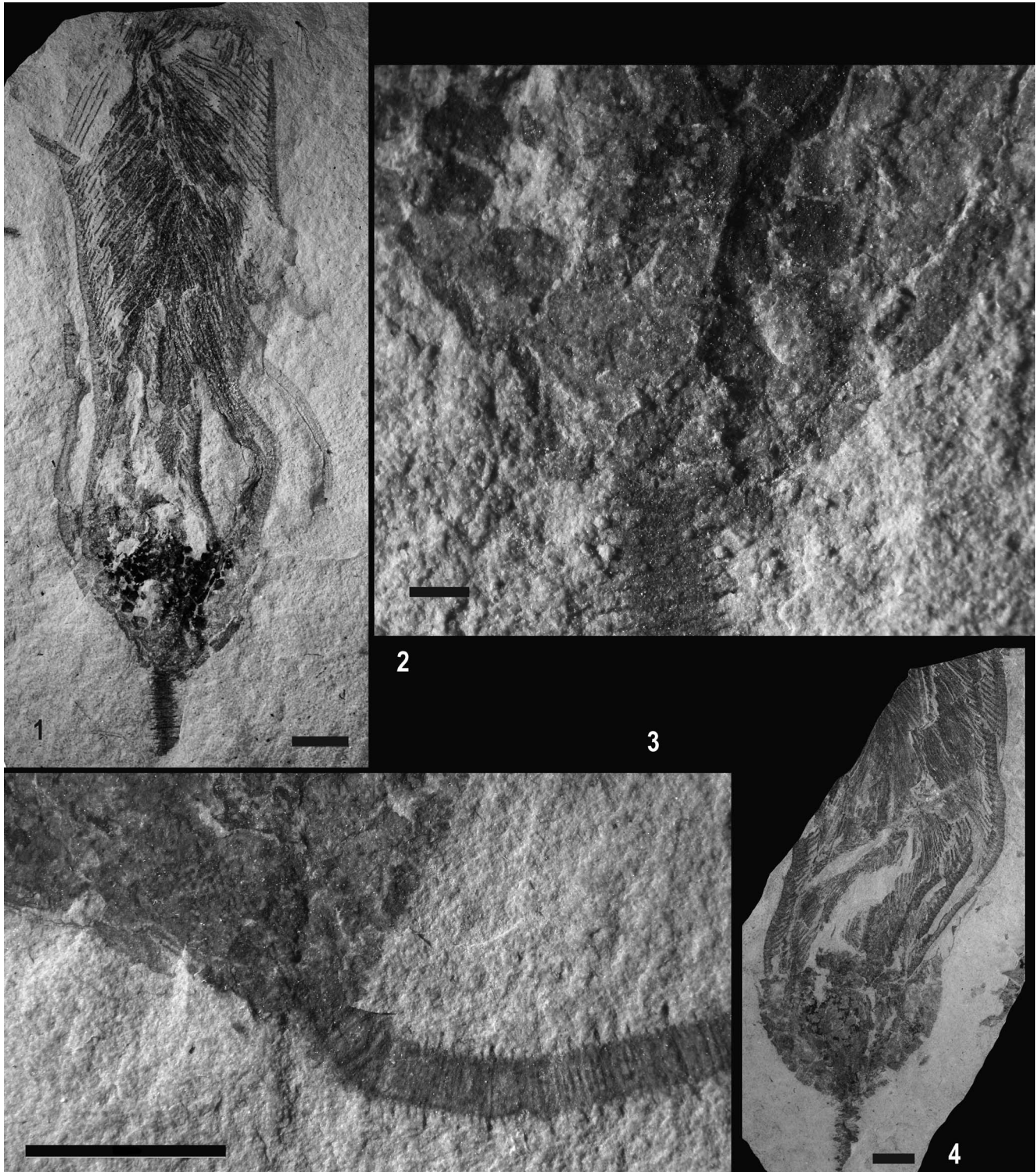


Figure 4. *Kalanacrinus mastikae* n. gen. n. sp. (1, 2) Lateral view of calyx, holotype, TUG 1736-3A; (2) enlargement of portion of (1) with low-angle highlight, note median ray ridge that divides on the radial plate and connects to like ridges on basal plates; short portion of infrabasal plate visible, scale bar 1.0 mm; (3) proximal calyx with length of column attached, note wide and very narrow epifacet extensions on most columnals; paratype, TUG 1736-7; (4) lateral view of calyx, paratype, TUG 1736-4; all scale bars, except part 2, 5.0 mm.

Class Crinoidea Miller, 1821
 Subclass Camerata Wachsmuth and Springer, 1885
 Infraclass Eucamerata Cole, 2017
 Order Diplobathrida Moore and Laudon, 1943

Family Dimerocrinitidae von Zittel, 1879
 Genus *Kalanacrinus* new genus

Type species.—*K. mastikae* n. gen. n. sp.; by monotypy.



Figure 5. *Kalanacrinus mastikae* n. gen. n. sp. (1) Lateral view of calyx with arms mostly closed, paratype, TUG 1736-5; (2) lateral view of calyx with arms open, paratype, TUG 1736-6A; all scale bars 5.0 mm.

Diagnosis.—As for the type species by monotypy.

Occurrence.—Silurian (Llandovery, Aeronian); Estonia (Baltica paleocontinent).

Etymology.—The genus name recognizes the village Kalana, Estonia, that is near the quarry where this specimen was collected.

Remarks.—As outlined in Ausich and Copper (2010, table 5), 10 genera of Silurian dimerocrinitids are known with only

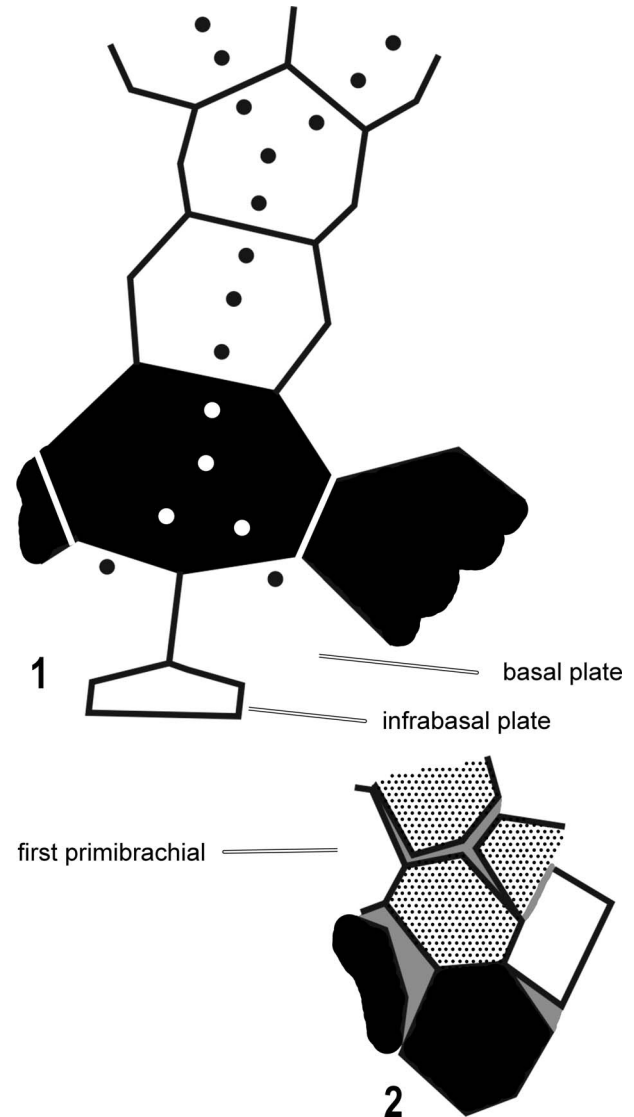


Figure 6. Camera lucida drawings of *Kalanacrinus mastickae* n. gen. n. sp. depicting calyx plating. (1) Aboral cup and ray plating in TUG 1736-3A (holotype); compare to Fig. 4.2; (2) calyx plating in TUG 1736-6B (paratype), despite plate displacement, first interradial plate on shoulders of two radial plates and two interradial plates directly above first interradial plate; compare to Fig. 7; black = radial plates; dotted pattern = interradial plates; large dots = along midline of ray ridges; gray = matrix.

four having 20 atomous free arms, including *Dimerocrinites* Phillips, 1839; *Eucrinus* Angelin, 1878; *Nexocrinus* Eckert, 1984; and *Cybelecrinus* Ausich and Copper, 2010. *Kalanacrinus* n. gen. is the fifth. *Kalanacrinus* n. gen. is most similar to *Cybelecrinus* and is compared below to these four genera. *Dimerocrinites* has a medium bowl-shaped calyx, stellate and smooth plate sculpturing, infrabasal plates not visible in lateral view, antitaxial ridge present, interray regions depressed with respect to rays, first primibrachial tetragonal (rarely hexagonal), primaxil shape pentagonal (rarely hexagonal), fixed pinnules absent, 10–20 free arms, atomous free arms, biserial pinnulate brachials, and circular columnals. *Eucrinus* has a medium cone-shaped calyx, smooth and nodose plate sculpturing, distal-most portions of infrabasal plates visible in lateral view, antitaxial ridge present, interray

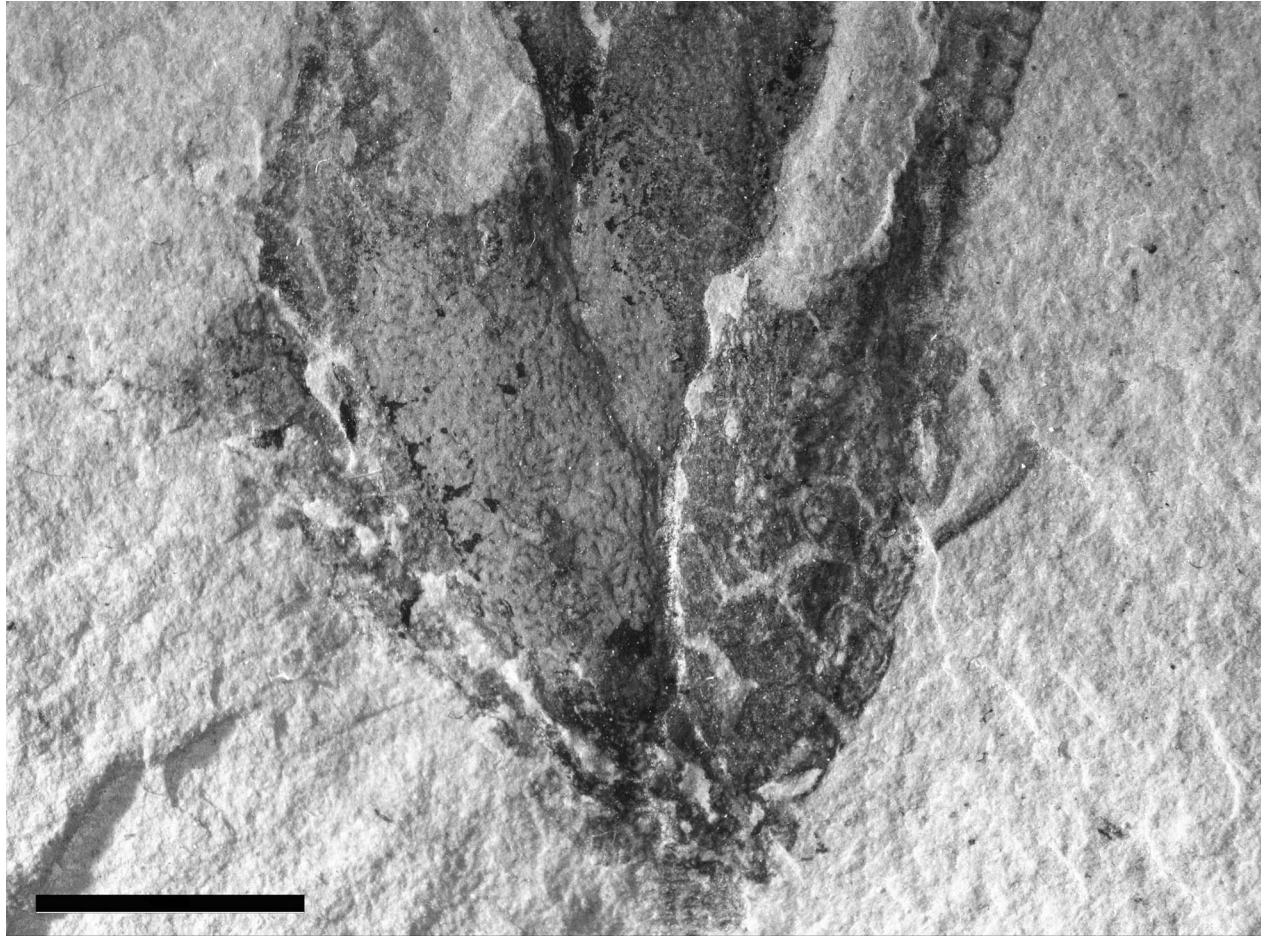


Figure 7. *Kalanacrinus mastikae* n. gen. n. sp. Enlargement of calyx with plate morphology preserved on the right and texture of internal mold on left; note inter-radial plates slightly separated and two plates above the proximal interradial plate in this regular interray; paratype, TUG 1736-6B; scale bar, 5.0 mm.

regions depressed with respect to rays, first primibrachial hexagonal, primaxil pentagonal, second secundibrachial axillary, fixed pinnules absent, 20 free arms, atomous free arms, biserial pinnulate brachials, and circular columnals. *Nexocrinus* has a medium cone-shaped calyx, smooth plate sculpturing, distal-most portions of infrabasal plates not visible in lateral view, anitaxial ridge present, interray regions inflated with rays, first primibrachial hexagonal, primaxil shape heptagonal, second to fourth secundibrachial axillary, fixed pinnules absent, 20 free arms, branching free arms, rectilinear uniserial pinnulate brachials, and circular columnals. *Cybelecrinus* has a medium cone-shaped calyx, stellate and smooth plate sculpturing, distal-most portions of infrabasal plates visible in lateral view, anitaxial ridge present, interray regions flush with rays, first primibrachial hexagonal, primaxil shape heptagonal, fourth secundibrachial axillary, fixed pinnules present, 20 free arms, atomous free arms, rectilinear uniserial pinnulate brachials, and circular columnals. In contrast, *Kalanacrinus* n. gen. has a high cone- to bowl-shaped calyx, nodose plate sculpturing, distal-most portions of infrabasal plates visible in lateral view (condition of the anitaxial ridge unknown), interray regions flush with rays, first primibrachial hexagonal, primaxil shape heptagonal, third secundibrachial axillary, fixed pinnules

present, 20 free arms, atomous free arms, rectilinear uniserial pinnulate brachials, and circular columnals.

Kalanacrinus mastikae new species

Figures 2–9

Type.—Holotype: TUG 1736-3A (part), TUG 1736-3B (counterpart).

Diagnosis.—Dimerocrinitid with high cone- to bowl-shaped calyx, nodose plate sculpturing, distal-most portions of infrabasal plates visible in lateral view (condition of anitaxial ridge unknown), interray regions flush with rays, first primibrachial hexagonal, primaxil shape heptagonal, third secundibrachial axillary, fixed pinnules present, 20 free arms, atomous free arms, rectilinear uniserial pinnulate brachials, and circular columnals.

Occurrence.—*Kalanacrinus mastikae* n. gen. n. sp. is only known from the Kalana Lagerstätte in the Kalana Quarry, near the village of Kalana, Jõgeva County, central Estonia; Silurian, Llandovery, middle Aeronian.

Description.—Crown medium in size; high bowl- or cone-shape (Fig. 4.1, 4.4); arms not grouped; calyx plate sculpturing nodose



Figure 8. *Kalanacrinus mastikae* n. gen. n. sp. Large set of arms articulated to partial calyx, arms preserved with especially dark coloration; note small noncalcareous alga adjacent to arms in upper right portion of image; TUG 1736-12; scale bar, 5.0 mm.

with median ray ridges (Figs. 4.2, 4.4, 5.1) that divide on radial plates and continue proximally to basal plates.

Infrabasal circling ~50% of calyx height; presumably five infrabasal plates, only distal corners of infrabasals visible in lateral view (Figs. 3.3, 4.2, 6.1). Basal circling ~20% of calyx height, entire circling visible in lateral view; presumably five basal plates, branches from ray ridges of the superjacent radial plates meet at center of basal plate and divides proximally presumably leading to the infrabasal plates. Radial circling ~24% of calyx height; radial plates presumably five, heptagonal, ~1.5 times wider than high; larger than basal plate.

Normal interrays do not interrupt radial plate circling, in contact with tegmen; plate sculpturing nodose. First interrarial plate hexagonal, wider than high, intermediate in size between radial and basal plates and first primibrachial plates; second range with two plates, proximal plating 1-2-2-? (Figs. 6.2, 7). Interrarial

regions and intrabrachial regions within half-rays connect to tegmen. Fixed pinnules may be present.

Primal present, CD interray wider than normal interrays, but other details of CD interray not preserved.

First primibrachial fixed, hexagonal wider than high, smaller than radial plates but approximately the same size as the first interrarial plate in normal interrays; second primibrachial axillary, heptagonal (Figs. 4.2, 6.2). Third secundibrachial fixed, axillary. Intrabrachial plates between adjacent secundibrachial half-rays, plating 1-2-2-?, in contact with tegmen. Tegmen unknown.

Free arms 20, become free on approximately the ninth tertribrachial, atomous (Fig. 5.1). Brachial plates rectilinear uniseriate, wider than high; densely pinnulate (Fig. 8) with very long pinnules (Fig. 9.1, 9.2).

Column circular, holomeric, and heteromorphic. In proximal mesistele, nudinodals separated by one priminternodal and two secundinternodals, in more distal parts of the mesistele, tertinternodals present (Figs. 2, 4.3). Nudinodals slightly higher than internodals that are all approximately the same height. Nudinodals with wide, narrow flange that doubles the width of the columnal, first internodal with narrower flange, secundinternodals and tertinternodals without flange (Fig. 4.3). Details of columnal facets and holdfast not known.

Etymology.—The species name recognizes Viirika Mastik for her contributions to understanding the Kalana Lagerstätte, which supported this study in many ways.

Materials.—Holotype: TUG 1736-3(A, B); paratypes: TUG 1736-4(A, B)–TUG 1736-11; other non-type specimens: TUG 1736-12(A, B), TUG 1736-17(A, B), and many unnumbered specimens.

Measurements.—Holotype: TUG 1736-3: CrH, 56.0*; CaH, 14.5; CaW, 9.5; CoH, 2.0*. Paratypes: TUG 1736-4: CrH 52.0*; CaH, 15.0; CaW, 12.7; CoH, 19.0*; TUG 1736-5: CrH, 75.0; CaH, 14.0; CaW, 12.1; CoH, 20.0*; TUG 1736-6: CrH, 70.0; CaH, 15.0; CaW, 7.0; TUG 1736-7: CrH, 59.0*; CaH, 14.0; CaW, 11.5; CoH, 20.0* (note the calyx width was calculated as described above).

Remarks.—As noted above, the preservation style of the Kalana Lagerstätte crinoids resulted in flattening of the calyces, making dimensions of the calyx other than that of the living organism. The true diameter was calculated assuming that the preserved diameter was one-half the circumference, and the true diameter was calculated with the standard geometric formula.

The proximal calyx of the Kalana specimens is rarely visible on specimens that were split during collection. Hence, the presence and characteristic of the infrabasal circling are only known on a few specimens. The part of the holotype (TUG 1736-3A) exposes a corner of an infrabasal below two basal plates, if the specimen is viewed with a very low-angle highlight (Fig. 4.1, 4.2). Also, TUG 1736-8 is a specimen that combines calcite and moldic/carbon impression preservation. In this specimen, plates of the calyx are preserved in cross section in calcite and reveal that a plate circling (the infrabasal circling) is present along the base of the calyx (Fig. 3.3).

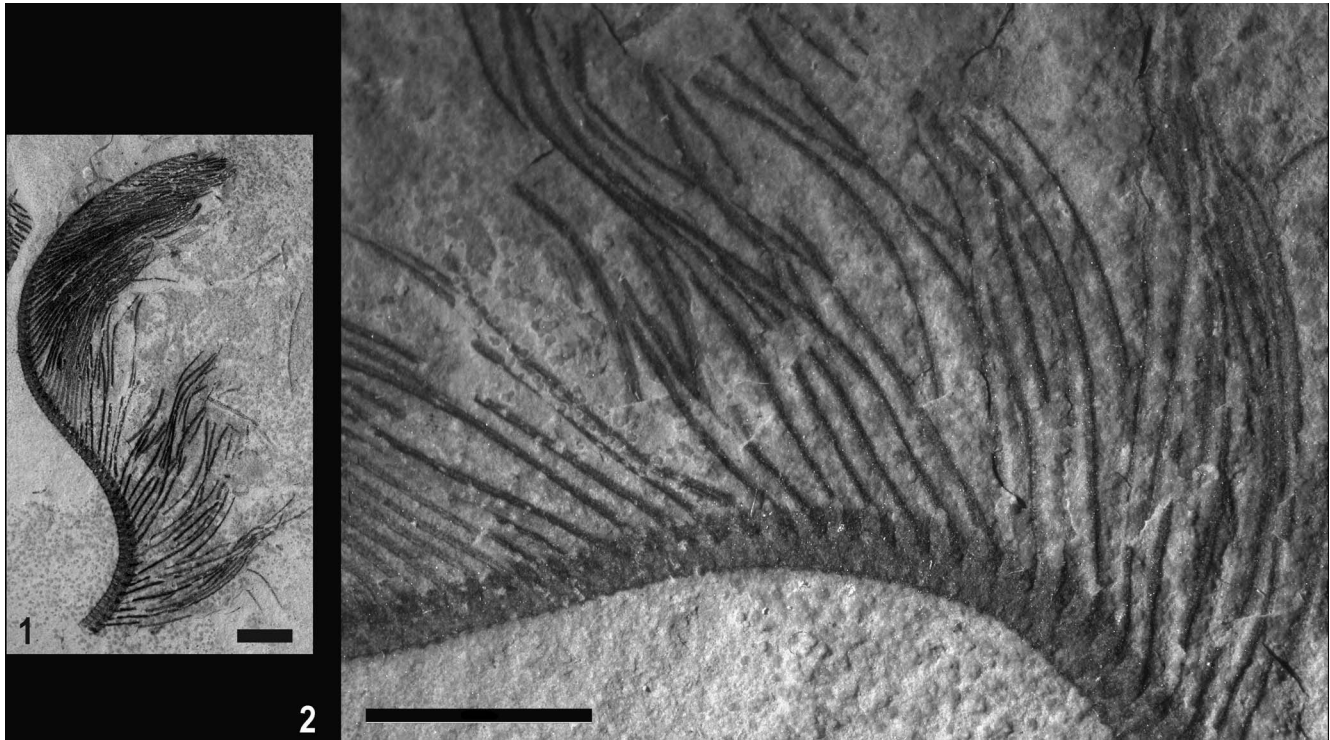


Figure 9. (1) Isolated arm of *Kalanacrinus mastikae* n. gen. n. sp., paratype, TUG 1736-10; (2) enlargement of proximal portion of arm; all scale bars 5.0 mm.

A few specimens preserve the plating in normal interradial regions. In TUG 1736-6A, the plates are slightly disarticulated and clearly demonstrate that the first interradial plate is sutured above shoulders of the two radial plates and that the second range consists of two plates (Figs. 5.2, 6, 7). The details of CD interray plating are not known. However, specimen TUG 1736-9 exposes an interray that is clearly wider than those on typical specimens, and that the radial plate positions are too far apart for them to be in lateral contact. This is interpreted to be evidence that the radial circlet is interrupted in the CD interray.

Subclass Pentacrinoidea Jaekel, 1918

Infraclass Inadunata Wachsmuth and Springer, 1885

Parvclass Disparida Moore and Laudon, 1943

Order Myelodactylida Ausich, 1998

Family Iocrinidae Moore and Laudon, 1943

Tartucrinus new genus

Type species.—*Tartucrinus kalanaensis*, n. gen. n. sp.

Diagnosis.—As for the type species by monotypy.

Occurrence.—Silurian (Llandovery, Aeronian); Estonia (Baltica paleocontinent).

Etymology.—This new crinoid is named for Tartu, the second largest city in Estonia and home of the University of Tartu.

Remarks.—GIT 405-257, GIT 405-258, and GUT 1736-20 each preserve two radial plates, and TUG 1736-20 has a single radial plate preserved. Some of the radial plates are complete and some are damaged. In all cases, these are simple radial plates. Where preserved, the first primibrachial is always a normal

brachial plate. What is known about the radial circlet is: (1) the CD interray is not exposed on any specimens (or is unrecognizable); (2) the iocrinid condition, with a C-ray suprarradial plate above the level of other radial plates, is either not exposed or absent; and (3) all three specimens with two radials preserved have two adjacent simple radial plates. This latter point is significant because the condition with two adjacent simple radial plates is present in only a limited number of Ordovician and Silurian disparids. Furthermore, if derived families with distinctive attributes (e.g., Calceocrinidae, Zophocrinidae) are eliminated, the only families with two adjacent simple radial plates are the Cincinnaticrinidae (Ordovician, 11212 radial circlet configuration), Columbocrinidae (Columbocrinidae (Ordovician, 11212), Iocrinidae (Ordovician–Silurian, 11211), Maennilicrinidae (Ordovician, 11111), and Tornatilicrinidae (Ordovician, 11211). Of these families, only the Iocrinidae and Cincinnaticrinidae have taxa with distinctive depressions along the aboral cup sutures; and most Cincinnaticrinidae have an aboral cup that is proportionally much higher than iocrinids and *Tartucrinus* n. gen. Further, taxa in the five families noted above nearly all have plenary radial facets, except iocrinids, of which *Iocrinus* Hall, 1866, *Muicrinus* Lin et al., 2018, *Schaldichocrinus* Rozhnov, 1997, and *Westheadocrinus* Donovan, 1989 have penepenary radial facets. Finally, and certainly not an absolute criterion, of the families considered, only the Iocrinidae ranges into the Silurian. Consequently, by the process of elimination, we regard *Tartucrinus* n. gen. as a member of the Iocrinidae, recognizing that this assignment is subject to verification when the morphology of *T. kalanaensis* n. gen. n. sp. can be more fully described.

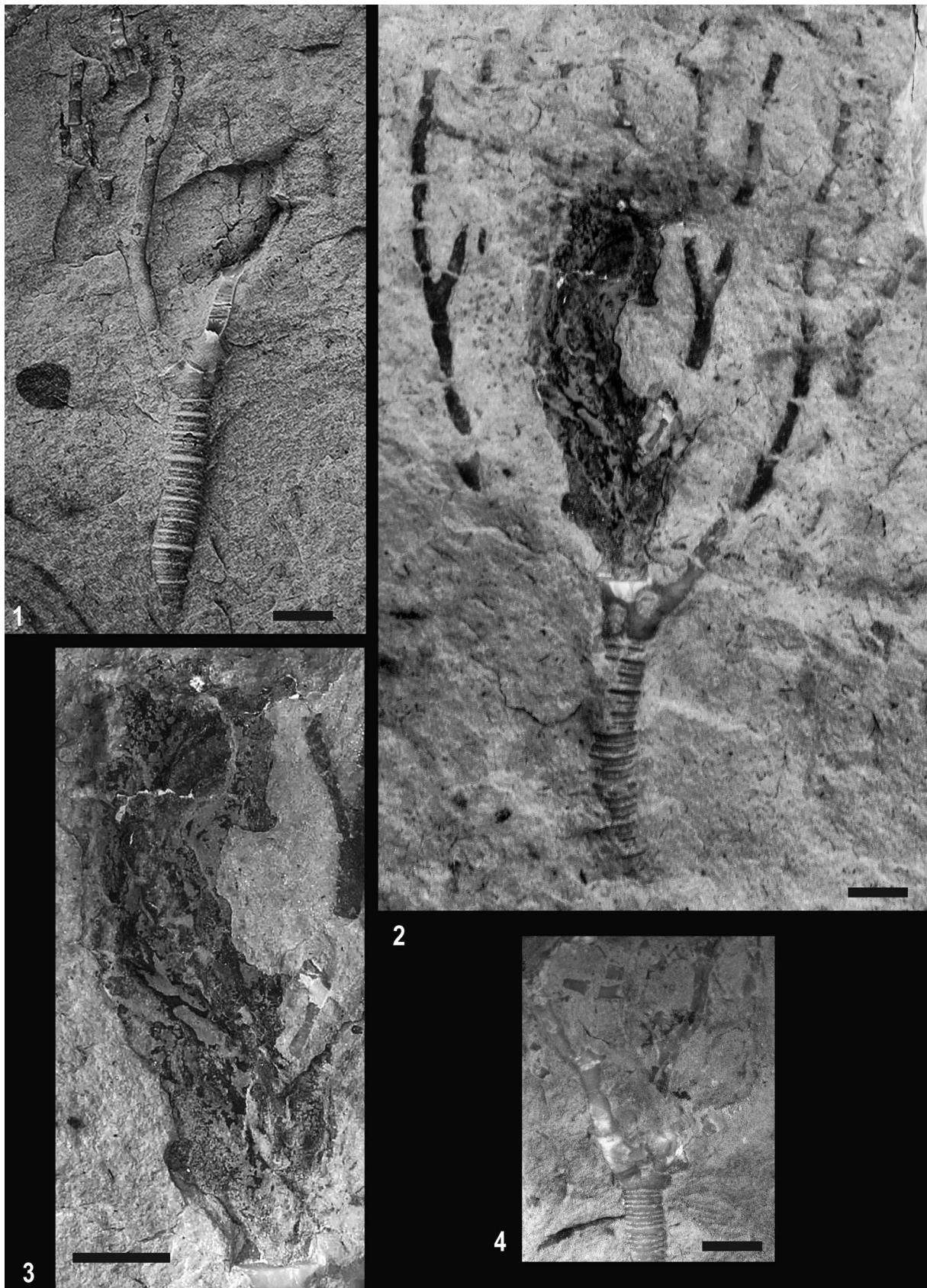


Figure 10. *Tartucrinus kalanaensis* n. gen. n. sp., note all crown plates preserved with calcite. (1) Lateral view of partial crown, specimen coated with ammonium chloride, holotype, GIT 405-257; (2, 3) lateral view; paratype, TUG 1372-20; (2) partial crown, note carbon film preserved between arms that may be the remnants of a non-calcified anal sac; (3) enlargement of carbon film between arms; scale bar 2.5 mm; (4) lateral view of calyx, paratype, GIT 405-254; all scale bars, except (3), 5.0 mm.

Within the Iocrinidae, *Tartuocrinus* n. gen. is unique with the following combination of characters: angustary radial facets, three primibrachials, two arm bifurcations, and a holomeric column (Supplemental Table 1). It is most similar to *Pariocrinus heterodactylus* Eckert, 1984, from the early Silurian (Llandovery, Rhuddanian) of North America. *Pariocrinus* has basal plates visible in lateral view, plenary radial facets, smooth anal sac plates, four or five primibrachials, at least three arm bifurcations, fixed interradsial plates absent, pentagonal to circular column, pentameric column, and a column spiral absent. In contrast, *Tartuocrinus* n. gen. has basal plates visible in lateral view, angustary radial facets, character of anal sac plates unknown, three primibrachials, two arm bifurcations, fixed interradsial plates absent, circular column, holomeric column, and a column spiral absent. *Tartuocrinus* n. gen. is compared to all genera in the Iocrinidae in Supplemental Table 1.

Tartuocrinus kalanaensis new species

Figure 10

Holotype.—GIT 405-257.

Diagnosis.—Iocrinid with basal cirlet visible in lateral view, angustary radial facets, (character of anal sac plates unknown), three primibrachials, two arm bifurcations, fixed interradsial plates absent, column circular, holomeric, and not spiraled (columnal facets in mesistele unknown).

Occurrence.—*Tartuocrinus kalanaensis* n. gen. n. sp. is only known from the Kalana Lagerstätte in the Kalana Quarry, near the village of Kalana, Jõgeva County, central Estonia; Silurian, Llandovery, middle Aeronian.

Description.—Crown small in size, subovate (Fig. 10.2). Aboral cup small, low cone in shape, maximum width 1.7 times wider than high; plates very convex with deep depressions, as noted below; other plate sculpturing smooth.

Basal cirlet ~40% of aboral cup height (less at basal plate-basal plate suture) (Fig. 10.1); basal plate width 2.0 times wider than high, deep depression at radial plate-radial plate-basal plate triple junction, elongate depression along basal plate-basal plate suture. Radial cirlet ~60% of aboral cup height (although much more through midline of radial plate), strongly convex proximal to distal middle of plate that continues to brachials, also strongly convex to form a broad ridge to adjoining radial plates; radial plates larger than basal plates. Radial facet angustary, ~40% of distal radial plate width; declivate (Fig. 10.4).

Number of basal plates, number of compound versus simple radial plates, posterior interray plating, and oral surface unknown.

Presumably five free arms, branch twice with isotomous divisions (as known) (Fig. 10.2); primibrachials 1.3 times higher than wide, strongly convex, third primibrachial axillary (Fig. 10.1); secundibrachials 2.5 times higher than wide, strongly convex, in single example second to fifth secundibrachial axillary; tertibrachials unbranched through as many as eight brachials. Arms apinnulate.

Column circular, holomeric, heteromorphic with N3231323 pattern; columnals low, all columnals have a convex

latus, but height and width of columnal and degree of convexity are greatest in nodals and become less with each subsequent internodal cycle (Fig. 10.1).

Materials.—Holotype: GIT 405-257; Paratypes: GIT 405-258, GIT 405-254B–GIT 405-254B, TUG 1372-20, and TUG 1736-20.

Measurements.—Holotype: GIT 405-257: CrH, 14.0*; CaH, 2.9; CaW, 4.3; CoH, 18.0*; Paratypes: GIT 405-258: CrH 50.0*; CaW, 10.0; CoH, 20.0*; TUG 1372-20: CrH, 48.0; CaH, 5.0; CaW, 9.0; CoH, 18.0*.

Remarks.—Four incomplete specimens of *Tartuocrinus kalanaensis* n. gen. n. sp. are known from the Kalana Lagerstätte. All expose a partial aboral cup, a few to several arms (Fig. 10.2, 10.3), and a length of column. Although strong convexity of aboral cup plates with depressions at the plate triple junctions and along the basal plate-basal plate, sutures are reminiscent of some Ordovician camerates (e.g., *Reteocrinus* and *Xenocrinus*), this Estonian taxon is monocyclic; the arms lack pinnules; and there are no fixed interradsial plates, all of which identify this as a disparid.

Despite the fact that Kalana Lagerstätte specimens of this species are preserved with 3D calcite, specimen TUG 1376-20 has an organic residue within the enclosed arms. There is no plating associated with this organic residue, so one can only speculate on its origin. Possibilities included a non-calcified anal sac (Kammer and Ausich, 2007), soft tissue from the ambulacra, or a combination of both.

Accessibility of supplemental data

Color images of Kalana crinoids and diagnostic tables are available from the Dryad Digital Repository: <https://doi.org/10.5061/dryad.68pv736>.

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References

- Ainsaar, L., Tinn, O., Männik, P., and Meidla, T., 2014, Stop B1: Kalana quarry, in Bauert, H., Hints, O., Meidla, T., and Männik, P., eds., 4th Annual Meeting of IGCP 591, The Early to Middle Paleozoic Revolution: Estonia, 10–19 June 2014, Abstracts and Field Guide, Tartu, p. 174–177.
- Ahmed, E., and Holmström, S.J.M., 2015, Microbe-mineral interactions: the impact of surface attachment on mineral weathering and element selectivity by microorganisms: *Chemical Biology*, v. 403, p. 13–23.
- Angelin, N.P., 1878, *Iconographia Crinoideorum in stratis Sueciae Siluricis fossilium*: Holmia, Samson and Wallin, 62 p.
- Ausich, W.I., 1984a, Calceocrinids from the Early Silurian (Llandoveryan) Brassfield Formation of southwestern Ohio: *Journal of Paleontology*, v. 58, p. 1167–1185.

- Ausich, W.I., 1984b, The genus *Clidochirus* from the Early Silurian of Ohio (Crinoidea, Llandoveryan): *Journal of Paleontology*, v. 58, p. 1341–1346.
- Ausich, W.I., 1985, New crinoids and revision of the superfamily Glyptocrinoidea (Early Silurian, Ohio): *Journal of Paleontology*, v. 59, p. 793–808.
- Ausich, W.I., 1986a, Early Silurian rhodocriniteacean crinoids (Brassfield Formation, Ohio): *Journal of Paleontology*, v. 60, p. 84–106.
- Ausich, W.I., 1986b, Early Silurian inadunate crinoids (Brassfield Formation, Ohio): *Journal of Paleontology*, v. 60, p. 719–735.
- Ausich, W.I., 1986c, New camerate crinoids of the Suborder Glyptocrinina from the Lower Silurian Brassfield Formation (southwestern Ohio): *Journal of Paleontology*, v. 60, p. 887–897.
- Ausich, W.I., 1987, Brassfield Compsocrinina (Lower Silurian crinoids) from Ohio: *Journal of Paleontology*, v. 61, p. 552–562.
- Ausich, W.I., 1998, Phylogeny of Arenig to Caradoc crinoids (Phylum Echinodermata) and suprageneric classification of the Crinoidea: *The University of Kansas Paleontological Contributions, New Series*, no. 9, 36 p.
- Ausich, W.I., 2018, Morphological paradox of disparid crinoids (Echinodermata): phylogenetic analysis of a Paleozoic clade: *Swiss Journal of Paleontology*, v. 137, p. 159–176.
- Ausich, W.I., and Copper, P., 2010, The Crinoidea of Anticosti Island, Québec (Late Ordovician to Early Silurian): *Palaeontographica Canadiana*, no. 29, 157 p.
- Ausich, W.I., and Deline, B., 2012, Macroevolutionary transitions in crinoids following the Late Ordovician extinction event (Ordovician to Early Silurian): *Palaeogeography, Palaeoclimatology, Palaeoecology*, v. 361–362, p. 38–48.
- Ausich, W. I., and Wilson, M.A., 2016, Llandovery (early Silurian) crinoids from Hiiumaa Island, western Estonia: *Journal of Paleontology*, v. 90, p. 1139–1147. <http://dx.doi.org/10.1017/jpa.2016.120>
- Ausich, W.I., Kammer, T.W., and Baumiller, T.K., 1994, Demise of the middle Paleozoic crinoid fauna: a single extinction event or rapid faunal turnover?: *Paleobiology*, v. 20, no. 3, p. 345–361.
- Ausich, W.I., Brett, C.E., Hess, H., and Simms, M.J., 1999, Crinoid form and function, in Hess, H., Ausich, W.I., Brett, C.E., and Simms, M.J., *Fossil Crinoids*: Cambridge, UK, Cambridge University Press, p. 3–30.
- Ausich, W.I., Wilson, M.A., and Vinn, O., 2012, Crinoids from the Silurian of western Estonia (Phylum Echinodermata): *Acta Palaeontologica Polonica*, v. 57, no. 3, p. 613–631.
- Ausich, W.I., Wilson, M.A., and Vinn, O., 2015, Wenlock and Pridoli (Silurian) crinoids from Saaremaa, western Estonia (Phylum Echinodermata): *Journal of Paleontology*, v. 89, p. 72–81.
- Bather, F.A., 1893, The Crinoidea of Gotland. Pt. 1, The Crinoidea Inadunata: *Kongliga Svenska Vetenskaps-Akademiens Handlingar*, v. 25, no. 2, 200 p.
- Baumiller, T.K., 1994, Patterns of dominance and extinction in the record of Paleozoic crinoids, in David, B., Guille, A., Féral, J. P., and Roux, M., eds., *Echinoderms through Time (Echinoderms Dijon)*: Rotterdam, A.A. Balkema, p. 193–198.
- Bogolepova, O.K., Donovan, S.K., Harper, D.A.T., Suyarkova, A.A., Yakupov, R., and Gubanov, A.P., 2018, New records of brachiopods and crinoids from the Silurian (Wenlock) of the southern Urals, Russia: *GFF*, v. 140, p. 323–331.
- Briggs, D.E.G., and McMahon, S., 2016, The role of experiments in investigating the taphonomy of exceptional preservation: *Palaeontology*, v. 59, p. 1–11.
- Cole, S.R., 2017, Phylogeny and morphological evolution of the Ordovician Camerata (Class Crinoidea, Phylum Echinodermata): *Journal of Paleontology*, v. 91, p. 815–828.
- Dong, H., 2010, Mineral-microbe interactions: a review: *Frontiers of Earth Science*, v. 4, p. 127–147.
- Donovan, S.K., 1989, *Pelmatozoan columnals from the Ordovician of the British Isles, Part 2*: *Palaeontographical Society Monograph*, v. 142, no. 580, p. 69–114.
- Donovan, S.K., 1993, A Rhuddanian (Silurian Llandovery) pelmatozoan fauna from south-west Wales: *Geological Journal*, v. 28, p. 1–19.
- Donovan, S.K., 1994, The Late Ordovician extinction of the crinoids in Britain: *National Geographic Research and Exploration*, v. 10, p. 72–79.
- Eckert, J.D., 1984, Early Llandovery crinoids and stelleroids from the Cataract Group (Lower Silurian), southern Ontario, Canada: *Royal Ontario Museum Life Sciences, Contributions*, no. 137, 83 p.
- Eckert, J.D., 1988, Late Ordovician extinction of North American and British crinoids: *Lethaia*, v. 21, p. 147–167.
- Erdtmann, B.-D., and Prezbindowski, D.R., 1974, Niagaran (Middle Silurian) interreef fossil burial environments in Indiana: *Neues Jahrbuch für Geologie und Paläontologie Abhandlungen*, v. 144, p. 342–372.
- Fearnhead, F.E., and Donovan, S.K., 2007, A cladid crinoid (Echinodermata) from the Llandovery (Lower Silurian) of the Girvan district, SW Scotland: *Scottish Journal of Geology*, v. 43, p. 74–82.
- Foerste, A.F., 1923, Notes on Medinan, Niagaran, and Chester fossils: *Denison University Bulletin*, v. 20, p. 37–210.
- Franzén, C., 1982, A Silurian crinoid thanatotope from Gotland: *Geologiska Föreningens i Stockholm Förhandlingar*, v. 103, p. 469–490.
- Franzén, C., 1983, Ecology and taxonomy of Silurian crinoids from Gotland: *Acta Universitatis Upsaliensis, Abstracts of Uppsala Dissertations from the Faculty of Science*, no. 665, 31 p.
- Hall, J., 1866, Descriptions of new species of Crinoidea and other fossils from the Lower Silurian strata of the age of the Hudson-River Group and Trenton Limestone: Albany, New York, privately published preprint, 17 p.
- Kammer, T.W., and Ausich, W.I., 2007, Soft-tissue preservation of the hind gut in a new genus of cladid crinoid from the Mississippian (Visean, Asbian) at St Andrews, Scotland: *Palaeontology*, v. 50, p. 951–955.
- Jaekel, O., 1918, *Phylogenie und System der Pelmatozoen*: *Paläontologische Zeitschrift*, v. 3, p. 1–128.
- Kluessendorf, J., 1994, Predictability of Silurian Fossil-Konservat-Lagerstätten in North America: *Lethaia*, v. 27, p. 337–344.
- Lane, N.G., and Ausich, W.I., 1995, Interreef crinoid fauna from the Mississippian Shale Member of the Wabash Formation (northern Indiana; Silurian; Echinodermata): *Journal of Paleontology*, v. 69, p. 1090–1106.
- Lin, J.-P., Ausich, W.I., Baliński, A., Bergström, S.M., and Sun, Y., 2018, The oldest iocrinid crinoids from the Early/Middle Ordovician of China: possible paleogeographic implications. *Journal of Asian Earth Sciences*, v. 151, p. 324–333.
- Lo Duca, S.T., 1990, *Medusagraptus mirabilis* Ruedemann as a noncalcified dasyclad alga: *Journal of Paleontology*, v. 64, p. 469–474.
- Lo Duca, S.T., and Brett, C.E., 1997, *The Medusaeagraptus* epibole and lower Ludlow Konservat-Lagerstätten of eastern North America, in Brett, C.E., and Baird, G.C., eds., *Paleontological Events Stratigraphic, Ecological, and Evolutionary Implications*: New York, Columbia University Press, p. 369–406.
- Männik, P., Tinn, O., Loydell, D.K., and Ainsaar, L., 2016, Age of the Kalana Lagerstätte, early Silurian, Estonia: *Estonian Journal of Earth Sciences*, 65, p. 105–114.
- Mastik, V., and Tinn, O., 2015, New dasycladalean algal species from the Kalana Lagerstätte (Silurian, Estonia): *Journal of Paleontology*, v. 89, p. 262–268.
- Mastik, V., and Tinn, O., 2017, *Leveilleites hartmageli* Foerste, 1923 (Rhodophyta?) from the Ordovician of Laurentia and Silurian of Baltica: redescription and designation of a neotype: *Palaeoworld*, v. 26, p. 602–611.
- Miller, J.S., 1821, *A Natural History of the Crinoidea or Lily-shaped Animals, with Observations on the Genera Asteria, Euryale, Comatula, and Marsupites*: Bristol, Bryan and Co., 150 p.
- Moore, R.C., and Laudon, L.R., 1943, Evolution and classification of Paleozoic crinoids: *Geological Society of America Special Paper* 46, 153 p.
- Moore, R.C., and Teichert, K., eds., 1978, *Treatise on Invertebrate Paleontology, Part T, Echinodermata 2, Crinoidea*: Boulder, Colorado, and Lawrence, Kansas, Geological Society of America, and University of Kansas, 1027 p.
- Peters, S., and Ausich, W.I., 2008, A sampling-adjusted macroevolutionary history for Ordovician–Early Silurian crinoids: *Paleobiology*, v. 34, p. 104–116.
- Phillips, J., 1839, Chapter 48. Encrinites and zoophytes of the Silurian System, in Murchison, R.I., *The Silurian System*: London, John Murray, p. 670–675.
- Rozhnov, S.V., 1997, New generic name *Schaldichocrinus* Rozhnov, nom. nov.: *Paleontological Journal*, v. 31, p. 437.
- Saunders, K.M., Bates, D.E.B., Kluessendorf, J., Loydell, D.K., and Mikulic, D.G., 2009, *Desmograptus micromenatodes*, a Silurian dendroid graptolite, and its ultrastructure: *Palaeontology*, v. 52, p. 541–559.
- Tinn, O., and Märss, T., 2018, The earliest osteostracan *Kalanaspis delectabilis* gen. et sp. nov. from the mid-Aeronian (mid-Llandovery, lower Silurian) of Estonia: *Journal of Vertebrate Paleontology*, v. 38, e1425212. doi: 10.1080/02724634.2017.1425212
- Tinn, O., Meidla, T., Ainsaar, L., and Pani, I., 2009, Thallophtic algal flora from a new Silurian Lagerstätte: *Estonian Journal of Earth Sciences*, v. 58, p. 38–42.
- Tinn, O., Mastik, V., Ainsaar, L., Meidla, T., 2015, *Kalana pusilla*, an exceptionally preserved non-calcified alga from the lower Silurian (Aeronian, Llandovery) of Estonia: *Palaeoworld*, v. 24, p. 207–214.
- Ubahgs, G., 1978, General morphology, in Moore, R.C., and Teichert, K., eds., *Treatise on Invertebrate Paleontology, Part T, Echinodermata*: Boulder, Colorado, and Lawrence, Kansas, Geological Society of America and University of Kansas, v. 2, p. T58–T216.
- Uroz, S., Calvaruzo, C., Turpault, M.-P., and Frey-Klett, P., 2009, Mineral weathering by bacteria: ecology, actors and mechanisms: *Trends in Microbiology*, v. 17, p. 378–387.
- Wachsmuth, C., and Springer, F., 1880–1886, *Revision of the Palaeocrinoidea. Pt. I. The families Ichthyocrinidae and Cyathocrinidae (1880)*, p. 226–378, (separate repaginated p. 1–153). Pt. II. Family Sphaeroidocrinidae, with the sub-families Platycrinidae, Rhodocrinidae, and Actinocrinidae (1881), p. 177–411, (separate repaginated, p. 1–237). Pt. III, Sec. 1. Discussion of

- the classification and relations of the brachiate crinoids, and conclusion of the generic descriptions (1885), p. 225–364, (separate repaginated, p. 1–138). Pt. III, Sec. 2. Discussion of the classification and relations of the brachiate crinoids, and conclusion of the generic descriptions (1886), p. 64–226 (separate repaginated, p. 139–302); Proceedings of the Academy of Natural Sciences of Philadelphia.
- Webster, G.D., and Webster, D.W., 2013, Bibliography and index of Paleozoic crinoids, coronoids, and hemistreptocrinids, 1758–2012. <http://crinoids.azurewebsites.net/> (accessed 12 October, 2018).
- Wright, D.F., 2017, Bayesian estimation of fossil phylogenies and the evolution of early to middle Paleozoic crinoids (Echinodermata): *Journal of Paleontology*, v. 91, p. 799–814.
- Wright, D.F., Ausich, W.I., Cole, S.R., Peter, M.E., and Rhenberg, E.C., 2017. Phylogenetic taxonomy and classification of the Crinoidea (Echinodermata). *Journal of Paleontology*, v. 91, p. 829–846. doi: <https://doi.org/10.1017/jpa.2016.142>; published online 02-22-17.
- Zamora, S., Rahman, I.A., and Ausich, W.I., 2015, A new iocrinid (Disparida) from the Ordovician (Darriwilian) of Morocco: *PeerJ*, 3:e1450, 10 p. doi: 10.7717/peerj.1450.
- Zittel, K.A. von, 1876–1880. *Handbuch der Palaeontologie*, v. 1, Palaeozoologie: München, Leipzig, R. Oldenbourg (1879), no. 1, 765 p.

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