

Zirconosilicates in the kakortokites of the Ilímaussaq complex, South Greenland: Implications for fluid evolution and high-field-strength and rare-earth element mineralization in apgaitic systems

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[Received 7 April 2015; Accepted 25 June 2015; Associate Editor: Kathryn Goodenough]

ABSTRACT

The layered apgaitic nepheline syenites (kakortokites) of the Ilímaussaq complex, South Greenland, host voluminous accumulations of eudialyte-group minerals (EGM). These complex Na-Ca-zirconosilicates contain economically attractive levels of Zr, Nb and rare-earth elements (*REE*), but have commonly undergone extensive autometasomatic/hydrothermal alteration to a variety of secondary mineral assemblages. Three EGM alteration assemblages are recognized, characterized by the secondary zirconosilicates catapleite, zircon and gittinsite. Theoretical petrogenetic grid models are constructed to assess mineral stabilities in terms of component activities in the late-stage melts and fluids. Widespread alteration of EGM to catapleite records an overall increase in water activity, and reflects interaction of EGM with late-magmatic Na-, Cl- and F-rich aqueous fluids at the final stages of kakortokite crystallization. Localized alteration of EGM and catapleite to the rare Ca-Zr silicate gittinsite, previously unidentified at Ilímaussaq, requires an increase in CaO activity and suggests post-magmatic interaction with Ca-Sr bearing aqueous fluids. The pseudomorphic replacement of EGM in the kakortokites was not found to be associated with significant remobilization of the primary Zr, Nb and *REE* mineralization, regardless of the high concentrations of potential transporting ligands such as F and Cl. We infer that the immobile behaviour essentially reflects the neutral to basic character of the late-magmatic fluids, in which *REE*-F compounds are insoluble and remobilization of *REE* as Cl complexes is inhibited by precipitation of nacareniobsite-(Ce) and various Ca-*REE* silicates. A subsequent decrease in F activity would furthermore restrict the mobility of Zr as hydroxyl-fluoride complexes, and promote precipitation of the secondary zirconosilicates within the confines of the replaced EGM domains.

KEYWORDS: Ilímaussaq, late-magmatic alteration, eudialyte, catapleite, gittinsite, HFSE and *REE* mineralization, nepheline syenites, chemographic analysis, apgaitic systems.

Introduction

BASED on studies of the peralkaline Ilímaussaq complex, South Greenland, Ussing (1912) introduced the term apgaitite to describe nepheline

* E-mail: anoukborst@gmail.com
DOI: 10.1180/minmag.2016.080.046

syenites with molar (Na + K)/Al ratios greater than 1.2. Agpaitic rocks are additionally characterized by the presence of complex Na-Zr-Ti silicates, such as eudialyte-group minerals (EGM), aenigmatite and rinkite, instead of the common Zr-Ti phases, zircon and ilmenite (e.g. Sørensen, 1997; LeMaitre, 2003; Marks *et al.*, 2011). The magmas that produce these unusual assemblages typically form in rift-related settings and are characterized by extreme enrichments in alkalis, halogens (F, Cl, Br) and high-field-strength elements (HFSE), e.g. Zr, Ti, Nb, Ta and REE at economically attractive levels (e.g. Sørensen, 1992; Bailey *et al.*, 2001).

Agpaitic melts generally crystallize over extended temperature intervals (c. 1000 to 450°C) under reduced conditions (e.g. Markl *et al.*, 2001; 2010). Low oxygen fugacity is inferred to suppress segregation of aqueous fluids enriched in Na, F and Cl and other water-soluble elements from the evolving melt until near-solidus temperatures (e.g. Konnerup-Madsen *et al.*, 2001; Markl *et al.*, 2001). This enables the melt to evolve to highly sodic and volatile-rich compositions, enhancing the solubility of HFSE in the melt until their sequestration in the characteristic Na- and HFSE-rich agpaitic minerals (e.g. Sørensen *et al.*, 1997; Andersen *et al.*, 2010; Marks *et al.*, 2011). In the final stages of crystallization, the primary minerals are commonly subjected to intense auto-metasomatic alteration and hydrothermal overprinting (e.g. Markl *et al.*, 2001; Mitchell and Liferovich, 2006; Marks and Markl, 2015). In many peralkaline complexes, such processes led to significant remobilization of the primary HFSE and REE mineralization (e.g. Boily and Williams-Jones, 1994; Salvi and Williams-Jones, 1995; 1996; 2005; 2006; Salvi *et al.*, 2000; Sheard *et al.*, 2012; Gysi and Williams-Jones, 2013; Estrade *et al.*, 2015).

With a growing demand and potential supply-risk for HFSE and REE in the high-tech industry (Roskill, 2011), peralkaline and agpaitic complexes are considered as important alternatives to more commonly exploited resources for these elements, such as carbonatites and heavy-mineral sands (Chakhmouradian and Zaitsev, 2012). This has also sparked an increased demand for better characterization of the magmatic and subsequent hydrothermal petrogenesis of these deposits and their often complex mineralogy. To contribute to the progress in this field, we characterize EGM alteration assemblages from the agpaitic nepheline syenites of the Mesoproterozoic Ilímaussaq complex; the type locality for e.g. agpaites (Ussing, 1912), eudialyte *s.l.* (Stromeyer, 1819) and rinkite (Lorenzen, 1884).

Alternating layers of arfvedsonite-, EGM-, and feldspar-rich nepheline syenites, known as the kakortokites, in the lower part of the well-preserved and exposed Ilímaussaq magma chamber host large volumes of cumulus EGM (general formula $\text{Na}_{15}(\text{Ca}, \text{REE})_6(\text{Fe}, \text{Mn})_3\text{Zr}_3\text{Si}(\text{Si}_{25}\text{O}_{72})(\text{O}, \text{OH}, \text{H}_2\text{O})_3(\text{Cl}, \text{OH})_2$, Johnsen *et al.*, 2003). The latter account for most of the exploitable Zr, Nb and REE mineralization within the kakortokites, and can readily be extracted by magnetic separation (e.g. Bohse *et al.*, 1971; Sørensen, 1992; Zakharov *et al.*, 2011). Wide-spread alteration of EGM, however, led to the formation of complex pseudomorphic aggregates of secondary zirconosilicates, aluminosilicates and various Nb and REE phases. The alteration effectively fractionates the HFSE and REE among the secondary mineral phases and has unconstrained consequences for the bulk-rock concentrations and economic potential of the kakortokites.

Previous studies have shown two distinct alteration types after EGM in the kakortokites and associated pegmatites, either dominated by zircon or catapleiite (e.g. Ussing, 1912; Graser and Markl, 2008; Karup-Møller *et al.*, 2010; Karup-Møller and Rose-Hansen, 2013;). We report a third and new EGM alteration assemblage, containing the rare Ca-zirconosilicate gittinsite ($\text{CaZrSi}_2\text{O}_7$). Gittinsite was first described in association with vlasovite and EGM from the Kipawa complex in Canada (Ansell, 1980); to our knowledge, the only known occurrence of gittinsite in agpaitic nepheline syenites. More often gittinsite is found in low-temperature alteration assemblages with Zr phases such as elpidite, armstrongite and zircon in Si-saturated peralkaline granites, such as Khan Bogd and Khaldzan Buregte in Mongolia (e.g. Kynicky *et al.*, 2011; Kempe *et al.*, 1999), Strange Lake and Thor Lake in Canada (e.g. Birkett *et al.*, 1992; Salvi and Williams-Jones, 1995, 1996; Sheard *et al.*, 2012) and additionally in carbonatites of the Afrikanda complex in Russia (Chakhmouradian and Zaitsev, 2002). Crystallization of the Ca-mineral gittinsite seems unusual in the highly perisodic syenites of Ilímaussaq, which generally display Ca-depletion with magmatic evolution (Marks *et al.*, 2011). Based on the observed EGM alteration assemblages, we constructed theoretical chemographic grid models to assess the physico-chemical conditions which control EGM replacement reactions. Mass- and volume-relations are discussed to identify bulk loss and gain of components. Together, these efforts contribute to the understanding of late-magmatic and

hydrothermal fluid evolution in the kakortokites and their potential impact on HFSE and REE mineralization in agpaitic complexes.

Geological Setting

The Ilímaussaq complex (1160 ± 5 Ma; Waight *et al.*, 2002; Krumrei *et al.*, 2006) is part of the Mesoproterozoic Gardar Province, South Greenland (Upton, 2013). The complex was emplaced during at least four intrusive pulses successively crystallizing augite syenites, alkali granites, quartz syenites and the volumetrically dominant agpaitic nepheline syenites (see reviews in Larsen and Sørensen, 1987; Sørensen, 2001; Sørensen *et al.*, 2006; Marks and Markl, 2015). The agpaitic sequence (c. 1500 m thick) consists of: (1) an upper series of downward-crystallized roof cumulates (pulaskite, foyaite, sodalite foyaite and naujaite); (2) a lower series of rhythmically layered floor cumulates (kakortokites); and (3) a ‘sandwich horizon’ of lujavrites that represent the most evolved melt fractions (e.g. Ussing, 1912; Ferguson, 1964; Sørensen *et al.*, 2006). All agpaitic units are textural varieties of nepheline syenites with varying proportions of the liquidus phases alkali feldspar, nepheline, sodalite, aegirine, arfvedsonite and EGM. Common intercumulus and accessory minerals are aenigmatite, rinkite-group minerals, fluorite, apatite and analcime.

The present work focuses on the kakortokite-lujavrite sequence exposed at Kringlerne in the southern part of the complex (Fig. 1). The lower ~200 m of the kakortokite sequence exhibit strong magmatic layering (lower layered kakortokites or LLK), comprising at least 29 triple-layered units (Fig. 2). Each unit consists of a basal arfvedsonite-rich layer (black), a thin EGM-rich layer (red, sometimes absent) and an upper white layer rich in nepheline and feldspar (e.g. Bohse *et al.*, 1971; Larsen and Sørensen, 1987). The units are numbered from –11 to +16, relative to the well-developed marker horizon of unit 0. The layering steepens towards the margins of the intrusion and gradually disappears into a marginal pegmatite zone of the kakortokite (up to 100 m wide; Sørensen *et al.*, 2006), ending with a sharp intrusive contact against the earlier augite syenite (Fig. 1).

The LLK are conformably overlain by c. 50 m of weakly-layered and poorly exposed kakortokites (slightly layered kakortokites or SLK). On the northern side of the Lakseelv fault-zone more kakortokites (c. 60 m) are exposed also showing the

modal layering similar to the LLK (Fig. 1). These ‘transitional layered kakortokites’ (TLK) are labelled downwards from A to F and are characterized by an upwards decrease in grain size and modal amount of arfvedsonite relative to aegirine (Bohse and Andersen, 1981). The TLK gradually evolve into aegirine-dominated lujavrites, which are subdivided into a lower aegirine lujavrite I and an upper, finer-grained aegirine lujavrite II. These are overlain by arfvedsonite-dominated lujavrites. Both the TLK and lujavrites exhibit pronounced magmatic fabrics, indicative of low-viscosity melts and near-solidus to solid state deformation during emplacement (e.g. Sørensen *et al.*, 2006; Ratschbacher *et al.*, 2015). Detailed discussions on the evolution of the kakortokite-lujavrite series are provided by Pfaff *et al.* (2008), Marks and Markl (2015) and Ratschbacher *et al.* (2015).

Analytical methods

The EGM and their pseudomorph assemblages were analysed in samples from units –11 to +16 in the LLK, unit E and B from the TLK and EGM-rich layer II from aegirine lujavrite I. Zirconosilicate compositions were determined by wavelength-dispersive spectroscopy (WDS) using a Cameca SX100 electron probe micro-analyser (EPMA) at the University of Oslo. EGM were analysed with an acceleration voltage of 15 kV, a beam current of 15 nA and a spot size of 20–25 μm to minimize sample damage and Na migration. Other zirconosilicates were analysed using a 15 kV acceleration voltage, 10 nA beam current and 10 μm beam size, because of their smaller grain size. The elements Na, Zr, Si and Cl were measured first and X-ray counts were monitored in-run for signal stability to prevent migration and associated beam damage effects. Counting times were 10–60 s on peak, and half of that on each baseline, with shorter times for volatile elements. Calibration standards were wollastonite ($\text{CaK}\alpha$, $\text{SiK}\alpha$), albite ($\text{NaL}\alpha$), corundum ($\text{AlK}\alpha$), Fe metal ($\text{FeK}\alpha$), MnTi-oxide ($\text{MnK}\alpha$, $\text{TiK}\alpha$), orthoclase ($\text{KK}\alpha$), synthetic alforsite ($\text{ClK}\alpha$), Monastery Mine zircon ($\text{ZrL}\alpha$), Hf metal ($\text{HfM}\alpha$), Nb metal ($\text{NbL}\alpha$) and synthetic orthophosphates (Jarosewich and Boatner, 1991) for REE and Y ($\text{LaL}\alpha$, $\text{CeL}\alpha$, $\text{NdL}\beta$, $\text{YL}\alpha$). Full-range spectra and energy-dispersive spectral (EDS) analyses were acquired on a few grains prior to analyses, to check for the presence of elements not included in the analyses. Peak and background positions were chosen carefully to avoid interferences from other elements in the

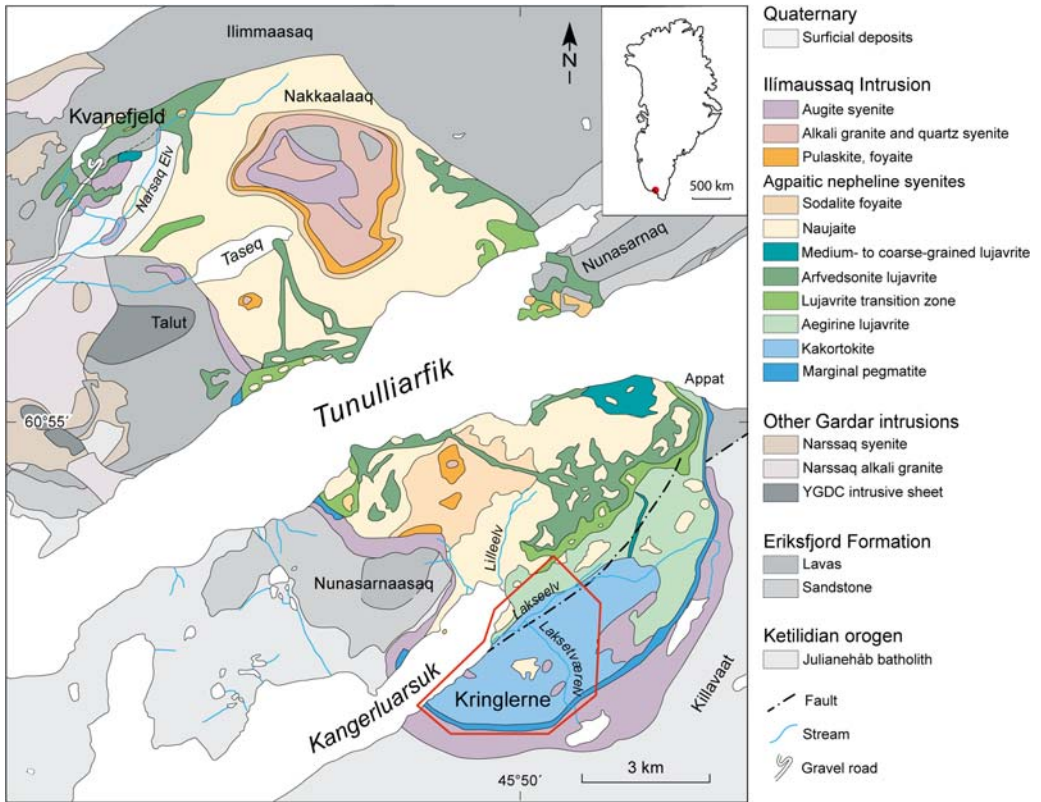


FIG. 1. Geological map of the Ilimaussaq intrusion and surrounding geology from Upton (2013), based on Ferguson (1964), Andersen *et al.*, (1988) and Sørensen (2001). Samples were obtained from the Kringlerne, Lakseelv and Laksetværelv areas in the southern part of the complex (red box).

spectral domains of interest. Discrete pulse-height analysis was applied to further minimize interferences from higher order spectral lines. Matrix corrections were applied according to the *PAP* procedure (Pouchou and Pichoir, 1984). EGM formulae were calculated based on $(Si + Zr + Ti + Nb + Al + Hf) = 29$ (Johnsen and Grice, 1999).

Backscattered electron images as well as analyses of aegirine were obtained using a JEOL JXA-8200 EPMA at the University of Copenhagen, using an acceleration voltage of 15 kV, beam current of 15 nA and beam size of 5 μ m. Counting times and X-ray lines were as described above. Calibration standards deviating from those above are corundum ($AlK\alpha$), hematite ($FeK\alpha$), baddelyite ($ZrL\alpha$) and rutile ($TiK\alpha$). Matrix corrections were carried out following the $\phi\rho z$ correction protocol. Secondary *REE* minerals were considered too small (<5 μ m) for accurate quantitative analyses using WDS, and were analysed semi-qualitatively using EDS.

A Siemens D5005 X-ray diffractometer equipped with a $CuK\alpha$ tube at the Natural History Museum of Oslo was used for mineral identification. A selection of Na and Ca zirconosilicates from the museum collection, i.e. catapleite, gaidonnayite, gittinsite, lemoynite, parakeldyshite, keldyshite, elpite, armstrongite and vlasovite, were checked by X-ray diffraction (XRD) (using only a small fraction of the grains) and embedded in epoxy to serve as external standards for EPMA analyses. Powders of catapleite- and gittinsite-bearing pseudomorphs after EGM and a crosscutting gittinsite-bearing vein were extracted for XRD analyses using a diamond-microdrill. The XRD confirmed the presence of catapleite and excluded the presence of related phases gaidonnayite ($Na_2Zr_3SiO_9 \cdot 2H_2O$) and hilairite ($Na_2ZrSi_3O_9 \cdot 3H_2O$). Moreover, the distinct platy or hexagonal crystal habit characteristic for catapleite (as opposed to mostly equidimensional habits for gaidonnayite) suggests that catapleite is the dominant phase in

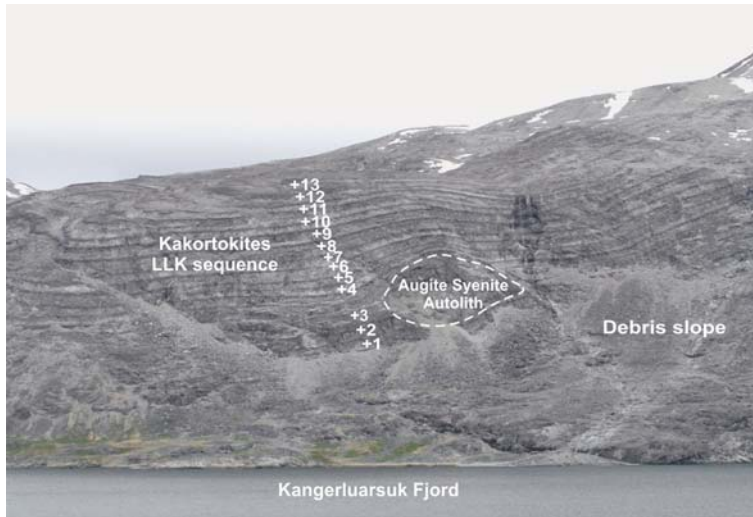


FIG. 2. Photograph of the lower layered kakortokite series (LLK). The cliff is ~300 m high and is exposed along the south-east coast of the Kangerluarsuk Fjord at Kringlerne. The visible units are numbered +1 to +13 (Bohse and Andersen, 1981). Unit +3 contains a large augite syenite autolith (dashed lines).

most EGM alteration assemblages. It should however be noted that the occurrence of gaiddonayite or hilairite cannot be excluded for samples not checked by XRD, given the difficulty of distinguishing these from catapleiite based on EPMA analyses alone.

Eudialyte-group minerals and their alteration paragenesis

In the kakortokites and lujavrites EGM occur as euhedral crystals which are part of the early crystallizing assemblage together with alkali feldspar, nepheline and arfvedsonite (e.g. Bohse *et al.*, 1971; Pfaff *et al.*, 2008; Lindhuber *et al.*, 2015; Ratschbacher *et al.*, 2015). Modal proportions of EGM vary from c. 5 to 20 vol.% in black and white kakortokites and lujavrites and 20 to 40 vol.% in red kakortokites. Average EGM grain sizes are 1 to 2 mm in the kakortokites (Fig. 3*a,b*) and ~0.3 mm in the TLK and aegirine lujavrite. Unaltered crystals typically exhibit sector-zoning, sub-micrometre oscillatory zoning and complex core-rim zoning (Fig. 3*a,b*). In the majority of samples EGM are partially or completely replaced by secondary assemblages. The degree of alteration ranges from initial replacement along crystal margins and along fractures to complete pseudomorphic replacement (Fig. 3*c-h*). Pseudomorphs

after EGM commonly cluster in zones of intense alteration, but also exist in direct contact with pristine EGM crystals (Fig. 3*c,d*). The degree of alteration is highly variable throughout the sequence and generally independent of lithology, although, EGM alteration is slightly more pervasive in the TLK and the aegirine lujavrites than in the LLK. Seventy-five percent of the investigated samples ($n \sim 80$) exhibit variable degrees of EGM alteration and in ~45% of the samples the majority of EGM are completely replaced. The degree and type of EGM alteration is generally independent of the surrounding mineral phases, although EGM are locally shielded from alteration when they are poikilitically enclosed by primary arfvedsonite.

Throughout the sequence, most EGM alteration assemblages contain catapleiite-group minerals as the major Zr phase; aegirine as the major Fe phase; fluorite and pectolite as major Ca phases; analcime, albite or K-feldspar as Na and K phases and nacareniobsite-(Ce) as the main Nb-REE phase. The REE are further distributed among micrometre-sized crystals of monazite, britholite-(Ce) and a related but poorly-defined group of Ca-REE silicates. The latter were studied by Karup-Møller and Rose-Hansen (2013), who described two dominant varieties and referred to them as Ca-rich A1 and Ca-poor A1, with the empirical formulae $\text{HCa}_3\text{REE}_6(\text{SiO}_4)_6(\text{F}\square)$ (where \square = vacancy) and $(\text{Fe,Mn,Ca})_{1.5}\text{REE}_6\text{Si}_6\text{FO}_{22}$, respectively. Both

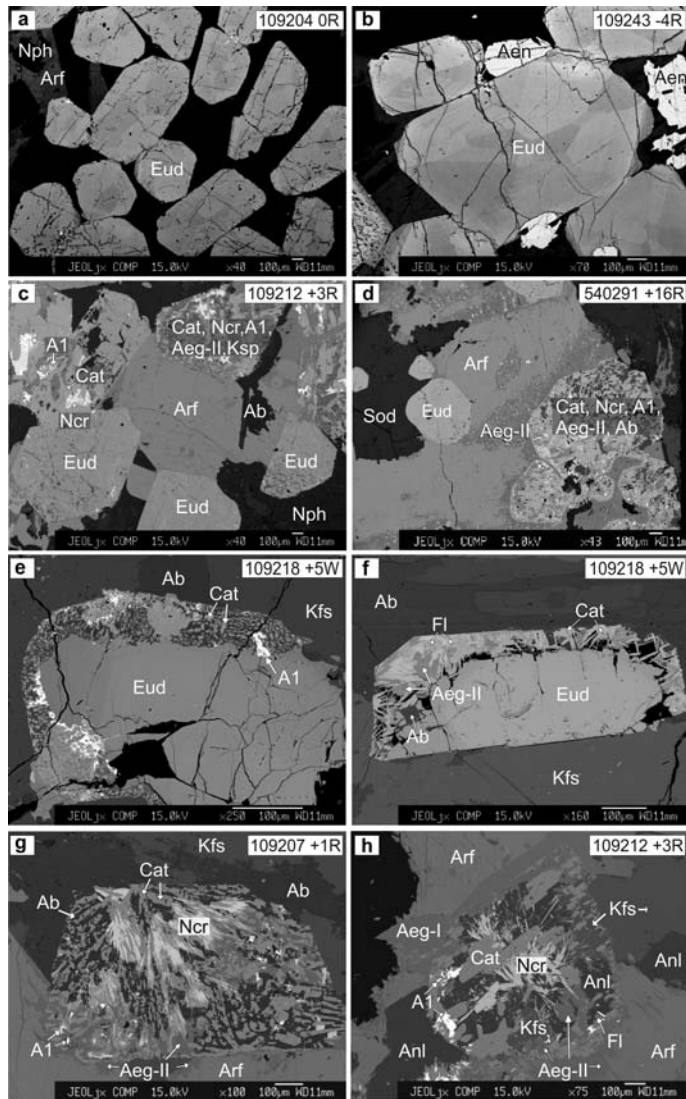


FIG. 3. Selection of backscatter electron images of EGM and their alteration products. Sample and layer number in top right corner. (a) Euhedral EGM crystals showing sector-zoning. (b) Sector- and oscillatory-zoned EGM with aenigmatite, arfvedsonite and nepheline. (c) Decomposed EGM adjacent to an unaltered EGM grain, surrounded by arfvedsonite and interstitial nepheline. (d) Catapleite dominated (type-I) pseudomorphs after EGM with replacement of adjacent arfvedsonite by secondary aegirine and Mn-Ti phases. (e) Symplectitic rim on EGM, consisting of finely-intergrown catapleite, albite, K-feldspar and the Ca-poor variety of A1. (f) Euhedral EGM with rim altered to catapleite, aegirine, albite and fluorite. Note the preservation of the original crystal faces. (g) Pseudomorph after EGM containing catapleite, nacareniobsite-(Ce), aegirine, pectolite, albite, A1 (Ca-rich variety) and minor Nb-oxides. (h) Type-I pseudomorph after EGM containing catapleite, nacareniobsite-(Ce), aegirine, analcime, K-feldspar and A1 (Ca-poor variety). Note that both K-feldspar and analcime have grown into the pseudomorph along their original contacts with the precursor eudialyte. Abbreviations: Arf = arfvedsonite, Nph = nepheline, Sod = sodalite, Eud = eudialyte, Aen = aenigmatite, Ab = albite, Kfs = K-feldspar, Cat = catapleite, Fl = fluorite, Aeg-I = magmatic aegirine, Aeg-II = secondary aegirine, Ncr = nacareniobsite-(Ce), A1 = Ca-rich or Ca-poor REE silicate.

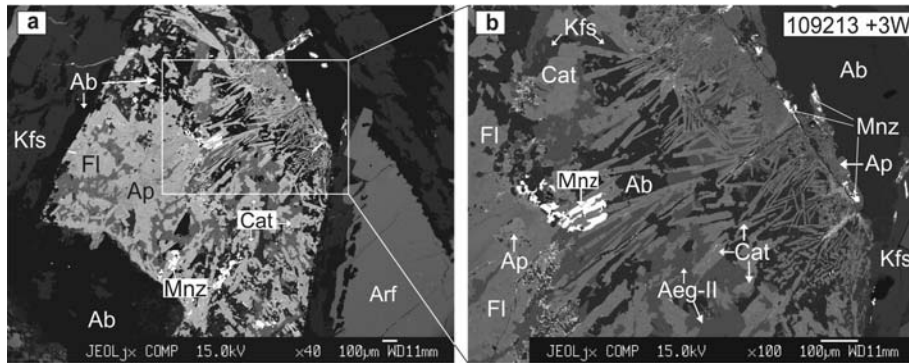


FIG 4. (a) Type-I pseudomorph after EGM in a white kakortokite, containing monazite (Mnz, bright mineral) as the major REE phase, together with abundant fluorapatite (Ap), fluorite and interstitial albite, K-feldspar and aegirine. (b) Thin laths of catapleite nucleating from the outer margins of the precursor eudialyte crystal.

varieties were observed in the current sample-set. Some pseudomorph assemblages contain dispersed grains (<5 µm) of additional Nb phases such as pyrochlore, fersmite and fergusonite-(Y).

Previous studies (Ussing, 1912; Karup-Møller *et al.*, 2010) also reported EGM alteration assemblages containing zircon, in addition to catapleite. These occurrences are limited to extensively altered kakortokites, lujavrites and related pegmatites, i.e. spatially restricted lithologies that were not investigated in the current study. Instead, we describe a third EGM alteration assemblage from the kakortokites, dominated by the rare Ca zirconosilicate gittinsite. Only the two observed parageneses characterized by catapleite and gittinsite (type-I and II, respectively) are described in further detail below.

Catapleite-bearing paragenesis

Catapleite-group minerals ((Na,Ca, \square)₂ZrSi₃O₉·2H₂O) are the most common alteration products after EGM throughout the complex. At small degrees of alteration, catapleite replaced EGM along rims and cracks, forming anhedral to subhedral grains, sometimes intimately intergrown with secondary feldspar and/or aegirine (Fig. 3c,d). At larger degrees of replacement, catapleite occurs as subhedral to euhedral elongate crystals, up to 0.5 mm long. In some alteration aggregates catapleite crystals nucleate inwards from the rim of the precursor EGM grain (Fig. 3f, Fig. 4b) and in some coarser grained pseudomorphs the well-developed catapleite laths are stacked and fully aligned. Aegirine, pectolite, fluorite, various Nb and REE

phases and occasionally apatite generally form subhedral aggregates between the catapleite crystals (Fig. 3). The remaining interstitial space is occupied by aluminosilicates, i.e. albite, K-feldspar or analcime, which can occur together in a single pseudomorph (Fig. 3e-h, Fig. 4).

Gittinsite-bearing paragenesis

Gittinsite is the dominant zirconosilicate in EGM pseudomorphs in a sample from layer +3 black (109211). It occurs as subhedral crystals of <15 µm in size, which are intergrown with Ca-rich catapleite and annite (Fe-biotite, Fig. 5). Textural relations suggest that gittinsite partially or completely replaced catapleite from a previous EGM alteration paragenesis (Fig. 5b,c,d). Other common phases in the gittinsite-bearing pseudomorphs are analcime, aegirine, fluorite, pectolite, allanite-(Ce), fergusonite-(Y) and Ca-REE silicates A1 and A2. A few grains of relict EGM are partially replaced by gittinsite, analcime and annite, with no evidence for intermediate catapleite formation (Fig. 5f,g,h). Primary aluminosilicates (e.g. feldspar, nepheline and sodalite) are replaced by analcime throughout the sample. Furthermore, a 1–2 mm-wide hydrothermal vein of pectolite, annite and analcime crosscuts the primary mineral assemblage. Aggregates of gittinsite, intergrown with catapleite, annite and an unidentified Ca-Y-HREE phase were found inside the hydrothermal vein, in close proximity of a fragmented precursor EGM crystal. In a separate sample (layer 0 red), an isolated crystal of gittinsite was identified within a catapleite-dominated pseudomorph after EGM,

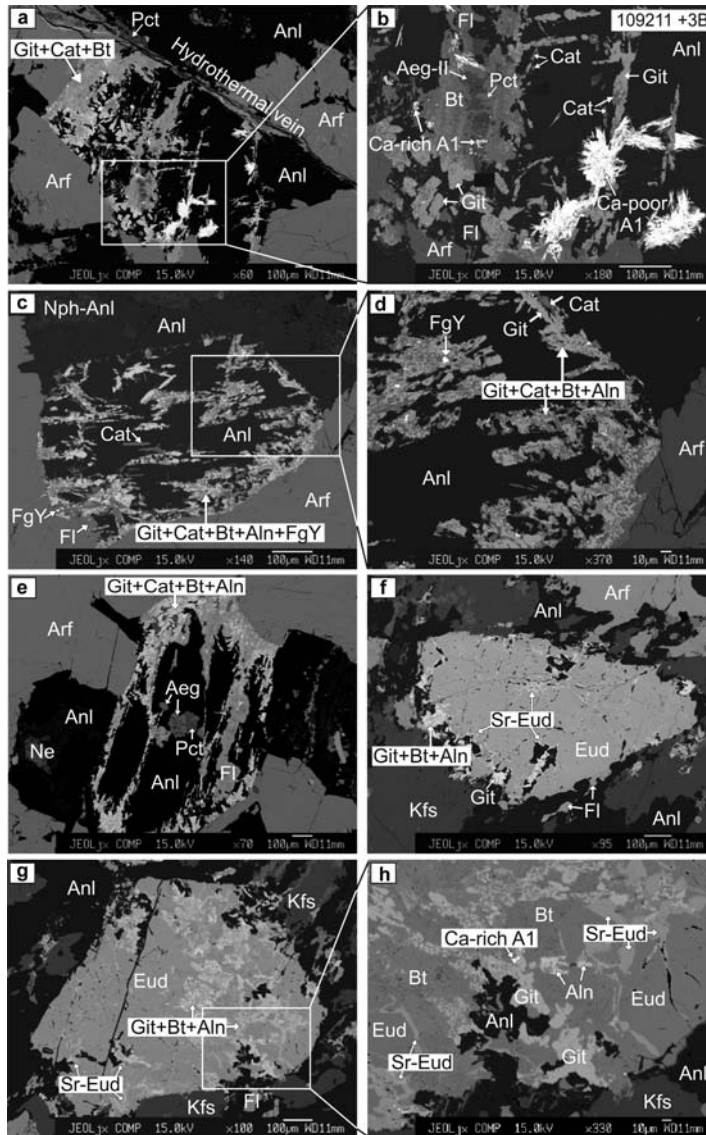


FIG. 5. Gittinsite-bearing (type-II) alteration assemblages after EGM in black kakortokite. Associated alteration minerals include catapleiite (Cat), Sr-rich eudialyte (Sr-Eud), analcime (Anl), fluorite (Fl), annite (Bt), aegirine (Aeg-II), pectolite (Pct), Ca-rich and Ca-poor A1, allanite-(Ce) (Aln) and fergusonite-(Y) (FgY). (a, b) Type-II pseudomorph after EGM in contact with a hydrothermal vein. Gittinsite is intergrown with annite and catapleiite. Radiating aggregates of Ca-poor A1 crystals, nucleating from gittinsite-catapleiite intergrowths. (c, d) Type-II pseudomorph after EGM showing an overprinted assemblage of gittinsite, catapleiite and annite on former catapleiite laths from type-I alteration paragenesis. Dominant Nb and REE phases are fergusonite-(Y) and allanite-(Ce). (e) Gittinsite-bearing pseudomorph after EGM containing intergrowths of gittinsite, catapleiite, allanite-(Ce) and annite, with ‘pillars’ of fluorite that precipitated around gittinsite-catapleiite intergrowths. (f, g) Remnant EGM crystals with secondary rims of Sr-rich eudialyte and partial decomposition to gittinsite, annite, analcime, allanite-(Ce) and Ca-rich A1. Catapleiite is absent. (h) Detail of partially altered EGM to gittinsite, annite (irregular surface, same shade as EGM), allanite-(Ce) and analcime. The rim of Sr-rich EGM (5–15 μm) separates remnants of primary EGM from the gittinsite-annite intergrowths.

potentially implying a more widespread occurrence of gittinsite in the kakortokites.

Mineral compositions

Eudialyte-group minerals

Eudialyte-group minerals belong to large group of Na-Ca-Zr silicates with complex crystallographic structures. The IMA-accepted formula (Johnsen *et al.*, 2003) for EGM is $N_{15-16}M(1)_6M(2)_3Z_3M(3)M(4)Si_{24}O_{66-73}(OH)_{0-9}X_2$, where $N = Na^+, K^+, Sr^{2+}, Ca^{2+}, REE^{3+}, \square, Ba^{2+}, Mn^{2+}, H_3O^+$; $M(1) = Ca^{2+}, REE^{3+}, Mn^{2+}, Fe^{2+}, Na^+, Sr^{2+}$; $M(2) = Fe^{2+}, Mn^{2+}, Na^+, \square, H_3O^+, Zr^{4+}, Ta^{5+}, Ti^{4+}, K^+, Ba^{2+}$; $M(3, 4) = Si^{4+}, Al^{3+}, Nb^{5+}, Ti^{4+}, W^{6+}, Na^+, Z = Zr^{4+}, Hf^{4+}, Ti^{4+}, Nb^{5+}, OH = H_2O, OH^-, O^{2-}, CO_3^{2-}, SO_4^{2-}, SiO_4^{4-}$; $X = Cl^-, F^-, OH^-$.

Throughout the kakortokite sequence EGM compositions are relatively constant, closely approaching compositions of eudialyte *sensu stricto*: $Na_{15}Ca_6Fe_3Zr_3Si(Si_{25}O_{72})(O,OH,H_2O)_3(Cl,OH)_2$, with minor deviations towards Mn-, REE- and Nb-bearing end-members (kentbrooksit component; Johnsen *et al.*, 2003; Pfaff *et al.*, 2008). Both within individual kakortokite units (Lindhuber *et al.*, 2015), and across the overall stratigraphy (Pfaff *et al.*, 2008), EGM exhibit decreasing molar Fe/Mn ratios, interpreted to indicate continued magmatic fractionation (e.g. Schilling *et al.*, 2011; Marks and Markl, 2015). From the base of the kakortokites towards the lujavrites, average EGM furthermore show decreasing Ca (6–5 a.p.f.u.), Cl (1.5–1 a.p.f.u.), Nb (0.23–0.17 a.p.f.u) and Hf contents (0.4–0.3 a.p.f.u.) and increasing REE + Y (0.45–0.5 a.p.f.u., Table 1).

Relics of primary EGM in partially replaced grains show little to no compositional modifications relative to unaltered EGM, except for occasional and minor loss of Na and Cl. This includes preservation of the primary sector zoning in the relict EGM domains. Metasomatic changes leading to distinct secondary EGM compositions have only been observed in the gittinsite-bearing assemblages from layer +3 black, where remnants of primary EGM are mantled by a thin rim (1–5 μ m) of a Sr-rich and Na-poor hydrothermal EGM variety (taseqite component; Johnsen *et al.*, 2003, Fig. 5f,g,h).

Catapleite-group minerals

Catapleite compositions range within the calcio-catapleite-catapleite solid solution series

(Table 2), with Ca contents ranging from 0.2 to 6.3 wt.% (0.02–0.47 a.p.f.u.) and Na contents from 7.8 to 14.3 wt.% (1–1.8 a.p.f.u.). The most Ca-rich catapleite co-exists with gittinsite in sample 109211. In some pseudomorphs, larger catapleite crystals (>30 μ m) are zoned with rims showing minor Ca enrichment relative to the core (+0.05 a.p.f.u.). Contents of Y_2O_3 and Hf_2O range between 0.1–0.9 wt.% and 0.4–0.7 wt.%, respectively, and are weakly inversely correlated with Y increasing and Hf decreasing across the sequence.

Gittinsite

Gittinsite exhibits no significant zoning or compositional variations. Average gittinsite compositions are in accordance with the ideal formula $CaZrSi_2O_7$ (Table 3). Calcium contents are near-constant (18–19 wt.% CaO) and the Na content is generally negligible (<0.1 wt.% Na_2O). Contents of Nb_2O_5 , Y_2O_3 and Hf_2O range between 0–0.4 wt.%, 0.1–0.3 wt.% and ~0.8 wt.%, respectively.

Clinopyroxene

Clinopyroxene in the kakortokites and lujavrites are compared to pyroxene from earlier Ilimaussaq rock types (e.g. augite syenites, Larsen, 1976; Markl *et al.*, 2001, Marks and Markl, 2015). At least two generations of pyroxenes are distinguished in most samples: (1) primary magmatic aegirine; and (2) secondary aegirine replacing primary Fe phases such as arfvedsonite, EGM and aenigmatite. In all EGM pseudomorphs, secondary aegirine forms the main Fe-bearing phase. Additionally, aegirine typically replaces primary arfvedsonite along contacts with altered EGM (Fig. 3e,f). Magmatic and secondary aegirine can be distinguished petrographically as they exhibit vivid green and brownish green pleochroism, respectively. Primary aegirine exhibits solid-solution trends in the aegirine–augite series with variable Ca (0.03–0.16 a.p.f.u.) and low Al (0.05 a.p.f.u., Table 4). In contrast, secondary aegirine is exclusively Na-rich (1 a.p.f.u.), with low Ca (0.02 a.p.f.u.) and variable Al (0.01–0.18 a.p.f.u.). Following the Morimoto classification (Morimoto, 1998) for Na-pyroxenes (i.e. Quad-jadeite-aegirine, where Quad represents all Ca-Fe-Mg pyroxenes) magmatic aegirine has relatively high Quad-components (up to 12 mol.%), whereas secondary aegirine has higher jadeite contents (up to 18 mol.%, averages in Table 4). Increasing jadeite-components are inferred to

TABLE 1. Average core compositions of eudialyte-group minerals from selected samples across the sequence.

Sample no.	520701		109242		109204		109211		109218		109236		540247	
Layer	-11 white		-4 black		0 red		3 black		5 white		13 white		E red	
<i>n</i>	Average	2σ	Average	2σ	Average	2σ	Average	2σ	Average	2σ	Average	2σ	Average	2σ
Na ₂ O	13.95	0.32	13.69	0.23	13.57	0.31	13.10	0.11	14.40	0.17	12.25	3.16	13.01	0.60
Al ₂ O ₃	0.27	0.08	0.22	0.06	0.23	0.04	0.25	0.02	0.25	0.04	0.31	0.20	0.31	0.21
SiO ₂	48.15	0.75	48.06	0.64	48.19	0.89	48.85	0.26	48.34	0.72	49.28	1.05	49.06	0.67
Cl	1.60	0.05	1.57	0.05	1.52	0.05	1.50	0.03	1.47	0.05	1.41	0.08	1.27	0.07
K ₂ O	0.28	0.05	0.27	0.03	0.23	0.03	0.27	0.03	0.32	0.03	0.28	0.07	0.24	0.04
CaO	10.59	0.30	10.49	0.39	10.17	0.38	10.55	0.39	10.31	0.28	9.90	0.37	9.43	0.44
FeO	6.20	0.26	7.09	0.17	6.31	0.22	6.32	0.23	6.12	0.16	5.54	0.19	6.37	0.39
MnO	0.77	0.07	0.57	0.05	0.75	0.05	0.65	0.13	0.73	0.11	0.80	0.05	0.92	0.08
Y ₂ O ₃	0.41	0.04	0.43	0.08	0.44	0.07	0.43	0.05	0.44	0.06	0.46	0.04	0.53	0.06
ZrO ₂	11.37	0.17	11.46	0.23	11.23	0.19	10.55	0.31	10.91	0.27	10.91	0.28	10.83	0.38
Nb ₂ O ₅	0.97	0.30	0.90	0.29	0.97	0.28	0.90	0.35	0.82	0.31	0.70	0.26	0.69	0.34
La ₂ O ₃	0.46	0.08	0.48	0.09	0.50	0.06	0.51	0.04	0.47	0.12	0.49	0.06	0.46	0.11
Ce ₂ O ₃	0.90	0.16	0.89	0.14	0.95	0.12	1.06	0.11	0.93	0.24	0.95	0.11	0.97	0.16
Nd ₂ O ₃	0.34	0.14	0.34	0.08	0.37	0.16	0.36	0.15	0.39	0.15	0.38	0.08	0.42	0.11
HfO ₂	0.25	0.04	0.24	0.05	0.25	0.04	0.24	0.05	0.25	0.05	0.24	0.06	0.21	0.07
TiO ₂	0.07	0.02	0.06	0.02	0.06	0.04	0.05	0.02	0.05	0.02	0.04	0.03	0.05	0.02
O = Cl	0.36		0.36		0.34		0.34		0.33		0.32		0.29	
Total	96.23		96.40		95.38		95.25		95.87		93.63		94.51	
Formulae based on (Si, Zr, Ti, Nb, Al, Hf) = 29														
Na	14.38		14.14		14.00		13.44		14.88		12.44		13.27	
Al	0.17		0.14		0.15		0.16		0.16		0.19		0.20	
Si	25.59		25.61		25.65		25.85		25.75		25.80		25.81	
Cl	1.44		1.42		1.37		1.34		1.33		1.25		1.14	
K	0.19		0.19		0.16		0.18		0.21		0.19		0.16	
Ca	6.03		5.99		5.80		5.98		5.89		5.56		5.32	
Fe	2.76		3.16		2.81		2.80		2.73		2.42		2.80	
Mn	0.35		0.26		0.34		0.29		0.33		0.35		0.41	
Y	0.12		0.12		0.12		0.12		0.13		0.13		0.15	
Zr	2.95		2.98		2.91		2.72		2.84		2.79		2.78	

Nb	0.23	0.22	0.23	0.22	0.17	0.17
La	0.09	0.09	0.10	0.10	0.09	0.09
Ce	0.17	0.17	0.19	0.18	0.18	0.19
Nd	0.06	0.06	0.07	0.07	0.07	0.08
Hf	0.04	0.04	0.04	0.04	0.04	0.03
Ti	0.03	0.02	0.02	0.02	0.02	0.02
O	75.84	76.00	75.47	75.89	73.98	74.56
Sum REE + Y	0.45	0.45	0.48	0.47	0.48	0.50

reflect decreasing silica activity with progressive crystallization and differentiation (Markl *et al.*, 2001), and as such fit a late-stage/hydrothermal origin. Zirconium contents range from 0 to 1.6 wt.% in magmatic aegirine, vs. 0 to 2.7 wt.% in secondary aegirine.

Nacareniobsite-(Ce) and rinkite-group minerals

Nacareniobsite-(Ce) ($\text{Na}_3\text{Ca}_3\text{REENb}(\text{Si}_2\text{O}_7)_2\text{OF}_3$) is common in most EGM pseudomorph assemblages and contains ~14 wt.% light rare-earth oxides (*LREO*; Ce + La + Nd) and ~15 wt.% Nb_2O_5 . In the absence of fluorite or pectolite, nacareniobsite-(Ce) incorporates most of the Ca released from EGM. Nacareniobsite-(Ce) occurs either as elongate prismatic crystals (<30 μm) or interstitially to catapleiite laths (Fig. 3c,g,h). Sector zoning was observed in some crystals. Nacareniobsite-(Ce) is related closely to the rinkite-mosandrite group minerals (ideally $\text{Na}_2(\text{Ca},\text{REE})_4\text{Ti}(\text{Si}_2\text{O}_7)_2\text{OF}_3$), and represents the Nb-rich analogue following the coupled substitution $\text{Nb}^{5+} + \text{Na}^+ \leftrightarrow \text{Ti}^{4+} + \text{Ca}^{2+}$ (e.g. Sokolova and Hawthorne, 2008; Cámara *et al.*, 2011; Rønso *et al.*, 2013). Both rinkite and nacareniobsite-(Ce) were first described from Ilímaussaq (Lorenzen, 1884; Petersen *et al.*, 1989). Rinkite is an accessory intercumulus mineral and is relatively abundant in white kakortokites. In contrast to magmatic rinkite, which shows significant solid solution between Ti and Nb and has low Na, Ca and F contents (Rønso *et al.*, 2013), secondary nacareniobsite-(Ce) shows near ideal end-member compositions. In common with the EGM, magmatic rinkite has been subjected to significant late-stage alteration (Rønso *et al.*, 2013).

Other REE phases

Besides nacareniobsite-(Ce) the most common *REE* minerals in the alteration assemblages are Ca-rich and Ca-poor varieties of the *REE* silicates A1 (Karup-Møller and Rose-Hansen 2013; Fig. 3c, e,g,h). These generally exhibit acicular or platy habits and are typically smaller than 5 μm . Semi-quantitative EDS analyses suggest that A1 minerals contain up to 65 wt.% *LREO* and thus represent a significant host for the released *REE*. Our analyses also indicate the presence of some varieties that are richer in P and poorer in *REE*, approaching compositions of britholite-(Ce) (following a $\text{REE}^{3+} + \text{Si}^{4+} \leftrightarrow \text{Ca}^{2+} + \text{P}^{5+}$ coupled

TABLE 2. Average compositions of catapleite-group minerals from selected samples across the sequence.

Sample no.	520704	109255	109208	109207	109211	109212	109213	109229	540247	540279
Layer	-10 white	-6 white	1 white	1 red	3 black	3 red	3 white	10 red	E red	Aeg Lj I
<i>n</i>	Average 2σ 6	Average 2σ 12	Average 2σ 6	Average 2σ 11	Average 2σ 2	Average 2σ 19	Average 2σ 9	Average 2σ 13	Average 2σ 12	Average 2σ 5
Na ₂ O	13.65 2.04	13.76 0.85	8.04 0.78	8.88 1.70	8.25 0.19	8.21 1.16	10.61 2.19	9.35 2.84	13.30 1.01	9.24 1.05
SiO ₂	43.52 0.43	43.44 0.51	43.05 0.99	43.20 1.32	42.39 0.87	43.27 0.35	43.60 0.90	43.25 1.18	43.40 0.80	43.82 0.58
CaO	0.24 0.26	0.23 0.34	5.80 1.08	4.50 1.82	6.31 1.99	5.80 1.38	1.28 1.37	3.86 3.72	0.38 0.43	3.12 1.38
FeO	0.16 0.19	0.05 0.14	0.08 0.12	0.19 0.10	0.25 0.40	0.05 0.08	0.14 0.35	0.27 1.32	0.29 0.31	0.09 0.06
MnO	0.01 0.04	0.00 0.05	0.01 0.04	0.03 0.12	0.01 0.06	0.01 0.03	0.01 0.04	0.11 0.30	0.06 0.15	0.04 0.04
Y ₂ O ₃	0.26 0.62	0.15 0.58	0.32 0.40	0.18 0.55	0.37 0.11	0.14 0.20	0.29 0.53	0.57 0.61	0.42 0.48	0.66 0.71
ZrO ₂	29.81 1.93	30.78 0.96	29.46 1.00	30.30 2.04	29.78 0.88	30.12 0.69	30.31 1.85	29.10 2.40	30.12 1.66	30.20 1.50
Nb ₂ O ₅	0.50 1.96	b.d.*	0.10 0.38	b.d.	0.18 0.26	b.d.	0.35 1.05	0.96 0.96	0.19 0.29	0.20 0.38
HfO ₂	0.62 0.13	0.68 0.15	0.60 0.15	0.69 0.14	0.61 0.22	0.67 0.08	0.67 0.16	0.59 0.16	0.55 0.08	0.44 0.14
Total**	88.80	89.09	87.47	87.96	88.15	88.28	87.26	88.07	88.71	87.79
Fomulae based on 9 oxygens										
Na	1.82	1.83	1.08	1.19	1.11	1.10	1.43	1.25	1.78	1.23
Si	2.99	2.99	2.99	2.99	2.94	2.98	3.03	2.98	2.99	3.02
Ca	0.02	0.02	0.43	0.33	0.47	0.43	0.10	0.29	0.03	0.23
Fe	0.01	0.00	0.00	0.01	0.01	0.00	0.01	0.02	0.02	0.00
Mn	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01	0.00	0.00
Y	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.02	0.02	0.02
Zr	1.00	1.03	1.00	1.02	1.01	1.01	1.03	0.98	1.01	1.01
Nb	0.02	0.00	0.00	0.00	0.01	0.00	0.01	0.03	0.01	0.01
Hf	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01

*b.d. = below detection limit.
 **Low totals due to high water contents.

TABLE 3. Average composition of gittinsite.

Sample no.	109211 (3 black)			
	Average*	2σ	Average**	2σ
<i>n</i>	14		7	
Na ₂ O	0.10	0.29	b.d	
SiO ₂	39.02	0.45	41.19	0.25
CaO	18.29	0.55	19.10	0.73
ZrO ₂	38.49	1.54	39.80	1.18
Nb ₂ O ₅	0.16	0.72	0.39	0.62
HfO ₂	0.78	0.12	n.a.	
Total	96.84		100.47	
Formulae based on 7 oxygens				
Na	0.01		–	
Si	2.00		2.02	
Ca	1.01		1.01	
Zr	0.96		0.95	
Nb	0.00		0.01	
Hf	0.01		–	

*Cameca analyses at the University of Oslo.

**JEOL analyses at the University of Copenhagen.

substitution scheme; cf. Karup-Møller *et al.*, 2010; A1 with apatite structure). An unknown Ba-rich REE silicate (c. 6–10 wt.% BaO, 0–1% CaO, 2% Nb₂O₅, 55% LREO, 20% SiO₂ and 2% F) was observed in six samples. These minerals exhibit similar molar proportions to the A1 REE silicates, with inversely correlated Ca and Ba contents. The Ba-REE silicates have similar habits (radiating fibrous or platy) and back-scatter index as the A1 minerals, making them indistinguishable without compositional data. Several of the above-mentioned REE phases can occur together in a single pseudomorph. Catapleite-type pseudomorphs containing secondary apatite and fluorite contain monazite as the main REE phase (Fig. 4). Monazite was only observed in white kakortokites. The gittinsite-bearing assemblages contain allanite-(Ce), as well as Ca-poor A1 minerals, fergusonite-(Y) and an unidentified Ca-HREE-Y silicate.

Discussion

EGM alteration; characterization of assemblages and mechanisms

Petrographic observations of the kakortokites at Ilímaussaq demonstrate substantial subsolidus alteration of primary minerals including

decomposition of the primary HFSE phases EGM and rinkite, zeolitization of primary feldspars and feldspathoids and alteration of Fe phases arfvedsonite and aenigmatite. Three types of EGM decomposition assemblages are distinguished based on the presence of the secondary zirconosilicates catapleite, zircon or gittinsite (Fig. 6). The gittinsite-bearing paragenesis (type-II, Fig. 6) represents a previously unidentified EGM alteration assemblage and was found to be pervasive in one sample and rarely dispersed in another, in both cases overprinting a previous catapleite-alteration paragenesis.

Catapleite-bearing assemblages (type-I, Fig. 6) are dominant throughout the kakortokites and most agpaite units of Ilímaussaq. Zircon-bearing assemblages (type-III, Fig. 6) were not observed in the current study, which confirms their restricted occurrence to the more extensively altered kakortokites, lujavrites and pegmatites (Karup-Møller *et al.*, 2010). Alteration of EGM to catapleite is commonly explained by interaction with hydrous interstitial melts, or late-magmatic aqueous fluids at the final stages of crystallization (Ussing, 1912; Karup-Møller *et al.*, 2010; Karup-Møller and Rose-Hansen, 2013). The existence of such late-stage fluids is confirmed by the presence of aqueous, commonly highly saline, fluid inclusions in secondary phases in the kakortokites (Konnerup-Madsen, 2001). In contrast, zircon-bearing assemblages were inferred to result from fluids expelled by late lujavrite melts or from external fluids with reduced alkalinity (Karup-Møller *et al.*, 2010). Both catapleite- and zircon-bearing EGM alteration assemblages have also been described from several other agpaite complexes, i.e. North Qôroq (Greenland), Pilanesberg (South Africa), Sushina Hill (India) and Tamazeght (Morocco) (e.g. Coulson, 1997; Khadem Allah *et al.*, 1998; Mitchell and Liferovich, 2006; Mitchell and Chakrabarty, 2012; Chakrabarty *et al.*, 2012; 2013). Other Zr phases reported from EGM alteration assemblages include hilairite (Na₂ZrSi₃O₉·3H₂O) and vlasovite (Na₂ZrSi₄O₁₁), from Tamazeght and Pilanesberg, respectively (Khadem Allah *et al.*, 1998; Mitchell and Liferovich, 2006).

Zirconosilicate assemblages are typically characterized as agpaite (containing complex alkali-zirconosilicates, e.g. catapleite/gaidonayite, hilairite or vlasovite) or miaskitic (containing simple zirconosilicates, e.g. zircon or baddelyite). Allanite-(Ce) was found in association with zircon-bearing assemblages, both at Ilímaussaq (Karup-Møller *et al.*, 2010) and other localities (e.g.

TABLE 4. Average compositions of magmatic and secondary aegirine.

Mineral <i>n</i>	Primary aegirine		Secondary aegirine	
	Average 102	2 σ	Average 54	2 σ
Na ₂ O	13.28	0.85	14.52	0.78
SiO ₂	52.70	2.49	53.45	1.86
Al ₂ O ₃	1.18	0.28	1.97	1.89
FeO	29.02	2.09	27.66	2.82
MgO	0.13	0.13	0.09	0.15
MnO	0.15	0.12	0.21	0.33
CaO	2.61	1.24	0.47	1.19
K ₂ O	0.01	0.02	0.02	0.11
TiO ₂	0.72	0.71	0.60	0.92
ZrO ₂	0.31	0.37	0.68	1.43
Total	100.11		99.69	
Formulae based on 4 cations				
Na	0.95		1.04	
Si	1.95		1.97	
Al	0.05		0.09	
Fe ³⁺ *	0.85		0.85	
Fe ²⁺	0.05		0.00	
Mg	0.01		0.01	
Mn	0.00		0.01	
Ca	0.10		0.02	
K	0.00		0.00	
Ti	0.02		0.02	
Zr	0.01		0.01	
O	6.00		5.95	
Morimoto end-members**				
Quad	5.49		1.13	
Jadeite	1.65		8.37	
Aegirine	92.87		90.50	

*Fe³⁺ and Fe²⁺ recalculated to 6 oxygens.

**Based on Morimoto (1988).

Coulson, 1997; Khadem-Allah *et al.*, 1998; Mitchell and Liferovich, 2006), but not with apgaitic alteration assemblages and is therefore characterized as miaskitic. With allanite-(Ce) being part of the gittinsite-bearing paragenesis, the latter can be regarded as transitional-miaskitic (Fig. 6).

Several stages of alteration can lead to multiple generations of superimposed apgaitic and miaskitic assemblages (Mitchell and Liferovich, 2006). The relative stability of the zirconosilicates depends on parameters such as pressure, temperature and the chemical potential of volatiles and other components in the melts and fluids (e.g. Marr and Wood, 1992; Markl *et al.*, 2001; Mitchell and Liferovich, 2006; Mitchell and Chakrabarty, 2012; Andersen *et al.*, 2010; Marks

et al., 2011). As such, EGM decomposition provides a useful tool for monitoring chemical changes in late- and post-magmatic fluids during the late-stage evolution of peralkaline complexes (e.g. Coulson, 1997; Mitchell and Liferovich, 2006; Mitchell and Chakrabarty, 2012). For example, Mitchell and Liferovich (2006) linked the sequence of apgaitic and miaskitic assemblages in the lujavrites of Pilanesberg to changes in the pH, governed by fluctuating Na/Cl in response to fluid-rock reactions involving nepheline, sodalite and EGM.

In addition to physico-chemical controls, the proportion and type of secondary phases, particularly those incorporating the released REE, Y, Fe, Mn, Sr, Nb, Hf and Ti, are governed by the primary

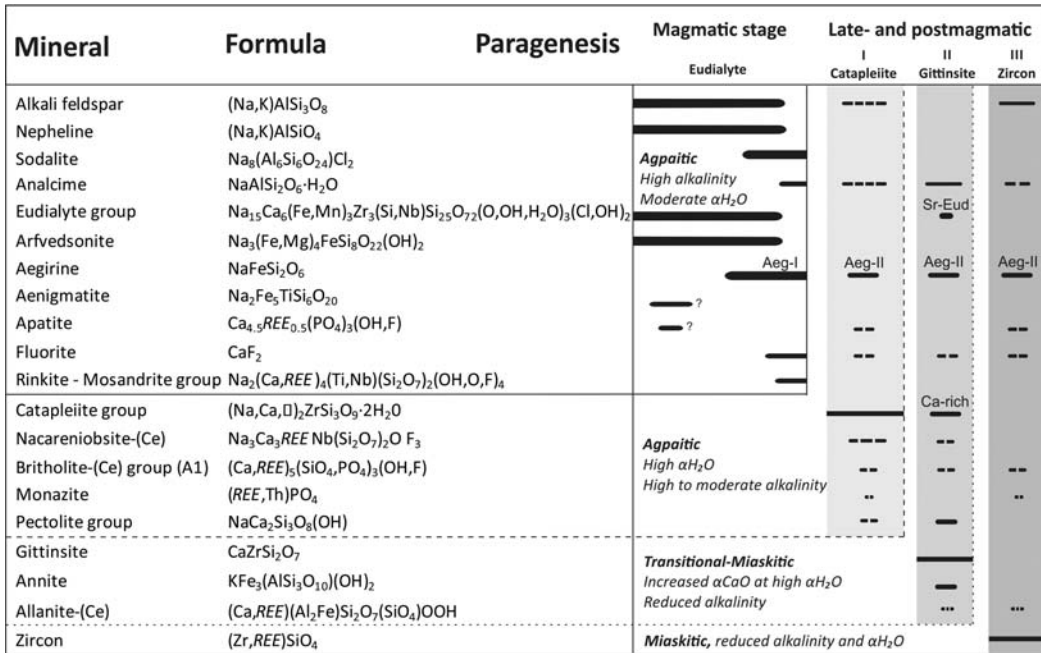


FIG. 6. Generalized paragenetic diagram for the kakortokites and EGM decomposition assemblages. Note that the crystallization sequence at the magmatic stage is slightly different for individual layers of black, red and white kakortokites. EGM alteration parageneses are distinguished based on the Zr phases catapleite, gittinsite and zircon, and are characterized, respectively, as apgaitic, transitional and miaskitic (see text for details). Zircon-bearing assemblages (type-III) are obtained from Karup-Møller *et al.* (2010). Catapleite-bearing assemblages (type-I) are prevalent throughout the kakortokites, whereas type-II and III are spatially restricted. The gittinsite-bearing assemblage (type-II) is superimposed on a type-I paragenesis. Length and continuation of bars indicate relative abundance and proportions within the pseudomorphs.

composition of the replaced EGM. Given the broad compositional range of EGM worldwide (e.g. Schilling *et al.*, 2011) it is not surprising that the proportion and composition of secondary phases in EGM alteration assemblages may be unique for each locality (see overview Mitchell and Chakrabarty, 2012). For example, EGM at Sushina Hill are particularly Fe-poor and Mn-rich and, accordingly, their alteration assemblages contain the Mn-mineral serandite instead of aegirine (Mitchell and Chakrabarty, 2012). Similarly, REE rich hydrothermal EGM at Pilanesberg produce alteration assemblages with high proportions of secondary REE phases such as britholite-(Ce) and fergusonite-(Y) (Olivo and Williams-Jones, 1999).

As described above, EGM compositions in the kakortokite-lujavrite sequence are relatively constant, except for minor fractionation trends in Ca, Mn, Hf, REE and Y contents across the stratigraphy (Pfaff *et al.*, 2008; Lindhuber *et al.*, 2015; this

work, Table 1). These variations, however, appear too small to cause systematic trends in the composition and type of EGM alteration assemblages across the stratigraphy. The only compositional feature which may be inherited from the precursor EGM is a weakly-defined increase in Y and decrease in Hf in catapleite. We infer that most variations in the alteration assemblages reflect different physico-chemical properties of the late-stage magmatic/hydrothermal fluids, and as such present an ideal case for modelling of zirconosilicate stabilities.

Chemographic analyses

Theoretical phase diagrams on the relative stability of catapleite, zircon and EGM in peralkaline assemblages were presented previously by Andersen *et al.*, 2010 and Marks *et al.*, 2011.

These show that increasing H_2O activity, decreased alkalinity (activity of Na in the form of $Na_2Si_2O_5$, Na_2SiO_3 or Na_2O) and lower Cl^- activity favours catapleite stabilization after EGM, whereas zircon formation requires a further decrease in alkalinity as well as a decrease in H_2O activity. The relative stability of Ca-rich gittinsite in these mineral assemblages has not been discussed so far.

To further explore the factors governing the observed assemblages in the kakortokites we applied Schreinemakers chemographic analyses to construct semi-quantitative 3D petrogenetic grids as a function of alkalinity, H_2O and CaO activities in the melt/fluid (see methods in Andersen *et al.*, 2010; 2013 and references therein). This was undertaken for a 13-component system including $NaO_{0.5}$ – $KO_{0.5}$ – CaO – FeO – $REEO_{1.5}$ – ZrO_2 – TiO_2 – $NbO_{2.5}$ – $AlO_{1.5}$ – SiO_2 – $HO_{0.5}$ – $FO_{0.5}$ – $ClO_{0.5}$, combining Mn and Mg with Fe, Hf with Zr, and Y with REE (Table 5a). In our diagrams, alkalinity is expressed by the activity of the water-soluble sodium metasilicate component $Nms = Na_2SiO_3$. Divariant reactions were calculated for zirconosilicate assemblages involving EGM, catapleite, gittinsite, vlasovite and zircon, assuming the co-existing non-HFSE phases analcime, annite, aegirine, pectolite and fluorite (Table 5b). Stoichiometric end-member compositions were used for all phases except catapleite and EGM, for which we used average compositions (Table 5a). Calculations (not presented) using the co-existing non-HFSE phases albite, K-feldspar, pectolite and fluorite (type-I and type-III alteration assemblages, Fig. 6) yielded similar configurations, with only minor differences in the slopes of the divariant planes involving catapleite and EGM.

External buffering was assumed for the components TiO_2 , REE_2O_3 and Nb_2O_5 by phases not included in the analyses (e.g. nacareniobsite-(Ce), fergusonite-(Y), monazite, Al minerals, etc). The model assumes isobaric-isothermal conditions (~ 1 kbar and $450^\circ C$, Markl *et al.*, 2001; Markl and Baumgartner, 2002). Pressure-temperature variations influence the relative sizes of the stability volumes, but not the general topology. The effect of oxygen fugacity was disregarded by excluding ferric iron from the calculations, but is known to increase during the final stages of evolution from Fe-mineral equilibria (Markl *et al.*, 2001). Similarly, the effects of F and Cl are not considered here but are relevant for the discussion on REE and HFSE mobility; i.e. F⁻ activity is assumed at the fluorite saturation level and a general decrease in Cl^- activity is inferred from the fact that all

secondary phases are devoid of Cl (e.g. Andersen *et al.*, 2010; Marks *et al.*, 2011). Finally, Si activity remains well below quartz saturation, corresponding to nepheline-alkali feldspar coexistence.

The calculated reactions were used to construct a grid of divariant planes and univariant lines in 3D $\log a_{Nms} - \log a_{H_2O} - \log a_{CaO}$ space (Fig. 7). Together, they separate trivariant single-phase stability volumes for EGM, zircon, catapleite, vlasovite and gittinsite. Vlasovite is neither observed nor reported from the kakortokites, but was included for comparison with the apgaitic type locality of gittinsite at Kipawa, where gittinsite occurs with apophyllite and vlasovite in EGM-rich pegmatites (Ansell, 1980). The resulting petrogenetic grid is internally consistent and gives a true representation of the relative slopes of the divariant planes and univariant lines, and the relative positions of the stability volumes in 3D log activity space. It is semi-quantitative in the sense that no absolute values can be assigned to the axes or the actual sizes of the stability fields without existing quantitative thermodynamic and experimental data. A petrogenetic grid in log activity space constructed in this way, however, is most useful when predicting the consequences of changes in component activities in melt or fluid. A grid configuration that agrees with the observed mineral assemblages is presented in Fig. 7.

Alteration of EGM to catapleite

Eudialyte-group minerals are stable at high levels of alkalinity and variable H_2O and CaO activities, whereas catapleite becomes the dominant zirconosilicate at high H_2O activity (Fig. 7). The reaction from EGM to catapleite (reaction 1 in Table 5b) is represented by the divariant plane between the red and blue volumes at high a_{H_2O} and a_{Nms} in Fig. 7 and can be achieved by either: (1) increasing the a_{H_2O} ; (2) decreasing the alkalinity in a system at high a_{H_2O} ; or (3) increasing a_{CaO} in a system at high a_{H_2O} and a_{Nms} (Fig. 7). Assuming that most EGM alteration occurred at relatively low to medium a_{CaO} , increasing water activity appears to be the dominant factor. This can be illustrated in a 2D section perpendicular to the $\log a_{CaO}$ axis, showing relative stabilities as a function of water activity and alkalinity (Fig. 8). The section is positioned just above the vlasovite field, which is restricted to lower a_{CaO} and a_{H_2O} . Given the presence of Ca phases fluorite, apatite and pectolite in both magmatic and post-magmatic assemblages, a_{CaO} may be too high to allow stabilization of vlasovite in

TABLE 5a. Compositions of minerals and components (in moles per formula unit) used in modelling.

Minerals	NaO _{0.5}	KO _{0.5}	CaO	FeO	FeO _{1.5}	REEO _{1.5}	ZrO ₂	TiO ₂	NbO _{2.5}	AlO _{1.5}	SiO ₂	HO _{0.5}	FO _{-0.5}	ClO _{-0.5}
Zirconosilicates														
Eudialyte*	14.04	0.20	5.68	2.78	0.00	0.48	2.85	0.02	0.21	0.16	25.74	2.00	0.00	1.29
Catapleite	1.20	0.00	0.40	0.00	0.00	0.00	1.00	0.00	0.00	0.00	3.00	4.00	0.00	0.00
Gitinsite	0.00	0.00	1.00	0.00	0.00	0.00	1.00	0.00	0.00	0.00	2.00	0.00	0.00	0.00
Zircon	0.00	0.00	0.00	0.00	0.00	0.00	1.00	0.00	0.00	0.00	1.00	0.00	0.00	0.00
Vlasovite	2.00	0.00	0.00	0.00	0.00	0.00	1.00	0.00	0.00	0.00	4.00	0.00	0.00	0.00
Non-HFSE phases														
Analcime	1.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.00	2.00	2.00	0.00	0.00
Aegirine	1.00	0.00	0.00	0.00	1.00	0.00	0.00	0.00	0.00	0.00	2.00	0.00	0.00	0.00
Pectolite	1.00	0.00	2.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	3.00	1.00	0.00	0.00
Annite	0.00	1.00	0.00	3.00	0.00	0.00	0.00	0.00	0.00	1.00	3.00	2.00	0.00	0.00
Fluorite	0.00	0.00	1.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.00	0.00
Melt/fluid components														
Na ₂ SiO ₃	2.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	1.00	0.00	0.00	0.00
CaO	0.00	0.00	1.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
H ₂ O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	2.00	0.00	0.00

*Average of 118 EGM core analyses.

TABLE 5b. Divariant reactions between zirconosilicates, co-existing minerals and components in melt/fluid.

Divariant reactions

- 1 **0.35 Eud** + 0.02 Anl + 2.50 H₂O + 0.67 CaO + 0.19 O₂ = **1.00 Cat** + 1.13 Pct + 0.07 Ann_{bi} + 0.76 Aeg + 0.01 TiO₂ + 0.04 Nb₂O₅ + 0.08 REE₂O₃ + 0.92 Nms + 0.45 HCl
- 2 **0.35 Eud** + 0.01 Anl + 0.78 H₂O + 1.39 CaO + 0.19 O₂ = **1.00 Zrc** + 1.69 Pct + 0.07 Ann_{bi} + 0.76 Aeg + 0.01 TiO₂ + 0.04 Nb₂O₅ + 0.08 REE₂O₃ + 1.24 Nms + 0.45 HCl
- 3 **0.35 Eud** + 0.01 Anl + 0.58 H₂O + 1.59 CaO + 0.19 O₂ = **1.00 Git** + 1.29 Pct + 0.07 Ann_{bi} + 0.76 Aeg + 0.01 TiO₂ + 0.04 Nb₂O₅ + 0.08 REE₂O₃ + 1.44 Nms + 0.45 HCl
- 4 **0.35 Eud** + 0.01 Anl + 0.38 H₂O + 0.19 O₂ = **1.00 Vla** + 0.89 Pct + 0.07 Ann_{bi} + 0.76 Aeg + 0.01 TiO₂ + 0.04 Nb₂O₅ + 0.08 REE₂O₃ + 0.64 Nms + 0.21 CaO + 0.45 HCl
- 5 **1.00 Zrc** + 0.56 Pct + 1.72 H₂O + 0.32 Nms = **1.00 Cat** + 0.72 CaO
- 6 **1.00 Zrc** + 0.40 Pct + 0.20 CaO = **1.00 Git** + 0.2 H₂O + 0.2 Nms
- 7 **1.00 Zrc** + 0.80 Pct + 0.60 Nms = **1.00 Vla** + 0.4 H₂O + 1.6 CaO

Bold indicates zirconosilicates.

Abbreviations given in Table 5a.

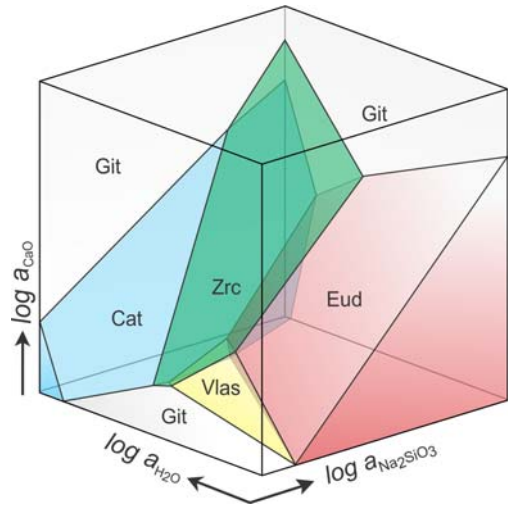


FIG. 7. Semi-quantitative 3D petrogenetic grid in $\log a_{\text{Nms}} - \log a_{\text{H}_2\text{O}} - \log a_{\text{CaO}}$ space illustrating the relative stability volumes for the Na-Ca zirconosilicates eudalyte (red), catapleite (blue), gittinsite (grey transparent), zircon (green) and vlasovite (yellow). The divariant reactions, calculated for the system of $\text{NaO}_{0.5}\text{-KO}_{0.5}\text{-CaO-FeO-REEO}_{1.5}\text{-ZrO}_2\text{-TiO}_2\text{-NbO}_{2.5}\text{-AlO}_{1.5}\text{-SiO}_2\text{-HO}_{0.5}\text{-FO}_{0.5}\text{-ClO}_{0.5}$ (Table 5a, b), each define a 2-phase stability plane in 3D log activity space. The stability field for gittinsite overlies all other stability fields at high $\log a_{\text{CaO}}$, low $\log a_{\text{H}_2\text{O}}$ and low $\log a_{\text{Nms}}$. Divariant planes between gittinsite and other phases are left semi-transparent to illustrate the 3D geometry of the stability volumes below.

the kakortokites. The stability of vlasovite in the kakortokites may be limited additionally by the relatively low T during hydrothermal alteration (<450°C), i.e. in Si-saturated systems vlasovite typically represents a high T assemblage (>600°C at 1 kbar, e.g. Currie and Zaleski, 1985; Marr and Wood, 1992; Birkett *et al.*, 1992).

Alteration of EGM and catapleite to gittinsite

The gittinsite stability volume (grey transparent field, Fig. 7) overlies all other fields at high $\log a_{\text{CaO}}$. Based on the abundance of analcime, the presence of a crosscutting hydrothermal vein and the fact that gittinsite overprints a catapleite assemblage, we assume that gittinsite formation also occurred at elevated levels of water activity, i.e. most probably in the presence of a free fluid phase. This can be illustrated with a vertical 2D section at high $\log a_{\text{H}_2\text{O}}$ (Fig. 9). Assuming constant water activity, gittinsite can stabilize after catapleite following: (1) a reduction in alkalinity; (2) an increase in CaO activity; or (3) a combination of

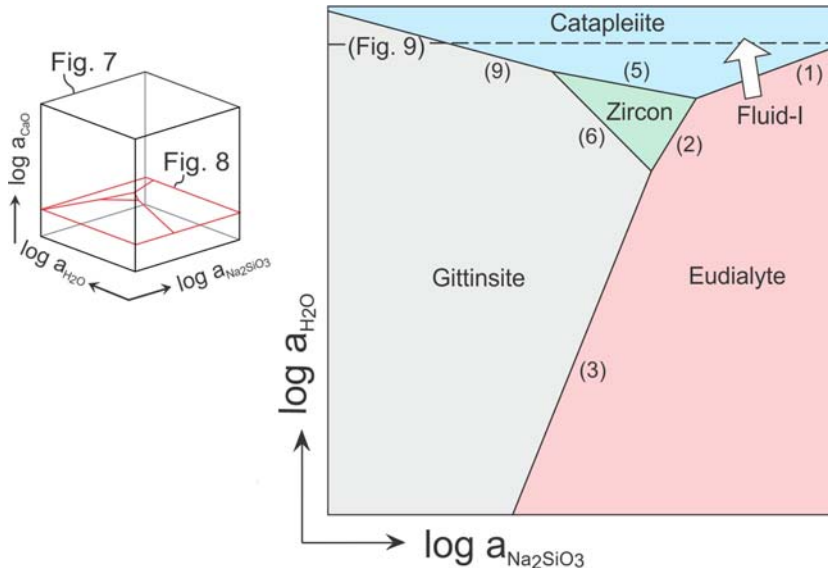


FIG. 8. Horizontal section through the petrogenetic grid in Fig. 7, illustrating the relative stability fields of EGM and alteration products as a function of $\log a_{\text{H}_2\text{O}}$ and $\log a_{\text{Nms}}$ at constant CaO activity. Catapleite formation is governed by increasing $a_{\text{H}_2\text{O}}$ (arrow, Fluid I). The relative location with respect to Fig. 7 is indicated on the right and lies just above the vlasovite stability volume (see text for details). Lines are labelled with reaction numbers in Table 5b. The dashed line marks the location of the vertical section in Fig. 9, which represents a theoretical level of water saturation.

both (Fluid-II in Fig. 9). Because of abundant pectolite and fluorite in the gittinsite-paragenesis and the rare occurrence of this paragenesis in general, we infer that an overall input of calcium was required to form gittinsite at the expense of EGM and catapleite. This was probably associated with a simultaneous decrease in alkalinity, thereby avoiding the zircon stability field prior to gittinsite stabilization (Fluid-II, Fig. 9). Direct replacement of EGM by gittinsite without the intermediate step of catapleite formation (reaction 3 in Table 5b and Fig. 8) requires an increase in a_{CaO} or decrease in a_{Nms} at locally reduced water activity, and probably reflects incomplete equilibration of EGM at higher water activities. Rims of Sr-rich secondary EGM in the gittinsite paragenesis (Fig. 5f,g,h) suggest that this fluid also added Sr.

Both Ca and Sr are depleted with continued magmatic evolution, i.e. of all intrusive units, the augite syenite contains the highest concentrations of Ca and Sr (c. 3–6 wt.% and 300–400 ppm, respectively; Bailey *et al.*, 2001). As such, alteration of earlier magmatic units, such as the augite syenites entrained as autoliths in unit +3 (Fig. 2) where gittinsite was found, could provide a potential source for Ca- and Sr-bearing fluids. Alternative mechanisms for a relative increase in

CaO activity and/or decrease in alkalinity within internal fluids may involve late- to post-magmatic breakdown of other Ca-bearing phases, such as apatite or fluorite, or segregation of complementary alkali-rich and alkali-poor fluids/vapours (e.g. Shmulovich and Churakov, 1998; Markl and Baumgartner, 2002). Such processes, however, may not satisfactorily explain the addition of Sr. The extremely rare occurrence of gittinsite in the kakortokites and agpaitic rocks in general is probably related to the overall high alkalinity, and the high buffering capacity of Ca in F-rich hydrothermal environments by precipitation of fluorite and phases like nacareniobsite-(Ce).

Limited evidence for Ca metasomatism by external fluids at Ilímaussaq includes a study of endo-skarn assemblages containing Ca-rich garnet, epidote (or allanite) and ilvaite (Graser and Markl, 2008). These assemblages were suggested to have formed from seawater which percolated through the overlying basalts and sandstones and entered the complex along fractures and faults upon cooling. Although the gittinsite-sample was found at a considerable distance from the margin of the complex, such a process could explain the input of both Ca and Sr. Calcic metasomatism has also been suggested as the main mechanism for

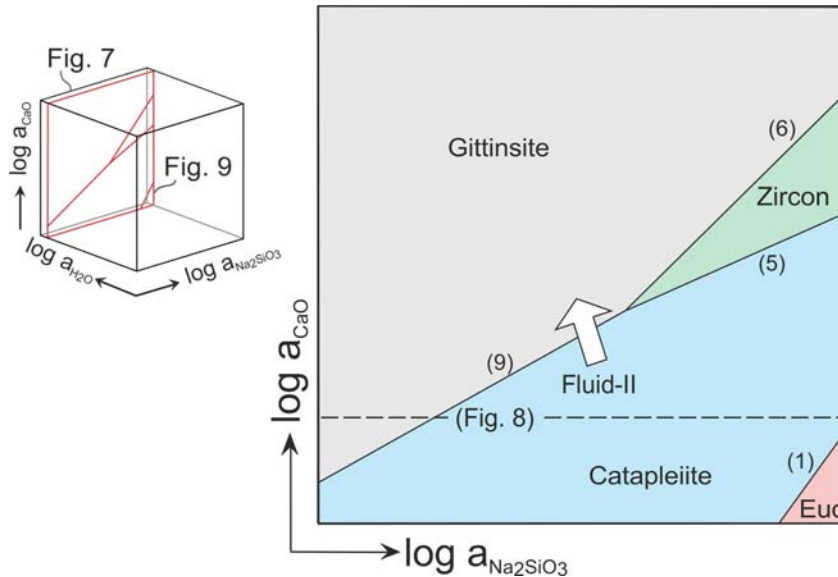


FIG. 9. Vertical section through the petrogenetic grid in Fig. 7, illustrating the relative stability fields of EGM and alteration products as a function of $\log a_{\text{CaO}}$ and a_{Nms} at constant $a_{\text{H}_2\text{O}}$. Lines are labelled with reaction numbers in Table 5b. Gittinsite formation at the expense of catapleite at high H_2O activity (or saturation) can be achieved by decreasing a_{Nms} and increasing a_{CaO} (arrow, Fluid-II).

gittinsite formation in some peralkaline granites (sometimes rock forming, e.g. Strange Lake), where gittinsite commonly replaces zircon or elpidite ($\text{Na}_2\text{ZrSi}_6\text{O}_{15} \cdot 3\text{H}_2\text{O}$) together with armstrongite ($\text{CaZrSi}_6\text{O}_{15} \cdot 3\text{H}_2\text{O}$) and quartz (i.e. high-Si zirconosilicate assemblages; Marks *et al.*, 2011). In these systems, Ca-metasomatism was associated with significant hydrothermal remobilization of primary Zr, Nb and REE and involved mixing of external Ca-rich fluids with late-magmatic F-rich fluids (e.g. Salvi and Williams-Jones, 1996; 2005; 2006; Kynicky *et al.*, 2011).

Alteration of EGM to zircon

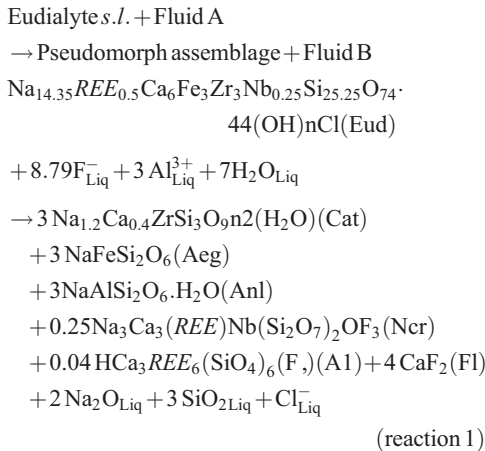
The zircon stability volume (green, Fig. 7) is bounded by EGM at high a_{Nms} and low a_{CaO} , catapleite at high $a_{\text{H}_2\text{O}}$ and gittinsite at high a_{CaO} and lower $a_{\text{H}_2\text{O}}$. The diagram suggests that an increase in a_{CaO} during gittinsite formation could also lead to zircon formation after both EGM and catapleite. The fact that zircon was not observed in EGM alteration assemblages of the main kakortokites may suggest that the zircon field is smaller than currently presented in Fig. 7, or simply absent due to variations in temperature, pressure, Si activity or other parameters not considered here. In that case, the zircon-stability volume is replaced

by a univariant line between catapleite, gittinsite and EGM. The general topology in Fig. 7, however, implies that the zircon-bearing assemblages do not reflect a continuous alteration sequence by progressive fluid-rock reactions, but rather reflect localized and separate melt/fluid conditions which were achieved only in pegmatitic and more evolved/altered apgaitic units.

Mass-balance and volume relations

The observed and reported EGM decomposition assemblages from Ilímaussaq suggest that most elements were retained within the precursor EGM domains, with Zr sequestered in catapleite, gittinsite or zircon, Nb in nacareniobsite-(Ce), fergusonite-(Y) and other Nb phases, and Fe in aegirine. Nacareniobsite-(Ce) additionally acts as a major host for released REE, while the remaining REE form fine-grained Ca-(Ba)-REE silicates (A1), or alternatively monazite when additional apatite is present. Using the most common type-I alteration assemblage (i.e. Cat + Aeg + Anl + Ncr + Ca-rich A1 + Fl, adapting a simplified EGM formula and treating REE as a single component) the alteration reaction can be expressed by the mass- and volume-

balanced equation (abbreviations as in Fig. 3):



Note that the equation is not charge-balanced and semi-balanced for O and H due to the unconstrained fluid speciation and volatile contents of EGM. The growth of catapleite, aegirine and nacareniobsite-(Ce) is confined by the availability of Zr, Fe and Nb, respectively. Their theoretical relative volumes within the pseudomorph volume can be calculated using the *Z*-values and unit-cell volumes, yielding proportions of 40%, 18% and 5% for catapleite, aegirine and nacareniobsite-(Ce), respectively. These are in good agreement with visual estimates in most pseudomorph assemblages (Fig. 3*g,h*). Residual *REE* are contained within A1 (unit-cell volumes unknown), while Ca re-precipitated as fluorite by addition of F. If all Ca is retained in this way, fluorite should account for ~9% of the pseudomorph volume. The proportions of fluorite, however, rarely exceed 5%. Some of the remaining Ca may be used to form pectolite or apatite, or is lost from the EGM micro-domains. The absence of secondary Cl phases in or around the pseudomorphs, i.e. no sodalite rims on primary nepheline (as observed in the Igaliko and Sushina complexes; Finch, 1991 and Mitchell and Chakrabarty, 2012, respectively), implies that all Cl was released to the fluid phase.

In all pseudomorph assemblages the remaining interstitial volume is occupied by analcime, albite and/or K-feldspar, requiring significant input of aluminum. In reaction 1, the amount of analcime is calculated based on a remaining pseudomorph volume of 27% (Cat + Aeg + Ncr + Fl = 73%, neglecting A1). A similar calculation for albite-dominated assemblages requires 2.5 mol of albite per mol of replaced EGM to fill 27% of the total

pseudomorph volume. This reduces the required input of Al and H₂O to 2.5 and 4 mol., respectively. Although Al is often considered to be immobile during fluid-rock interactions, it was undoubtedly soluble in the system considered here. A potential source of Al is the associated constant-volume replacement of primary nepheline and alkali feldspar by analcime (Markl and Baumgartner, 2002). As EGM contain very little P and Ba, these elements must additionally be added to form EGM alteration assemblages containing monazite, apatite, britholite-(Ce) and Ba-*REE* silicates. The restricted occurrence of these assemblages suggests that P and Ba contents in the fluids were highly variable, which may reflect local interaction of the fluids with (or precipitation of) phases enriched in these elements, e.g. apatite or K-feldspar.

In summary, the observed relative proportions combined with estimated volumes of secondary phases from reaction 1 suggest that the alteration of EGM to catapleite-type assemblages was not associated with significant remobilization of *REE*, Zr and Nb. The reaction reflects a net uptake of F, H₂O, Al, and occasionally P and Ba from the fluid phase, whereas Na, Si and Cl were lost from the assemblage.

The sequential alteration of EGM, first to catapleite and then to gittinsite, is associated with a significant volume reduction. With preservation of Zr, gittinsite consumes 57% of the replaced catapleite volume, and only 23% of the total EGM volume. This again correlates well with the observed proportions of gittinsite and the comparatively large proportion of analcime-filled ‘pore space’ in the gittinsite-bearing pseudomorphs (Fig. 5). In terms of *REE*, however, the gittinsite-bearing pseudomorphs contain relatively small proportions of *REE*-bearing phases in general, and a comparably high proportion of heavy rare-earth element (*HREE*)/Y-bearing phases (Fergusonite-(Y) and unidentified Ca-*HREE*-Y minerals, Fig. 5*c,d*) with respect to the catapleite-type paragenesis. This could suggest that the gittinsite-forming fluid was more capable of fractionating the light and heavy *REE*, potentially reflecting a difference in pH and/or F and Cl contents between the two fluids.

Mobility of HFSE and REE and comparison to peralkaline granites

Although HFSE and *REE* are generally considered to be highly immobile in hydrothermal systems,

studies of HFSE-*REE* mineralization in peralkaline granites and syenites such as Strange Lake, Thor Lake, Tamazeght, Khan Bogd and Khaldzan Buregte have shown that late-magmatic hydrothermal fluids enriched in ligands such as Cl^- , F^- , PO_4^{3-} , SO_4^{2-} and OH^- can lead to significant remobilization of primary HFSE and *REE* (e.g. Boily and Williams-Jones, 1994; Salvi and Williams-Jones, 1996; 2005; 2006; Salvi *et al.*, 2000; Kynicky *et al.*, 2011; Sheard *et al.*, 2012; Gysi and Williams-Jones, 2013). Recent numerical and experimental results have shown that *REE* (especially light rare-earth elements (*LREE*)) are most effectively transported as Cl complexes (e.g. Migdisov *et al.*, 2009; Migdisov and Williams-Jones, 2014), whereas Zr and Nb remobilization is most effective via complexation with mixed F-OH compounds (e.g. Migdisov *et al.*, 2011; Timofeev *et al.*, 2015).

It is clear that most EGM alteration in the kakortokites occurred by aqueous fluids carrying significant amounts of F and Cl, the latter being released to the fluid upon EGM breakdown. By analogy with the peralkaline granites listed above, this could provide a potential mechanism for significant hydrothermal remobilization of the primary *REE*-HFSE. However, in the presence of F, the mobility of *REE*-chloride complexes is strongly limited to fluids with low pH (<4 at 400°C) due to precipitation of insoluble *REE*-fluoride phases at higher pH (e.g. Migdisov *et al.*, 2009; Migdisov and Williams-Jones, 2014; Gysi and Williams-Jones, 2013). Markl and Baumgartner (2002) argued that late-stage fluids in Ilimaussaq are characterized by neutral to high pH and that the pH is strongly controlled by the Na to Cl ratio (pH > 8 at Na/Cl > 1). With the inferred insolubility of *REE*-F species at neutral to high pH (Migdisov and Williams-Jones, 2014), the released *REE*, Nb and Ca from the EGM micro-domains would immediately react with fluorine from the fluid to form fluorite, Al and nacareniobsite-(Ce). This would subsequently lead to a decrease in the F^- activity of the fluid, and thereby also decrease the potential mobility of Zr and Nb as fluoride and hydroxyl-fluoride complexes, even if such complexes would be stable at the considered temperature and pH (Migdisov *et al.*, 2011; Timofeev *et al.*, 2015).

The influence of pH on the solubility of *REE*-chloride and -fluoride complexes (e.g. Migdisov *et al.*, 2009; Migdisov and Williams-Jones, 2014; Gysi and Williams-Jones, 2013) may be crucial in understanding why fluid-driven remobilization of

HFSE and *REE* is relatively ineffective with respect to orthomagmatic enrichment processes in highly Si-undersaturated and peralkaline systems such as the kakortokites, as opposed to inherently more acidic systems such as the peralkaline granites of Thor Lake, Strange Lake and Khan Bogd (e.g. Sheard *et al.*, 2012; Gysi and Williams-Jones, 2013; Kynicky *et al.*, 2011) or related skarn-deposits (e.g. Ambohimirahavavy complex in Madagascar, Estrade *et al.*, 2015). In granitic systems, magmatic fluids enriched in F and Cl are more likely to evolve towards low pH and may thus be more capable of remobilizing HFSE and *REE* and concentrating them to potentially economic levels when magmatic mineralization processes are insufficient.

Conclusions and economic implications

Pseudomorph assemblages after primary eudialyte-group minerals in the kakortokites of the Ilimaussaq complex are invariably dominated by the Na zirconosilicate catapleite. The catapleite-assemblages record an overall increase in H_2O activities, which reflects the presence of Na-, Cl- and F-rich aqueous fluids at the final stages of kakortokite crystallization following a near-continuous evolution from melt to fluid. EGM alteration is slightly more pervasive in the TLK sequence and lujavrites than in the main LLK sequence, which probably reflects higher volatile contents in the more evolved melts from which they crystallized. The pseudomorph assemblages generally require an input of H_2O , Al and F, and occasionally Ba and P, to the EGM micro-domains, whereas Cl was released to the fluid phase. Zircon-bearing EGM alteration assemblages were not encountered in the current sample set, confirming their restricted occurrence to pegmatitic and more strongly altered units of the complex (Karup-Møller *et al.*, 2010). Zircon formation is therefore interpreted to record localized melt-fluid evolution in units not representative for the main kakortokite body. Similarly, stabilization of the rare Ca zirconosilicate gittinsite, associated with analcime, pectolite, fluorite, annite, Ca-*REE* silicates and allanite-(Ce), was triggered by a local increase in CaO activity along with a complementary decrease in alkalinity and is inferred to result from a separate, low *T* post-magmatic alteration event by a Ca- and Sr-bearing aqueous fluid.

Alteration of EGM has primarily affected the HFSE and *REE*-mineralization through increased

dissemination and fractionation of the prospective elements among the secondary minerals, while leaving the overall ore grade and resources unaffected. We infer that the immobile behaviour of HFSE and REE during autometasomatic alteration of EGM in the kakortokites relates to a relatively high pH (<8) of the late-magmatic fluids, triggering immediate precipitation of F-bearing phases such as fluorite, nacareniobite-(Ce) and Ca-REE silicates (A1) upon interaction of fluorine with the released Ca, REE and Nb. The resulting decrease in F⁻ activity would subsequently prevent the potential mobility of Zr and Nb as F and mixed OH-F complexes. A somewhat lower pH for the post-magmatic gittinsite-forming fluid compared to the late-magmatic catapleiite-forming fluids could explain the minor fractionation between LREE and HREE by preferential remobilization of LREE as chloride-complexes.

Despite the fact that most HFSE and REE are retained within the pseudomorphs, the mineralogical changes associated with the wide-spread decomposition of EGM in the kakortokites led to an increased complexity of the mineralization. A significant reduction in grain size (<10 µm) and different physical properties for the secondary Zr, Nb and REE phases may complicate liberation or chemical separation of the metals from the rock. The latter was observed during dissolution studies of EGM concentrates from the Lovozero complex (Russia), where the presence of secondary Na-Zr silicates resulted in low Zr recovery (Zakharov *et al.*, 2011). As such, further studies on the mineralogical and physical properties of EGM alteration assemblages will be of great value for optimal beneficiation of the kakortokites and other EGM occurrences, and enhance their potential value as future ore deposits.

Acknowledgements

This research was carried out as part of AMB's PhD project with financial support from the Geocenter Denmark (project number 4/2014) and benefited from collaborations in the EURARE project, funded by the European Community's Seventh Framework Programme (FP7/2007–2013, grant 309373). The authors thank Henning Bohse for kindly hosting and assisting them during field sessions at Kringlerne, Muriel Erambert for help with EPMA analyses in Oslo and John Bailey for providing a collection of samples. Per Kalvig, Alexander Bartels, Lotte Larsen and Quinten van der Meer are thanked for fruitful discussions. Constructive comments from reviewers

Stefano Salvi, Ian Coulson and Principal Editor Roger Mitchell are gratefully acknowledged and contributed to a significantly improved manuscript.

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