

A new cobia (Teleostei, Rachycentridae) species from the Miocene St. Marys Formation along Calvert Cliffs, Maryland, USA

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Abstract.—The highly fossiliferous St. Marys Formation is exposed along Calvert Cliffs, Maryland, and comprises one of the best available records of late Miocene life in the northeastern United States. *Rachycentron stremphaencus* new species, a cobia from the late Miocene (Tortonian) of the St. Marys Formation is described herein on the basis of a single three-dimensional neurocranium. This fossil represents the earliest known occurrence of neurocranial remains of the genus *Rachycentron* in the record. *Rachycentron stremphaencus* differs from *Rachycentron canadum* (Linnaeus, 1766) in many ways. The most obvious include a different ornamentation of the outer surface of the cranial bones; a notably pronounced lateral ridge resulting in a considerable gradient from the dorsal-medial exposure of the frontal to its lateralmost supraorbital margin; the size, shape, and position of the sphenotic that is located in the posterior half of the neurocranium and its lateralmost edge being adjacent to the anteriormost extent of the wedge-shaped trough in the dorsal surface of the skull formed by the lateral and medial ridges; the two contralateral medial ridges forming a proportionately much wider trough on either side of the supraoccipital; the epioccipitals not reaching the rear edge of the neurocranium; and the lack of a conspicuous posterolateral prong of the intercalar.

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Introduction

The cobia *Rachycentron canadum* (Linnaeus, 1766) is an extant carangiform fish, the sole representative of the monotypic family Rachycentridae (Nelson, 2006; Sajeevan and Madhusoodana Kurup, 2014). Cobias are marine teleosts that occur nearly worldwide in a variety of tropical, subtropical, and warm temperate habitats and occasionally enter estuaries and other brackish-water environments feeding on crustaceans, cephalopods, and other fishes (Shaffer and Nakamura, 1989). Other than the review by Carnevale and Godfrey (2018) of a single partially complete specimen from the Calvert Formation referred to *Rachycentron* sp., the genus apparently has no fossil record based on skeletal remains. Neither is it known on the basis of fossil otoliths (Nolf, 2013). Here, we describe a new fossil species of the genus *Rachycentron* from the Tortonian of the St. Marys Formation, Calvert Cliffs, Calvert County, Maryland, USA, on the basis of a single nearly complete neurocranium. This fossil therefore represents the second known record of the family Rachycentridae, thereby providing relevant information about the evolutionary history of these fishes.

Geology

The specimen derives from the Little Cove Point Member of the Miocene St. Marys Formation from north of Driftwood Beach along Calvert Cliffs, Maryland, USA. The Miocene deposits of the Atlantic Coastal Plain of North America include the

siliciclastic sediments of the St. Marys Formation. These fossiliferous shallow marine deposits constitute the uppermost portion of the Neogene Chesapeake Group as exposed in the Calvert Cliffs along the Chesapeake Bay shoreline in Calvert County (Shattuck, 1904; Kidwell, 1984; Kidwell et al., 2015). The deposits of the St. Marys Formation are located along the western margin of the Salisbury Embayment, a paleogeographic domain that was occupied by a shallow arm of the Atlantic Ocean during the Miocene (Ward and Andrews, 2008), structurally representing a westward extension of the Baltimore Canyon Trough (see Poag, 1979; Kidwell, 1984; Kidwell et al., 2015). Overall, the Miocene succession of the Chesapeake Group records a gradual shallowing within the Salisbury Embayment (e.g., Kidwell, 1984, 1988, 1989, 1997; Ward and Strickland, 1985; Ward, 1992; Ward and Andrews, 2008). In this context, the deposits of the St. Marys Formation (specifically, the Little Cove Point Member) are quite different from those of the Calvert and Choptank formations, which are the other two Miocene formations making up Calvert Cliffs. Sediments of this St. Marys Formation member show more rapid lateral (and vertical) facies changes and include facies from shallower water and more tide-dominated environments, reflecting a transition from open marine to tidally influenced low salinity muddy embayments (Kidwell, 1988, 1997; Kidwell et al., 2015). The age of the Little Cove Point Member has been estimated at 9–10 Ma (Perez et al., 2019, fig. 1), which is entirely within the Tortonian stage.

The St. Marys Formation has yielded a rich record of marine and terrestrial organisms (Vogt and Eshelman, 1987; Carnevale

and Godfrey, 2018; Godfrey, 2018). A large number of predominantly marine species have been described from these deposits, with representatives from several groups of microfossils and invertebrates as well as an array of vertebrate taxa, including sharks and rays, actinopterygian fishes, turtles, crocodiles, pelagic birds, seals, sea cows, odontocete and mysticete cetaceans, and rare isolated remains of large terrestrial mammals (e.g., Clark et al., 1904; Ward and Andrews, 2008; Carnevale et al., 2011; Godfrey, 2018).

Remains of actinopterygian fishes are relatively common in the St. Marys Formation, consisting of isolated bones, partially articulated skeletons, teeth, and otoliths (Carnevale and Godfrey, 2018). However, until recently, our knowledge of the ichthyofaunal assemblage was limited to short descriptions of isolated remains incidentally collected from several localities in Maryland and Virginia (Cope, 1867, 1869; Leidy, 1873; Eastman, 1904; Smith, 1909; Lynn and McLelland, 1939; Blake, 1940; Leriche, 1942; Dante, 1953; Kimmel and Purdy, 1984; Weems, 1985). More recently, Müller (1999) provided a comprehensive overview of the otolith assemblage. Finally, on the basis of the analysis of the skeletal remains in the collections of the Calvert Marine Museum, Solomons, Maryland, and the Department of Paleobiology, National Museum of Natural History, Smithsonian Institution, Washington, DC, Carnevale and Godfrey (2018) presented a broad overview of the fish diversity of the successions of the Chesapeake Group, recognizing on the basis of skeletal elements 19 genera from the St. Marys Formation. Thirty-six genera were reported from the St. Marys Formation on the basis of the presence of otoliths (Müller, 1999; Carnevale and Godfrey, 2018). Carnevale and Godfrey (2018, fig. 3.8) also described a set of associated bones (CMM-V-3694), (sclerotic ring, right premaxilla, left maxilla, left suspensorium and preopercle, right second pharyngobranchial, right posttemporal, vertebrae) from the Calvert Formation that were referred to an indeterminate species of the genus *Rachycentron*.

Materials and methods

The nearly complete neurocranial remains were collected on 27 and 28 April 2019 from the St. Marys Formation outcropping along Calvert Cliffs, Maryland. Detailed osteological examinations were made possible following the removal of the surrounding matrix. This was achieved using needles working under a stereomicroscope. Measurements were taken with a caliper to the nearest 0.1 mm. To improve contrast in the figures, the specimens were lightly dusted with sublimed ammonium chloride. After the specimens were photographed under fluorescent lighting, the ammonium chloride was removed by rinsing them under running water.

Repositories and institutional abbreviations.—Type, figured, and one other specimen examined in this study are deposited in the Calvert Marine Museum comparative osteology collection (CMM-O) and vertebrate paleontology collection (CMM-V), Solomons, Maryland, USA.

Systematic paleontology

Division Percomorphacea Wiley and Johnson, 2010
Order Carangiformes Jordan, 1923

Family Rachycentridae Gill, 1895
Genus *Rachycentron* Kaup, 1826

Type species.—*Rachycentron canadum* (Linnaeus, 1766).

Rachycentron stremphaencus new species
Figures 1, 2, 3.2, 3.4

Holotype.—CMM-V-10,000, relatively complete neurocranium lacking the ethmoid portion and the anterior half of the basicranium (Figs. 1, 2, 3.2, 3.4).

Diagnosis.—A *Rachycentron* species with outer surface of the bones of the skull roof smooth or scarcely ornamented; skull roof with two longitudinal bony ridges, a lateral ridge formed by the frontal and pterotic and a flat medial ridge formed by the frontal, parietal, and epioccipital bordering a shallow medial frontal depression; sphenotic located in the posterior half of the neurocranium and not extending laterally at the level of the pterotic; epioccipital not extending posteriorly to the level of the articular facet of the exoccipital; trough formed by the medial ridges on either side of the supraoccipital; intercalary prong apparently absent.

Occurrence.—Calvert Cliffs north of Driftwood Beach, Chesapeake Ranch Estates, Calvert County, Maryland, USA. Shattuck Zones 22–23, Little Cove Point Member of the St. Marys Formation, Tortonian Stage, late Miocene. GPS coordinates are approximately 38.3538591°N, 76.3929062°W.

Description.—Measurements (in millimeters): neurocranial length = 119.0; minimum neurocranial width across the frontals at the supraorbital margin of the orbit = 46.5; maximum neurocranial width across the pterotics = 67.0; neurocranial width across the epioccipitals = 39.0; neurocranial width across the exoccipitals = 22.5.

CMM-V-10,000 consists of a dorsoventrally compressed neurocranium (Figs. 1, 2). Overall, the lateral profile of the neurocranium is wedge shaped, deepest at the back of the skull. The width of the skull across the pterotics is approximately equal to 1.9 times the maximum height of the back of the neurocranium (which is 35 mm deep). There are two paired continuous bony ridges and a median supraoccipital crest (damaged in CMM-V-10,000) on the dorsal face of the neurocranium. The lateral ridge is formed by the frontal, sphenotic, and pterotic; the medial ridge is formed by the frontal, parietal, and epioccipital. The lateral ridge arises at the level of the anterior third of the orbit and continues posteriorly along the central margin of the pterotic, ending as the posterolateral corner of the neurocranium. The flattened medial ridge begins slightly anteriorly, runs longitudinally through the whole length of the frontal, and continues posteriorly on the parietal and epioccipital. In the posterior third of the neurocranium, a deep and oblong trough is bounded by the lateral and medial ridges. Another shorter trough is demarcated by the medial ridge and the supraoccipital crest.

The right lateral ethmoid is partially preserved in the specimen along the ventral surface of the neurocranium (Fig. 2); it is a wing-like flattened bone, almost quadrangular in shape, with a

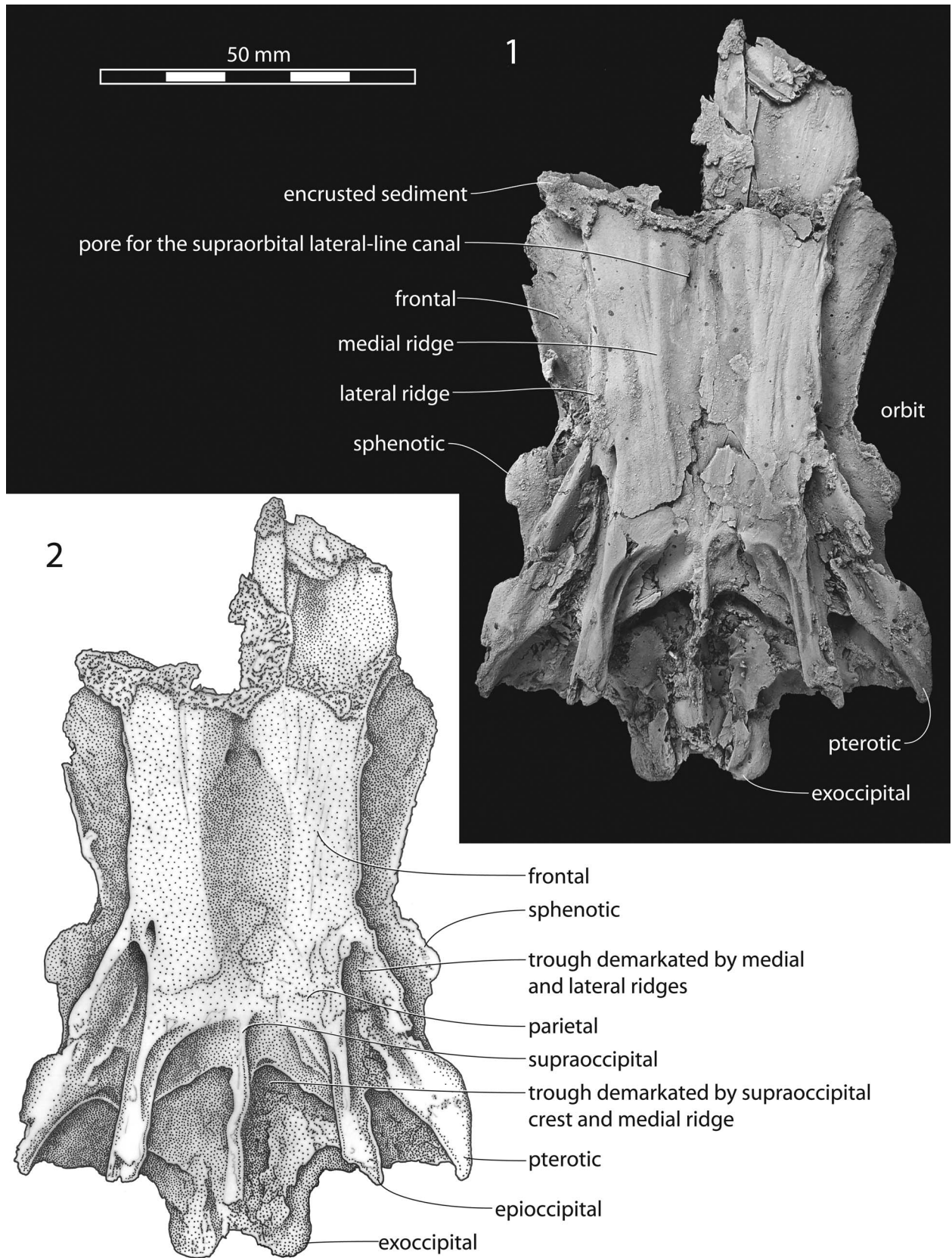


Figure 1. *Rachycentron stremphaencus* n. sp., CMM-V-10,000, neurocranium: (1) dorsal view; (2) interpretive drawing. Specimen dusted with sublimed ammonium chloride to improve contrast.

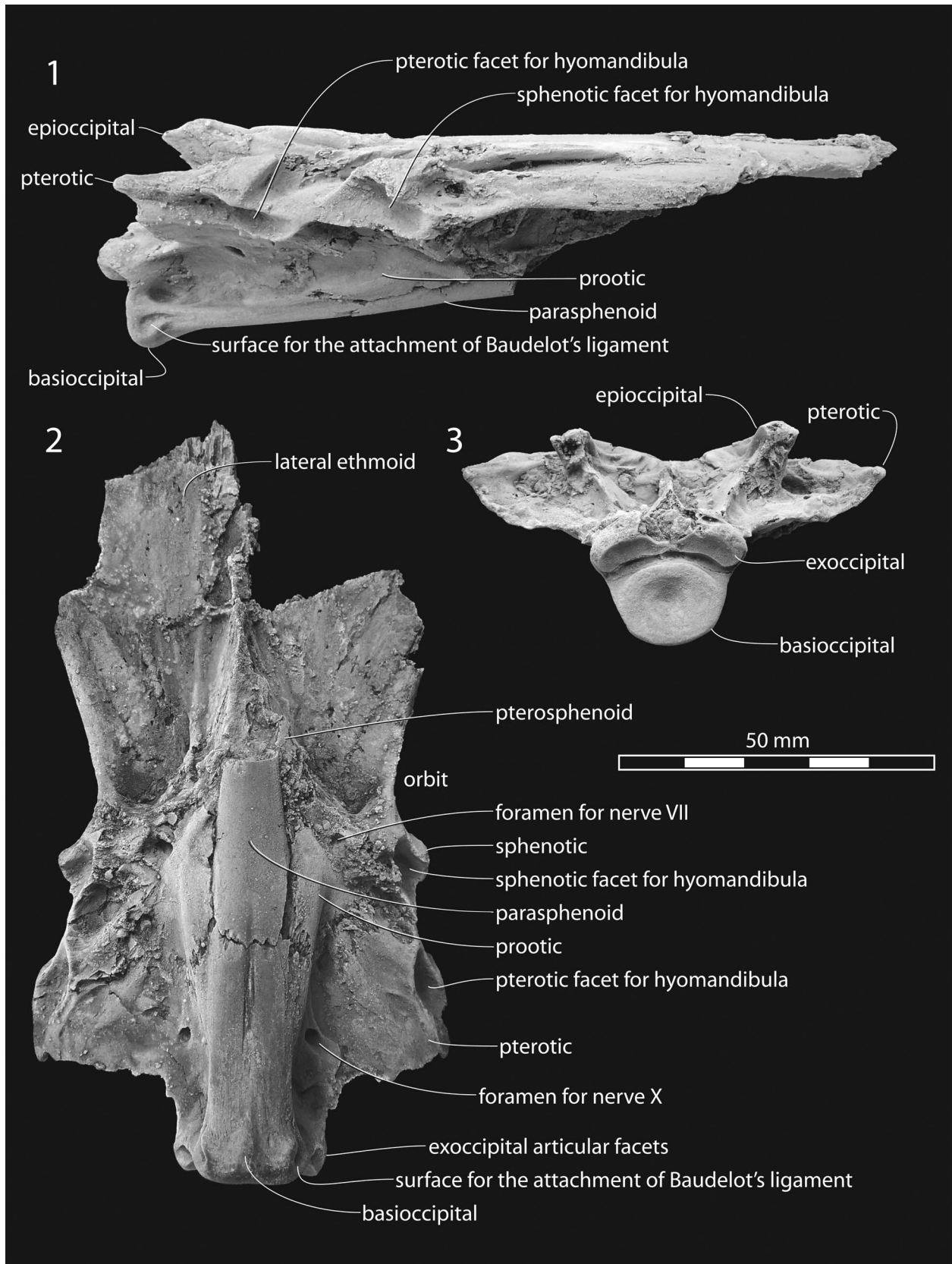


Figure 2. *Rachycentron stremphaencus* n. sp., CMM-V-10,000, neurocranium: (1) right lateral view; (2) ventral view; (3) posterior view. Specimen dusted with sublimed ammonium chloride to improve contrast.

smooth ventral surface. The lateral ethmoid articulates posteriorly with the frontal.

The frontals are the largest bones of the skull roof. They comprise more than half of the length of the preserved portion of the neurocranium and the full width of the neurocranium anteriorly. There is a shallow medial depression that reaches its maximum depth at the level of the suture of the two contralateral frontals. The contralateral pores of the supraorbital lateral-line canal can be observed in the anterior portion of this longitudinal depression. Lateral to the medial ridge, the lateral ridge stands above the wide supraorbital margin of the frontal, and a deep furrow separates the lateral ridge from the supraorbital margin of the frontal. In ventral view (Fig. 2.2), the frontal exhibits a relatively smooth and wide supraorbital surface. Each frontal articulates anteriorly with the lateral ethmoid, posterolaterally with the sphenotic and pterotic, posteriorly with the parietal and supraoccipital, and ventromedially with the pterosphenoïd.

The parietal occurs on the lateral side of the supraoccipital and articulates anteriorly with the frontal, medially with the supraoccipital, posteriorly with the epioccipital, and posterolaterally with the pterotic. The supraoccipital occupies the median posterodorsal portion of the skull roof and supports a shallow and relatively thin supraoccipital crest (Fig. 1); the supraoccipital crest is incomplete posteriorly, but in life, it would have risen above the exoccipitals.

From the posterolateral margin of the frontal, the tongue-shaped sphenotic projects laterally to form the dorsolateral margin of the orbit. In dorsal view, a ridge crosses the bone diagonally (in an anteromedial to posterolateral direction) to join with one on the pterotic. Medial to this ridge, the sphenotic preserves the anterior end of a wedge-shaped depression that broadens posteriorly onto the pterotic and is bounded medially by the bold ridge of the epioccipital. In lateral and ventral views, the sphenotic displays a posteroventrally directed deep cavity that represents the articular surface for the anterior dorsal head of the hyomandibula (Fig. 2.1, 2.2).

The sphenotic articulates posteriorly with the pterotic and ventrally with the prootic.

The pterotic forms the posterolateral corner of the neurocranium. In dorsal view, it bears a posterolaterally projecting tubular ridge; medial to this ridge, the surface of the bone is depressed into the longitudinal furrow that started on the sphenotic; ventrally, the pterotic has a deep articular surface for the posterior dorsal head of the hyomandibula. The pterotic articulates medially with the epioccipital and posteriorly with the exoccipital.

The two contralateral epioccipitals are separated from each other by the supraoccipital. Each epioccipital is robust, is expanded posteriorly, and forms the posterior portion of the medial ridge of the dorsal face of the neurocranium (Fig. 1); the epioccipital portion of the medial ridge splits the difference in distance between the supraoccipital crest and the lateral ridge on the pterotic.

The exoccipitals are visible in dorsal, ventral, and posterior views (Figs. 1, 2). Posteriorly, each exoccipital bears a shallow and medially oriented concave facet that overhangs the basioccipital for the articulation with the fovea of the first vertebra. Ventrally, each exoccipital shows a large foramen for the nervus vagus. The basioccipital occupies the posteroventral portion of

the neurocranium. In posterior view, the basioccipital is concave and subcircular in outline. The attaching surface for Baudelot's ligament is well recognizable ventrolaterally in the posterior sector of this bone. There is a 2 mm wide trough that separates the basioccipital from the exoccipitals (Fig. 2.3).

In a ventral view of CMM-V-10,000, the parasphenoid is deeply sutured to the basioccipital, showing no clear physical separation between these two bones. The elongate parasphenoid is unremarkable in its smooth ventral surface except for a medially situated shallow longitudinal trough in the midsection of its preserved length. The intercalar is difficult to recognize. The bilaterally swollen prootics are the largest bones of the lateral wall of the neurocranium; they are preserved dorsolateral to the anterior end of the preserved length of the parasphenoid (Fig. 2.2). The structure of the trigemino-facialis chamber is difficult to determine although a large facial foramen can be easily recognized. Each prootic contacts the sphenotic dorsolaterally and the pterotic posterodorsally, the exoccipital and basioccipital posteriorly, and the parasphenoid ventrally. The pterosphenoïd articulates dorsally with the frontal and sphenotic, and posteriorly with the prootic.

Etymology.—The specific epithet honors Stephen Groff and Marcus Jones, who co-discovered the specimen. The name '*stremphaencus*' is derived from a combination of the letters of 'Stephen' and 'Marcus' with some rearrangement to make the specific name mellifluous.

Remarks.—Despite being represented solely by a single partially complete neurocranium, the attribution of CMM-V-10,000 to the family Rachycentridae and, more particularly to the genus *Rachycentron*, is clearly justified by its overall morphology and configuration, as well as by the mutual relationships of the recognizable skull bones (see O'Toole, 2002). In particular, the remarkable posterior elongation of the epioccipitals and the medially oriented articular facets of the exoccipitals are considered as diagnostic of the genus *Rachycentron* (O'Toole, 2002; Friedman et al., 2013). Compared with the extant cobia *Rachycentron canadum*, the neurocranium of *R. stremphaencus* n. sp. exhibits a number of differences that unquestionably support its assignment to a new previously undescribed species. These differences mostly regard the relative proportions and development of the neurocranial portions, which are well exemplified in Figure 3. First, the external bony texture of the skulls is broadly different (Fig. 3, feature A). While that of the extant cobia consists of fine to bold anastomosing sharp ridges, the outer surface of the neurocranial bones of *Rachycentron stremphaencus* n. sp. is nearly smooth or scarcely ornamented.

In CMM-V-10,000, the lateral ridge is pronounced, resulting in a considerable gradient from the dorsal-medial exposure of the frontal to its lateralmost supraorbital margin (Fig. 3, feature B). In the extant cobia, the typical carangoid lateral ridge is weakly developed, and the frontal presents a gently curving outer surface.

The size, shape, and relative position of the sphenotic are conspicuously different between these two taxa (Fig. 3, feature C). In *Rachycentron stremphaencus* n. sp., the sphenotic is located in the posterior half of the neurocranium, and its lateralmost edge is adjacent to the anteriormost extent of the wedge-

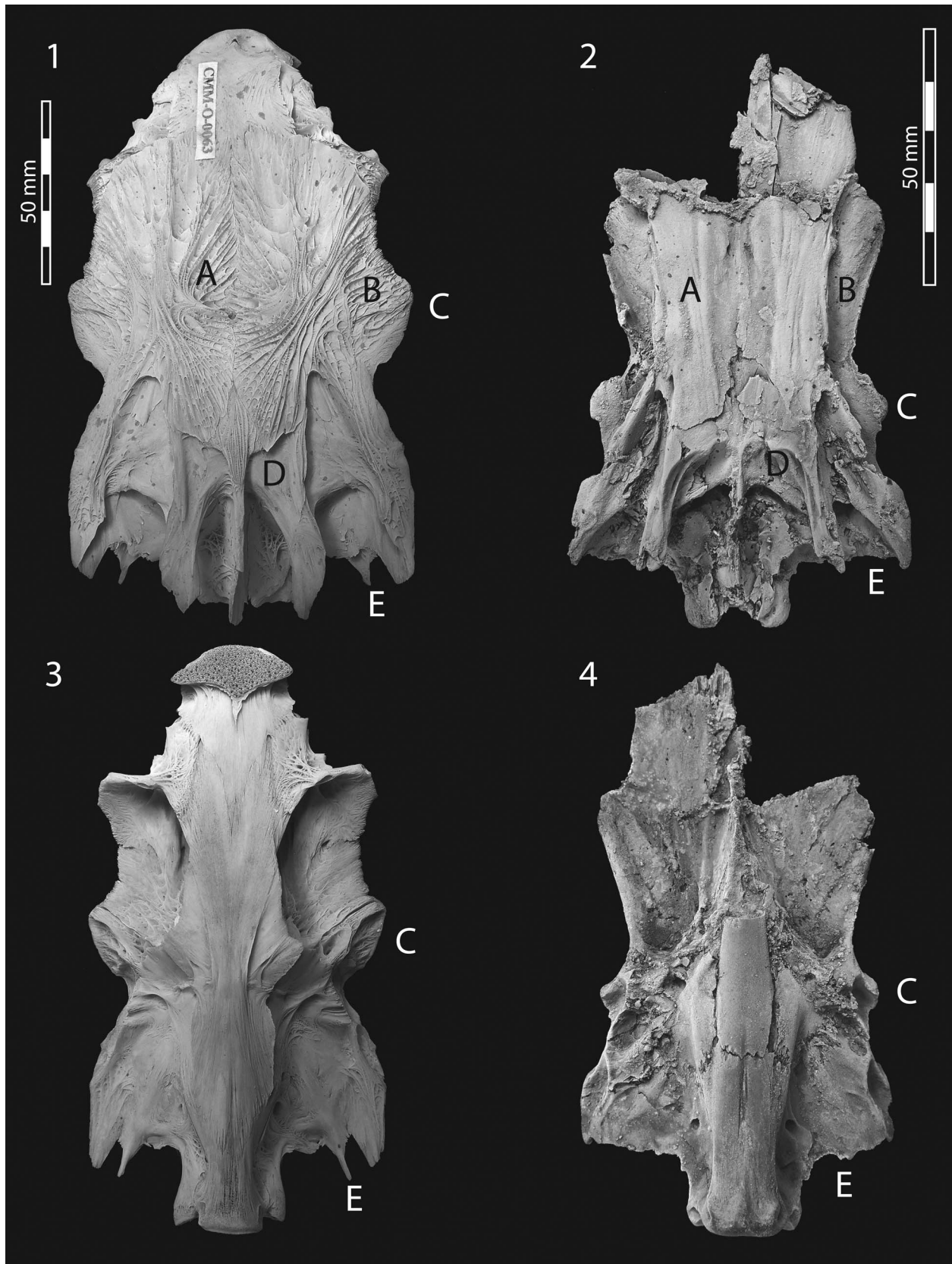


Figure 3. (1, 3) Neurocranium of *Rachycentron canadum*, CMM-O-0063: (1) dorsal view; (3) ventral view. (2, 4) Neurocranium of *Rachycentron strempaencus* n. sp., CMM-V-10,000: (2) dorsal view; (4) ventral view. The letters A–E enumerate the major differences between the neurocranium of *Rachycentron canadum* and that of *Rachycentron strempaencus* n. sp. Specimens dusted with sublimed ammonium chloride to improve contrast.

shaped trough in the dorsal surface of the skull formed by the lateral and medial ridges. In *Rachycentron canadum*, the sphenotic extends laterally at the same level of the pterotic and located further forward on the skull, approximately at the midpoint of its maximum length.

In *Rachycentron stremphaencus* n. sp., the two contralateral medial ridges form a proportionately much wider trough on either side of the supraoccipital than in *Rachycentron canadum*, in which the epioccipitals are much elongated posteriorly, reaching the rear edge of the neurocranium (Fig. 3, feature D). Finally, in *Rachycentron canadum*, there is a small but conspicuous posterolateral prong of the intercalar, which is apparently not present in *Rachycentron stremphaencus* n. sp. (Fig. 3, feature E).

Although the partial remains of one individual of *Rachycentron* have been described from the Calvert Formation (Carnevale and Godfrey, 2018), because there was no neurocranium preserved in that specimen, there is no way of knowing whether the Calvert cobia is conspecific with the one described herein from the St. Marys Formation. In comparing the fossil *Rachycentron* remains from the Calvert Formation with those of extant cobia, significant differences were not found.

As documented in the preceding, a set of associated bones (sclerotic ring, right premaxilla, left maxilla, left suspensorium and preopercle, right second pharyngobranchial, right posttemporal, vertebrae) from the Calvert Formation has been assigned by Carnevale and Godfrey (2018) to an indeterminate species of *Rachycentron*. Unfortunately, it is not possible to provide a reliable scrutiny of the potential affinities between *Rachycentron stremphaencus* n. sp. and the fossil cobia from the Calvert Formation due to the absence of comparable structures. As a consequence, any possible attribution of the fossil cobia from the Calvert Formation to the new species described herein must be excluded until additional comparable material would be available.

In summary, on the basis of the ecological requirements of its living congeners (e.g., Shaffer and Nakamura, 1989) and the paleoenvironmental reconstructions of the depositional context of the St. Marys Formation (Kidwell, 1988, 1997; Carnevale et al., 2011), it is reasonable to conclude that *Rachycentron stremphaencus* n. sp. was a predatory fish that inhabited the subtropical to warm temperate Atlantic coast of North America during the late Miocene, where it occurred in shallow marine and paralic waters, in a way similar to the other fossil and extant cobias.

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References

- Blake, S.F., 1940, *Paralbula*, a new fossil fish based on dental plates from the Eocene and Miocene of Maryland: Journal of the Washington Academy of Sciences, v. 30, p. 205–209.
- Carnevale, G., and Godfrey, S.J., 2018, Miocene bony fishes of the Calvert, Choptank, St. Marys, and Eastover Formations, Chesapeake Group, Maryland and Virginia: Smithsonian Contributions to Paleobiology, v. 100, p. 161–212.
- Carnevale, G., Godfrey, S.J., and Pietsch, T.W., 2011, Stargazer (Teleostei, Uranoscopidae) cranial remains from the Miocene Calvert Cliffs, Maryland, USA (St. Marys Formation, Chesapeake Group): Journal of Vertebrate Paleontology, v. 31, p. 1200–1209.
- Clark, W.B., Shattuck, G.B., and Dall, W.H., 1904, The Miocene Deposits of Maryland: Baltimore, Maryland Geological Survey, 543 p.
- Cope, E.D., 1867, An addition to the vertebrate fauna of the Miocene period, with a synopsis of the extinct Cetacea of the United States: Proceedings of the Academy of Natural Sciences of Philadelphia, v. 1867, p. 142.
- Cope, E.D., 1869, Descriptions of some extinct fishes previously unknown: Proceedings of the Boston Society of Natural History, v. 12, p. 310–317.
- Dante, J.H., 1953, Otoliths of a new fish from the Miocene of Maryland: Journal of Paleontology, v. 27, p. 877–885.
- Eastman, C.R., 1904, Pisces, in Clark, W.B., Shattuck, G.B., and Hall, W.D., eds., The Miocene of Maryland: Baltimore, Maryland Geological Survey, p. 71–93.
- Friedman, M., Johanson, Z., Harrington, R.C., Near, T.J., and Graham, M.R., 2013, An early fossil remora (Echeneoidea) reveals the evolutionary assembly of the adhesion disc: Proceedings of the Royal Society B: Biological Sciences, v. 280, 20131200.
- Gill, T.N., 1895, The nomenclature of *Rachycentron* or *Elacate*, a genus of acanthopterygian fishes: Proceedings of the United States National Museum, v. 18, p. 217–219.
- Godfrey, S.J. (ed.), 2018, The Geology and Vertebrate Paleontology of Calvert Cliffs. Smithsonian Contributions to Paleobiology, v. 100: Washington, DC, Smithsonian Institution Scholarly Press, 274 p.
- Jordan, D.S., 1923, A Classification of Fishes Including Families and Genera as Far as Known. Stanford University Publications, University Series, Biological Sciences, v. 3, no. 2, p. 77–243, Stanford, California, Stanford University Press.
- Kaup, J.J., 1826, Beitrage zur Amphibiologie und Ichthyologie: Isis (Oken), v. 19, p. 87–90.
- Kidwell, S.M., 1984, Outcrop features and origin of basin margin unconformities in the lower Chesapeake Group (Miocene), Atlantic Coastal Plain, in Schlee, J.S., ed., Interregional Unconformities and Hydrocarbon Accumulation: American Association of Petroleum Geologists Memoir 36, p. 37–58.
- Kidwell, S.M., 1988, Reciprocal sedimentation and noncorrelative hiatuses in marine-paralic siliciclastics: Miocene outcrop evidence: Geology, v. 16, p. 609–612.
- Kidwell, S.M., 1989, Stratigraphic condensation of marine transgressive records: origin of major shell deposits in the Miocene of Maryland: Journal of Geology, v. 97, p. 1–24.
- Kidwell, S.M., 1997, Anatomy of extremely thin marine sequences landward of a passive-margin hinge zone: Neogene Calvert Cliffs succession, Maryland, USA: Journal of Sedimentary Research, v. 67, p. 222–240.
- Kidwell, S.M., Powars, D.S., Edwards, L.E., and Vogt, P.R., 2015, Miocene stratigraphy and paleoenvironments of the Calvert Cliffs, Maryland: The Geological Society of America Field Guide, v. 40, p. 231–279.
- Kimmel, P.G., and Purdy, R., 1984, Fossil fish of the Calvert and Eastover Formations, in Ward, L.W., and Krafft, K., eds., Stratigraphy and Paleontology of the Outcropping Tertiary Beds in the Pamunkey River Region, Central Virginia Coastal Plain: Guidebook for Atlantic Coastal Plain Geological Association 1984 Field Trip: Newark, Delaware, Atlantic Coastal Plain Geological Association, p. 205–215.
- Leidy, J., 1873, Contributions to the extinct vertebrate fauna of the Western Territories. Report of the United States Geological Survey of the Territories, v. 1: Washington, DC, Government Printing Office, 358 p.
- Leriche, M., 1942, Contribution à l'étude des faunes ichthyologiques marines des terrains Tertiaires de la Plaine Côtière Atlantique et du centre des États Unis. Le synchronisme des formations Tertiaires des deux cotés de l'Atlantique: Mémoires de la Société Géologique de France (Nouvelle Série), v. 45, p. 1–111.

- Linnaeus, C., 1766, *Systema naturae sive regna tria naturae, secundum classes, ordines, genera, species, cum characteribus, differentiis, synonymis, locis*, v. I, no. 1, Editio duodecima, reformata: Holmiae, Laurentii Salvii, 532 p.
- Lynn, W.G., and Melland, A.M., 1939, A fossil catfish (*Felichthys stauroforus*) from the Maryland Miocene: *Journal of the Washington Academy of Sciences*, v. 30, p. 205–209.
- Müller, A., 1999, Ichthyofaunen aus dem atlantischen Tertiär der USA. *Leipziger Geowissenschaften*, v. 9/10: Leipzig, Institut für Geophysik und Geologie, 360 p.
- Nelson, J.S., 2006, *Fishes of the World* (fourth edition): Hoboken, New Jersey, Wiley, 601 p.
- Nolf, D., 2013, *The Diversity of Otoliths Past and Present*: Bruxelles, Royal Belgian Institute of Natural Sciences, 584 p.
- O'Toole, B., 2002, Phylogeny of the species of the superfamily Echeneoidea (Perciformes: Carangoidei: Echeneidae, Rachycentridae, and Coryphaenidae), with an interpretation of echeneid hitchhiking behavior: *Canadian Journal of Zoology*, v. 80, p. 596–623.
- Perez, V.J., Godfrey, S.J., Kent, B.W., Weems, R.E., and Nance, J.R., 2019, The transition between *Carcharocles chubutensis* and *Carcharocles megalodon* (Otodontidae, Chondrichthyes): lateral cusplet loss through time: *Journal of Vertebrate Paleontology*, v. 38, e1546732.
- Poag, C.W., 1979, Stratigraphy and depositional environments of Baltimore Canyon Trough: *American Association of Petroleum Geologists Bulletin*, v. 63, p. 1452–1466.
- Sajeevan, M.K., and Madhusoodana Kurup, B., 2014, Osteological features of cobia, *Rachycentron canadum* (Linnaeus 1766): *Journal of the Ocean Science Foundation*, v. 11, p. 40–49.
- Shaffer, R., and Nakamura, E.L., 1989, Synopsis of Biological Data on the Cobia *Rachycentron canadum* (Pisces: Rachycentridae): NOAA Technical Report NMFS 82, 21 p.
- Shattuck, G.B., 1904, Geological and paleontological relations, with a review of earlier investigations, in Clark, W.B., Shattuck, G.B., and Dall, W.H., eds., *The Miocene of Maryland: Baltimore, Maryland Geological Survey*, p. 71–93.
- Smith, B., 1909, Note on the Miocene drum fish—*Pogonias multidentatus* Cope: *American Journal of Science*, ser. 4, v. 28, p. 275–282.
- Vogt, P.R., and Eshelman, R.E., 1987, Maryland's Cliffs of Calvert: a fossiliferous record of mid-Miocene inner shelf and coastal environments, in Roy, D.C., ed., *Geological Society of America Centennial Field Guide, Northeastern Section*, p. 9–14.
- Ward, L.W., 1992, Molluscan biostratigraphy of the Miocene, Middle Atlantic Coastal Plain of North America: *Virginia Museum of Natural History Memoir*, v. 2, 159 p.
- Ward, L.W., and Andrews, G.W., 2008, Stratigraphy of the Calvert, Choptank, and St. Marys Formation (Miocene) in the Chesapeake Bay area, Maryland and Virginia: *Virginia Museum of Natural History Memoir*, v. 9, 60 p.
- Ward, L.W., and Strickland, G.L., 1985, Outline of Tertiary stratigraphy and depositional history of the US Atlantic Coastal Plain, in Poag, C.W., ed., *Geological Evolution of the United States Atlantic Margin*: New York, Van Nostrand Reinhold, p. 87–123.
- Weems, R.E., 1985, Miocene and Pliocene Molidae (*Ranzania*, *Mola*) from Maryland, Virginia, and North Carolina (Pisces: Tetraodontiformes): *Proceedings of the Biological Society of Washington*, v. 98, p. 422–438.
- Wiley, E.O., and Johnson, G.D., 2010, A teleost classification based on monophyletic groups, in Nelson, J.S., Schultze, H.-P., and Wilson, M.V.H., eds., *Origin and Phylogenetic Interrelationships of Teleosts*: München, Verlag Dr. Friedrich Pfeil, p. 123–182.

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