

COMPARISON OF ^{14}C COLLECTED BY PRECIPITATION AND GAS-STRIP METHODS FOR DATING GROUNDWATER

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ABSTRACT. Dissolved inorganic carbon (DIC) in groundwater is used to estimate the residence time based on radiocarbon concentration. DIC is usually extracted by a gas-strip or precipitation (SrCO_3 or BaCO_3) method. In this study, the gas-strip and precipitation methods of DIC were applied to both artificially prepared NaHCO_3 solutions and natural groundwater to estimate the certainty of the two methods for ^{14}C dating. ^{14}C values obtained by the gas-strip method for NaHCO_3 solutions with distinct salinity, DIC, and ^{14}C concentrations were close to the theoretically predicted ^{14}C value based on the ^{14}C value of NaHCO_3 powder. Conversely, the ^{14}C value obtained by the precipitation method always showed higher values than the predicted values. The difference in ^{14}C value between the gas-strip and precipitation methods was assumed to be caused by the contamination of modern carbon in the NaOH solution used in the precipitation method. The contamination of modern carbon derived from the NaOH solution during precipitation was found to range from less than 1 mg/L to about 1 mg/L. The applicability of the precipitation method for groundwater should be considered carefully according to the DIC, ^{14}C concentration of groundwater, and purpose of the study being conducted.

KEYWORDS: groundwater, dissolved inorganic carbon, precipitation, gas strip.

INTRODUCTION

The radiocarbon value of dissolved inorganic carbon (DIC), which is one of the most powerful tools for dating groundwater samples, has been applied in many previous studies (Geyh 2000; Iwatsuki et al. 2000). DIC in water samples is usually extracted and collected by either a gas-strip method or a precipitation method. Both of these methods have frequently been used for groundwater studies (gas-strip method: Dorsett et al. 2011; Stewart 2012; Dulinski et al. 2013; precipitation method: Sültenfuß et al. 2011; Stewart 2012; Corcho Alvarado et al. 2013).

In the precipitation method, water samples are alkalized with NaOH or ammonium followed by the addition of BaCl_2 or SrCl_2 , after which DIC is precipitated as BaCO_3 or SrCO_3 . These precipitation procedures are occasionally conducted during groundwater sampling in the field. This procedure enables the amount of samples to be reduced significantly, and facilitates the transport of the samples from the field to the laboratory. However, the precipitation might be contaminated by the alkaline solution because CO_2 readily dissolves in alkaline solutions (Aggarwal et al. 2014). In addition, it could be difficult to extract carbon from groundwater samples with very low DIC concentrations because sometimes precipitation cannot be visually confirmed in such samples and this makes the collection of the precipitation difficult. Furthermore, precipitation may not quantitatively remove carbon due to interfering reactions (Minami and Takahashi 2015). In such cases, the ^{14}C might be fractionated. Possible contamination and potential ^{14}C fractionation need to be quantitatively determined to assess the reliability of the precipitation method in groundwater DI^{14}C dating.

In the gas-strip method, water samples are acidified by phosphoric acid and DIC is collected as CO_2 (Atekwana and Krishnamurthy 1998). This method can be applied to a wide variety of groundwater samples, although they must be brought to the lab for processing. This can be the

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only method to collect DIC in cases where the precipitation method cannot be applied. To validate the method, the possibility of contamination during the gas-strip extraction and DIC recovery from solutions with low DIC concentration needs to be investigated.

In this study, the reliability of these two methods in the determination of ^{14}C groundwater ages is examined, with special attention given to potential contamination. To accomplish this, (1) a preliminary comparison of ^{14}C values obtained by the two preparation methods using a set of NaHCO_3 solutions was made, (2) contamination in sample preparations was identified, and (3) the two methods were applied to natural groundwater samples.

METHODS

Preliminary Comparison of ^{14}C Values from Two Preparation Methods Using NaHCO_3 Solutions

A comparison of ^{14}C results obtained by the two preparation methods for NaHCO_3 solutions allows us to determine the precision and accuracy of each preparation method. This is accomplished by preparing ^{14}C DIC standards, given the DIC concentrations of the solutions and ^{14}C content of the NaHCO_3 powder. Four types of NaHCO_3 standard solutions were prepared (Table 1), and both the gas-strip and precipitation methods were applied to extract and collect DIC from them. The ^{14}C content of the DIC extracted from the solutions was measured by accelerator mass spectroscopy (AMS). Solutions 1–4 were prepared to simulate the following: (1) an old (~25,000 BP) groundwater with low salinity, (2) younger (~10,000 BP) groundwater with low salinity, (3) younger groundwater with low DIC and salinity, and (4) younger groundwater with high salinity. For each solution, two samples, one obtained by gas-strip and one by precipitation, were compared.

Table 1 Solutions for preliminary comparison of ^{14}C values obtained by the two preparation methods.

Solution name	DIC concentration (mg/L)	NaHCO_3 used for preparation	Salinity (g/L)	Extraction method	Sample name for ^{14}C measurement
Solution 1	9.0	Purchased (with low ^{14}C concentration)	0.06	Gas-strip	Sample1-1-1
				precipitation	Sample1-2-1
Solution 2	9.0	Prepared* (with relatively high ^{14}C concentration)	0.06	Gas-strip	Sample1-2-2
				precipitation	Sample2-1-1
Solution 3	4.5	Prepared* (with relatively high ^{14}C concentration)	0.03	Gas-strip	Sample2-2-1
				precipitation	Sample2-2-2
Solution 4	9.0	Prepared* (with relatively high ^{14}C concentration)	41.16	Gas-strip	Sample3-1-1
				precipitation	Sample3-2-1
					Sample3-2-2
					Sample4-1-1
					Sample4-2-1
					Sample4-1-2
					Sample4-2-2

* NaHCO_3 used for Solutions 2 to 4 was prepared as follows: NaOH solution was bubbled with air to obtain precipitation of NaHCO_3 and this NaHCO_3 was mixed with the purchased solution.

All procedures except weighing the NaHCO₃ powder were carried out in a glove box (GB) under an Ar atmosphere. A CO₂ absorbent (Lithoryme: Allied Healthcare Products, St. Louis, MO, USA) was placed in the GB to ensure the CO₂ concentration was low. Deionized water was placed in the GB and purged with Ar gas for at least 12 hr to reduce the DIC in the water. A fraction of this water was used for the DIC measurements after purging. NaHCO₃ powder was subsequently added to the purged water and stirred until completely dissolved. A commercial NaHCO₃ powder was used for Solution 1, while a second batch of NaHCO₃ was prepared by bubbling air through NaOH for Solutions 2–4. In the case of Solution 4, a simulated seawater (SSW) was prepared by adding NaCl, KCl, MgSO₄ 7H₂O, MgCl₂ 6H₂O, and CaCl₂ into the purged water and a NaHCO₃ solution was prepared by dissolution of NaHCO₃ into the SSW.

The four NaHCO₃ solutions were sent to Beta Analytic (Miami, FL, USA) for analysis. Two sample bottles were used for the gas-strip method, while the remaining samples were extracted by the precipitation method for ¹⁴C analyses. Both the DIC collection and ¹⁴C measurements were conducted by Beta Analytic.

Identification of Contamination during Sample Preparation

A preliminary comparison of ¹⁴C in the NaHCO₃ solutions showed the possibility of significant contamination with modern carbon during the precipitation procedure (see Results and Discussion). Therefore, a gas-strip line was employed to test the NaOH and SrCl₂ solutions used in the DIC precipitation procedure. This step was taken to identify the source and amount of contamination associated with the precipitation method. The NaOH and SrCl₂ solutions were prepared as follows. Dissolved gas in deionized water was purged with Ar in a GB, as described above. Two types of NaOH solutions were then prepared, one under atmospheric conditions and the other in the GB with the addition of granular NaOH to deionized water. SrCl₂ solutions were subsequently prepared by adding the SrCl₂ reagent to deionized water under atmospheric conditions.

DIC extraction using the gas-strip method was optimized by application of a vacuum gas-strip line (Figure 1) in the JAEA (Japan Atomic Energy Agency) and the following procedure was established. Prior to sample preparation, the air in the gas-strip line was removed under a vacuum of less than 0.1 Pa. Next, 2 mL of phosphoric acid was added to the water sample to convert the DIC to CO₂ gas. CO₂ gas was subsequently stripped from the water sample by bubbling with carrier gas (pure N₂ or Ar) at a flow rate of 500 cc/min and collected in a cold trap of liquid N temperature. After 5 min of gas circulation, the carrier gas was evacuated and trapped CO₂ was purified and directed into a glass tube at the gas sampling port. The yield of CO₂ relative to DIC in the water samples was 90%. The DIC concentration in NaOH and SrCl₂ solutions were estimated by the gas (CO₂) pressure in the gas-strip line.

Application of the Two Methods to Natural Groundwater Samples

Both the precipitation and gas-strip methods were applied to natural groundwater samples to identify ¹⁴C differences between each method. Groundwater samples were obtained from depths of 200–500 m in granitic rock at the Mizunami Underground Research Laboratory (MIU), Gifu Prefecture, Japan (Figure 2). The groundwater chemistry was Na-Ca-Cl or Na-Cl dominant with a salinity of less than 1 g/L (Iwatsuki et al. 2005, 2015). Groundwater samples were collected into 1-L airtight glass vessels for precipitation. Next, 10 mL of 5N NaOH solution and 10 mL of 2N SrCl₂ solution were added into the vessel to precipitate SrCO₃, after which the samples were stored at 4°C. Precipitated SrCO₃ was collected into small

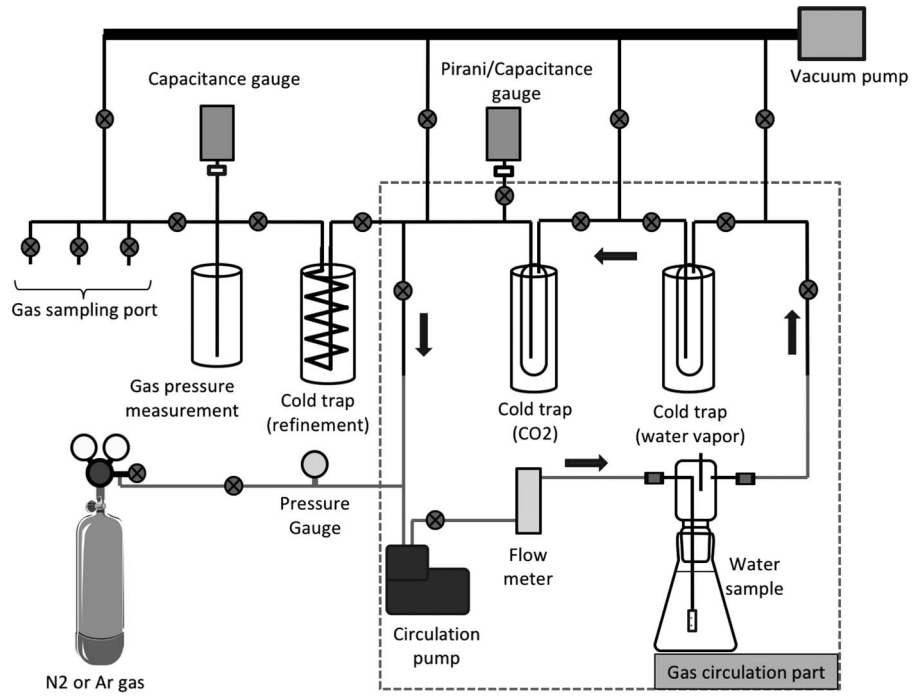


Figure 1 Gas-strip line to extract DIC from groundwater

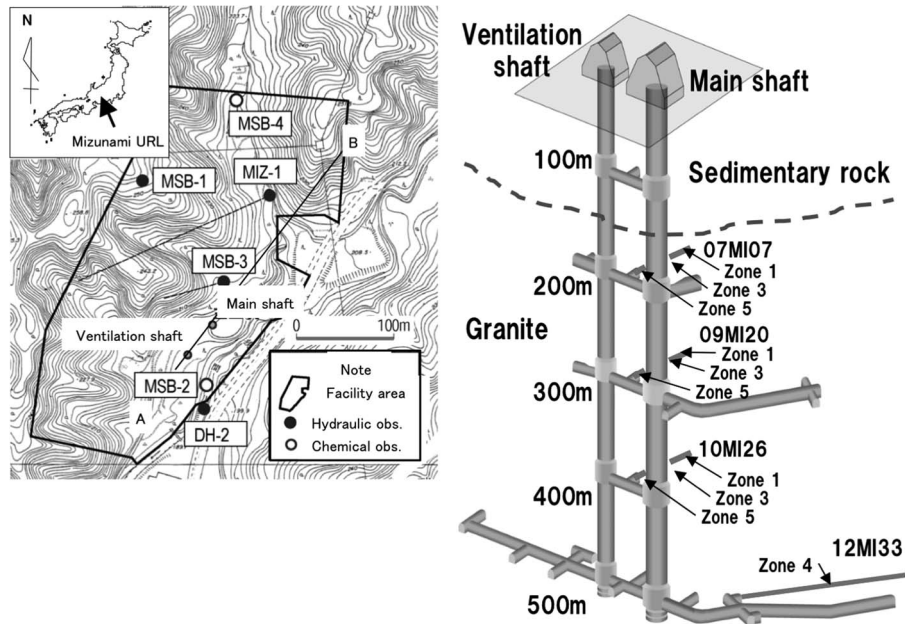


Figure 2 Groundwater sampling point at Mizunami underground research laboratory

sample vessels by pipette in a CO₂-free GB for freeze-dry storage. In parallel, groundwater samples for gas-strip method preparation were collected into 0.5-L airtight glass vessels and stored at 4°C, while CO₂ gas was collected by the aforementioned gas-strip method for ¹⁴C measurement. In the case of natural samples, the DIC collection was carried out in the Tono Geoscience Center of JAEA and ¹⁴C was measured by AMS in JAEA (JAEA-AMS-TONO).

RESULTS AND DISCUSSION

Preliminary Comparison of ¹⁴C Values from Both Preparation Methods Using NaHCO₃ Solutions

¹⁴C and ¹³C values of the NaHCO₃ powder used for Solution 1 were measured twice and the results were 0.5 pMC, -18.6‰ and 0.4 pMC, -17.1‰, respectively. The ¹⁴C and ¹³C values of NaHCO₃ powder prepared for Solutions 2–4 were also measured twice and found to be 26.6 pMC, -18.9‰ and 26.1 pMC, -18.1‰, respectively. NaHCO₃ solutions were prepared by dissolving NaHCO₃ powders into deionized water or SSW. Therefore, the ¹⁴C in the DIC sample solutions can be calculated by Equation 1:

$$^{14}\text{C}_s = \frac{^{14}\text{C}_p \times [\text{C}]_p + ^{14}\text{C}_w \times [\text{C}]_w}{[\text{C}]_p + [\text{C}]_w} \quad (1)$$

where ¹⁴C_w is the ¹⁴C value (pMC) of DIC in the deionized water or SSW used for preparation of the NaHCO₃ solutions; ¹⁴C_p is the ¹⁴C value of NaHCO₃ powder; [C]_p and [C]_w are the concentrations of DIC (mg/L) provided from NaHCO₃ powder and DIC in the deionized water, respectively; and [C]_p is found from the difference between the DIC concentration of the sample solutions and [C]_w.

¹⁴C_w is considered to be the ¹⁴C of CO₂ in air. According to previous studies (Taylor 2004; Levin et al. 2013), the value of ¹⁴C_w is between 100 and 110 pMC. The [C]_w and DIC concentration of sample solutions before the experiment (i.e. sum of [C]_p and [C]_w) were determined using a TOC analyzer and the measurements included a maximum error of 5%. The ¹⁴C concentrations of the powder NaHCO₃ reagent were 0.4 and 0.5 pMC; thus, the correct value of ¹⁴C_w was considered to be between 0 and 1 pMC. Similarly, the correct values of ¹⁴C_w for prepared NaHCO₃ Solutions 2–4 were assumed to be between 26 and 27 pMC. These values were put into Equation 1 and a possible range of ¹⁴C_s (correct value of ¹⁴C in the sample solution) was calculated. The possible ranges of ¹⁴C_s for Solutions 1–4 were 3.1 to 5.2, 26.0 to 28.0, 26.0 to 29.0, and 26.0 to 28.0 pMC, respectively. The calculated ranges of ¹⁴C_s are compared to the measured ¹⁴C values in Table 2.

Table 2 shows the ¹⁴C and ¹³C values of DIC measured by the gas strip (_{gas strip}) and precipitation (_{precip}) methods for NaHCO₃ solutions. In Solutions 1 and 2, the differences between ¹⁴C_{gas strip} and ¹⁴C_{precip} ranged from 4 to 7 pMC. However, the ages estimated by the gas-strip and precipitation methods differed significantly (by 7000 to 12,000 yr) for Solution 1, and by 1000 to 1200 yr for Solution 2. These results indicate that the uncertainty of ¹⁴C age becomes larger when the age of old groundwater is estimated, as expected. The difference between ¹⁴C_{gas strip} and ¹⁴C_{precip} in Solution 3 was about 10 pMC, which was twice that of Solution 1 and 2. This result indicates the difference becomes more significant in the case of solution with low DIC. Furthermore, the difference in Solution 4 was 17 to 33 pMC, indicating a very large difference compared to other solutions. This result shows that the ionic strength and/or ion composition of the sample water can significantly affect ¹⁴C values of the

Table 2 Comparison of measured ^{14}C and possible ranges of ^{14}C for NaHCO_3 solutions.

Sample name	Extraction method	$\delta^{13}\text{C}$ (‰)	^{14}C (pMC)	^{14}C yr BP	Possible range of $^{14}\text{C}_s^*$ (pMC)	Deviation from $^{14}\text{C}_s$ (pMC)
Sample-1-1-1	Gas-strip	-20.0	2.0 ± 0.1	$31,430 \pm 180$	3.1 to 5.2	-1.1
Sample-1-2-1		-19.3	3.7 ± 0.1	$26,320 \pm 110$		—
Sample-1-1-2	Precipitation	-19.6	9.0 ± 0.1	$19,290 \pm 70$		+3.8
Sample-1-2-2		-20.2	11.4 ± 0.1	$17,340 \pm 60$		+6.2
Sample-2-1-1	Gas-strip	-19.9	29.0 ± 0.1	9947 ± 32	26.0 to 28.0	+1.0
Sample-2-2-1		-17.6	28.7 ± 0.1	$10,025 \pm 32$		+0.7
Sample-2-1-2	Precipitation	-19.0	33.3 ± 0.1	8842 ± 33		+5.3
Sample-2-2-2		-17.1	32.8 ± 0.1	8956 ± 34		+4.8
Sample-3-1-1	Gas-strip	-19.3	30.5 ± 0.1	9534 ± 37	26.0 to 29.0	+1.5
Sample-3-2-1		-19.3	30.3 ± 0.1	9591 ± 37		+1.2
Sample-3-1-2	Precipitation	-19.9	39.5 ± 0.2	7459 ± 32		+10.5
Sample-3-2-2		-19.7	39.3 ± 0.2	7502 ± 32		+10.3
Sample-4-1-1	Gas-strip	-17.8	29.2 ± 0.1	9890 ± 38	26.0 to 28.0	+1.2
Sample-4-2-1		-17.4	29.9 ± 0.1	9706 ± 38		+1.9
Sample-4-1-2	Precipitation	-23.7	62.6 ± 0.2	3770 ± 22		+34.6
Sample-4-2-2		-27.3	47.4 ± 0.2	5994 ± 27		+19.4

*Calculated by Equation 1.

precipitation method. There is also a possibility that the ^{14}C values might be controlled by relative relationships between carbonate and the other ion concentrations.

As shown in Table 2, the $^{14}\text{C}_{\text{gas strip}}$ values were close to the possible range of $^{14}\text{C}_s$ values calculated from Equation 1, although measured $^{14}\text{C}_{\text{gas strip}}$ values were always slightly (~ 1 pMC) higher. Conversely, $^{14}\text{C}_{\text{precip}}$ always showed much higher values relative to the calculated range of $^{14}\text{C}_s$ values. Thus, the gas-strip method is shown to be reliable for many types of groundwater samples with minimal error. $^{14}\text{C}_{\text{precip}}$ always showed significantly higher values than $^{14}\text{C}_s$. Therefore, the precipitation method should be applied carefully to groundwater samples and potential errors in $^{14}\text{C}_{\text{precip}}$ results should be explicitly assessed.

The effects of DIC concentration were investigated by comparing the ^{14}C values in Solutions 1, 2, and 3. The averaged differences between $^{14}\text{C}_{\text{precip}}$ and $^{14}\text{C}_s$ in Solutions 1, 2, and 3 were 5.0, 5.1, and 10.4 pMC, respectively. These results clearly indicate that the relative difference in ^{14}C estimated by the two methods increases with decreasing DIC concentration. These findings can be explained if we assume that a certain amount of modern carbon contaminated the sample during precipitation. If 0.6 to 0.7 mg of modern C contaminates 1 L of sample solution during the precipitation procedure, the observed differences between $^{14}\text{C}_{\text{precip}}$ and $^{14}\text{C}_s$ will occur.

A comparison of Solutions 2 and 4 shows the effect of salinity on $^{14}\text{C}_{\text{precip}}$. The difference between $^{14}\text{C}_{\text{gas strip}}$ and $^{14}\text{C}_{\text{precip}}$ was greater in Solution 4 than in Solution 2. If contamination by modern carbon from the air was the dominant cause of this phenomenon, we would expect that the ^{13}C values would be close to -8‰ (Taylor 2004). However, the $^{13}\text{C}_{\text{precip}}$ values in the Solution 4 test samples were -23.7 and -27.3‰ , significantly lower than the ^{13}C value of

NaHCO₃ powder (−18.5‰). Thus, the difference between ¹⁴C_{gas strip} and ¹⁴C_{precip} in saline water cannot be explained by contamination of modern carbon from the air alone.

Identification of Contamination in Samples Prepared by the Precipitation Method

The preliminary comparison experiments described earlier indicate that contamination by modern carbon might occur during the precipitation procedure. Modern carbon contamination from air to alkaline and SrCl₂ solutions was assumed to be the most likely source of contamination during the precipitation procedure. The background DIC concentrations in NaOH and SrCl₂ solutions were analyzed using a gas-strip line to confirm this possibility. Table 3 shows the DIC concentrations of each solution. Degassed deionized water with H₃PO₄ contained 0.01 mg/L DIC. The SrCl₂ solution is estimated to contain 0.04 mg/L DIC. Similarly, the DIC content of the NaOH solution (1 L of degassed deionized water with 5 mL of 5N NaOH solution) was estimated to range from 0.2 to 0.3 mg/L, when 5N NaOH solutions were prepared under atmosphere conditions. On the other hand, NaOH solutions prepared under inert conditions contained approximately 0.2 mg/L DIC, which was the same as prepared in atmosphere condition. Accordingly, the modern carbon contamination of the solutions appears to originate from the NaOH granules. Such background contamination of modern carbon probably also occurs for other alkaline solutions, such as ammonium solutions. Contamination with modern carbon in preliminary experiments using NaHCO₃ solutions could also originate from the NaOH solution used in the precipitation method.

Conversely, modern carbon contamination might occur in the gas-strip method during the addition of H₃PO₄ into the groundwater prior to carrier gas circulation because deionized water with H₃PO₄ contained minor but measurable DIC. This would be expected to influence the ¹⁴C_{gas strip} result from the NaHCO₃ solutions, which deviate by about 1 pMC from ¹⁴C_s (Table 2).

Application of the Two Methods to Natural Groundwater Sample

¹⁴C values of DIC in groundwater samples prepared by both the gas-strip and precipitation methods are shown in Table 4. ¹³C values of DIC are also listed in Table 4. ¹⁴C_{precip} values were approximately 0.4–9.5 pMC higher than those obtained using the gas-strip method. By assuming that ¹⁴C_{gas strip} is less contaminated by modern carbon and the ¹⁴C value of

Table 3 Background DIC concentration of solutions used for precipitation method.

Sample name	Preparation of solution	Composition of solution			DIC concentration in solution (mg/L)
		5N NaOH (mL/1 L water)	2N SrCl ₂ (mL/1 L water)	H ₃ PO ₄ (mL/1 L water)	
Pure water	In the atmosphere	0.0	0.0	2.0	0.01
SrCl ₂ solution		0.0	5.0	2.0	0.04
NaOH solution1		5.0	0.0	2.0	0.21
NaOH solution2		5.0	0.0	4.0	0.29
NaOH solution3		10.0	0.0	8.0	0.45
NaOH solution4	In inert atmosphere	5.0	0.0	4.0	0.18

Table 4 Comparison of ^{13}C and ^{14}C of groundwater samples prepared by gas-strip and precipitation methods.

Borehole-zone No. (date)	$\delta^{13}\text{C}_{\text{precip}}^*$ (‰)	$\delta^{13}\text{C}_{\text{gas strip}}^*$ (‰)	$^{14}\text{C}_{\text{precip}}^*$ (pMC)	$^{14}\text{C}_{\text{gas strip}}^*$ (pMC)	$(^{14}\text{C}_{\text{precip}}) -$ $(^{14}\text{C}_{\text{gas strip}})$ (pMC)	Concentration of DIC in sample (mg/L)	DIC _{contami} **(mg)
07MI07-1 (2012/10)	-17.35	-16.83	34.57	29.07	5.50	17.6	1.48
07MI07-3 (2012/10)	-16.13	-17.43	32.16	24.23	7.93	12.1	1.41
07MI07-5 (2012/10)	-15.61	-14.91	26.45	16.95	9.50	8.9	1.15
09MI20-1[1] (2012/10)	-14.98	-14.88	29.27	24.34	4.93	15.4	1.07
09MI20-1[2] (2012/10)	-14.28	-14.47	27.46	23.38	4.08	15.4	0.87
09MI20-1[3] (2012/10)	-14.34	-14.77	27.24	22.05	5.19	15.4	1.10
09MI20-1 (2014/9)	-14.02	-13.88	23.97	18.14	5.83	14.3	1.10
09MI20-3 (2012/10)	-13.38	-14.04	24.97	20.12	4.85	15.4	1.00
09MI20-3 (2014/9)	-13.19	-10.89	24.12	10.43	13.69	13.2	2.38
09MI20-5 (2012/10)	-11.73	-11.52	21.20	15.50	5.70	15.4	1.11
09MI20-5 (2014/9)	-12.31	-11.95	19.82	12.02	7.80	10.3	1.01
10MI26-1 (2012/10)	-7.39	-6.79	10.25	9.86	0.39	15.4	0.07
10MI26-3 (2012/10)	-8.89	-8.09	15.33	7.90	7.43	15.4	1.35
10MI26-5 (2012/10)	-13.47	-12.77	27.18	19.09	8.09	15.4	1.71
12MI33-4 (2015/2)	No precipitation	-11.43	No precipitation	2.10	—	4.9	—
						Average	1.04

* $\delta^{13}\text{C}_{\text{precip}}$, $\delta^{13}\text{C}_{\text{gas strip}}$, $^{14}\text{C}_{\text{precip}}$, and $^{14}\text{C}_{\text{gas strip}}$ are measured ^{13}C or ^{14}C of samples prepared by precipitation and gas-strip method, respectively.

**DIC_{contami} is the amount of DIC contamination of modern carbon calculated by Equation 2.

contaminated modern carbon in samples obtained by the precipitation method is 100 pMC, the mass balance equation describing contamination in the precipitation methods ($^{14}\text{C}_{\text{precip}}$) can be expressed by the following equation:

$$\text{DIC}_{\text{gw}} \times ^{14}\text{C}_{\text{gas strip}} + \text{DIC}_{\text{contami}} \times 100 = (\text{DIC}_{\text{gw}} + \text{DIC}_{\text{contami}}) \times ^{14}\text{C}_{\text{precip}} \quad (2)$$

where $\text{DIC}_{\text{contami}}$ is the amount of DIC contamination with modern carbon (mg) and DIC_{gw} is the amount of DIC (mg) in groundwater samples analyzed before addition of NaOH solution. The amounts of modern carbon contamination during the precipitation procedure (described as $\text{DIC}_{\text{contami}}$ above) were estimated to be 1 mg for all groundwater samples, regardless of the $^{14}\text{C}_{\text{gas strip}}$ and DIC_{gw} values (Table 4). These findings suggest that modern carbon contamination during preparation of the precipitation method significantly influenced the ^{14}C value according to the DIC content and ^{14}C concentration of sampled water.

The $^{14}\text{C}_{\text{gas strip}}$ value of groundwater at depths of 200 to 500 m were estimated to range from 2 to 29 pMC, with a tendency to become lower with increasing depth. We were not able to collect the DIC in groundwater at 500 m depth (12MI33 borehole) by the precipitation method. The reason that the DIC did not precipitate is not clear, but is likely related to the low DIC concentration (4.9 mg/L) and high salinity (about 450 mg/L) of the sample. The groundwater residence time at that depth is estimated to be about 31,000 yr by the $^{14}\text{C}_{\text{gas strip}}$ value. However, the groundwater flow around this large underground facility (MIU) has been influenced by water drainage for more than 10 yr (Iwatsuki et al. 2015), and the analyzed ^{14}C values might reflect mixed shallow and deep groundwater sources.

When we discuss the groundwater ages with ^{14}C , the ^{14}C values should be corrected considering the geochemical reactions that could affect to the ^{14}C values (Clark and Fritz 1997; Kalin 2000). However, in this article groundwater ages are estimated without correction of geochemical reactions because our main purpose is to indicate the difference between the precipitation and gas-strip methods.

Applicability of the Precipitation Method for Groundwater Dating

As discussed earlier, contamination with modern carbon probably occurs during the DIC precipitation procedure. The effects of contamination varied significantly depending on the DIC and ^{14}C concentrations of the groundwater samples. Equation 2 was applied to estimate the error in cases in which modern carbon contaminated groundwater with various DIC concentrations. For example, Figures 3a and 3b show ^{14}C values after contamination of 1 L of groundwater with 0.5 or 1 mg modern carbon as estimated by Equation 2. The error from the true ^{14}C value became more significant as the DIC and ^{14}C concentrations of sampled groundwater decreased. Constraints for DIC and ^{14}C of groundwater that should be analyzed within 10% error are estimated in Figure 3c. Because background contamination with modern carbon was estimated for each preparation method, the uncertainty in ^{14}C analysis could be inferred by the DIC concentration and probable ^{14}C content of the groundwater. As described, 1 L of NaHCO_3 solutions and groundwater samples were contaminated with about 0.7 mg and 1 mg of modern carbon, respectively.

As shown in Figure 3b, if we want to control the ^{14}C age error from the precipitation procedure within 10%, the concentration of DIC and $^{14}\text{C}_{\text{gw}}$ contents should be higher than 20, 10, and 1 mg/L and 32, 47, and 83 pMC, respectively. If the error has to be controlled within 20%, the concentration of DIC and $^{14}\text{C}_{\text{gw}}$ should be higher than 20, 10, and 1 mg/L and 19, 31, and 71 pMC, respectively.

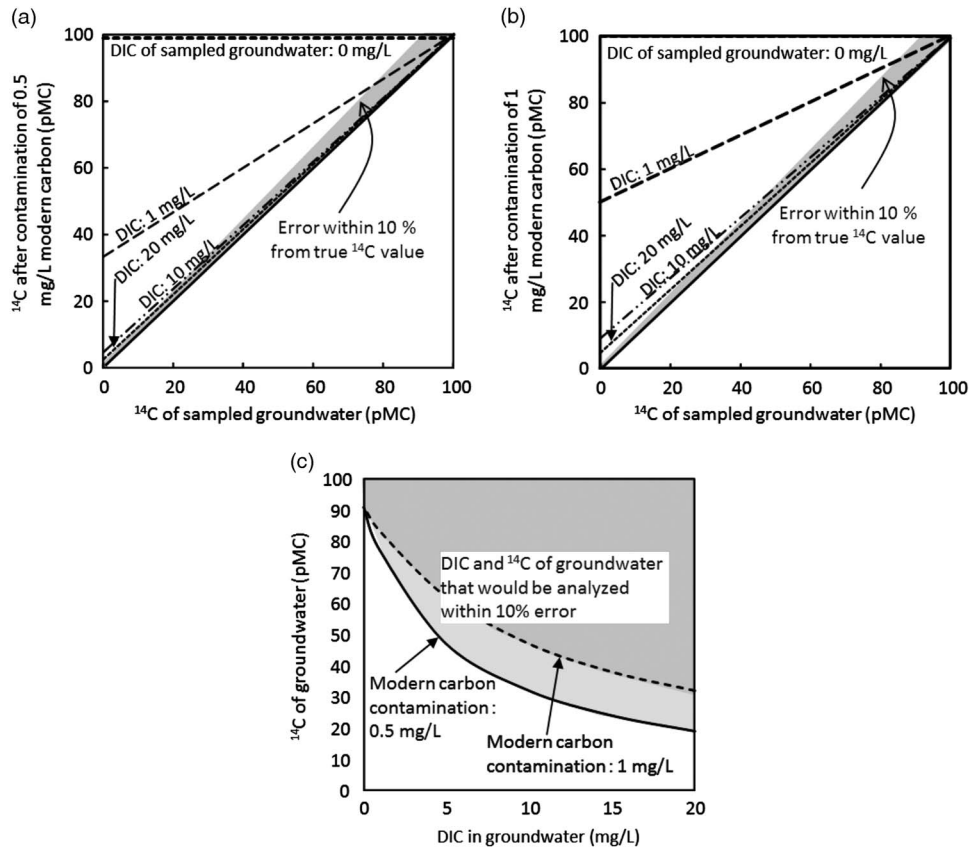


Figure 3 Error in ^{14}C analysis for contamination by modern carbon for groundwater with initial DIC concentrations of 0–20 mg/L. ^{14}C value of contaminated modern carbon is assumed to be 100 pMC. (a) Contamination by modern carbon is 0.5 mg/L. (b) Contamination by modern carbon is 1 mg/L. The shaded area shows error within 10% of the known ^{14}C value. (c) Constraint for DIC and ^{14}C of groundwater that should be analyzed within 10% error in the case of modern carbon contamination with 0.5 and 1 mg/L.

Because our purpose was to determine the effects of both methods on estimation of groundwater age, the relationships between expected groundwater age estimated by ^{14}C (without considering the effect of geochemical reaction) using the precipitation method and DIC concentrations are compared in Figure 4. The expected groundwater age estimated by ^{14}C using the precipitation method was calculated by the following equations:

$$^{14}\text{C age}_{\text{ex}}(y) = \frac{5730}{\ln 2} \times \ln \left(\frac{^{14}\text{C}_{\text{precip}}}{100} \right)^{(-1)x} \tag{3}$$

$$^{14}\text{C}_{\text{precip}} = \frac{^{14}\text{C}_{\text{gw}} \times [\text{DIC}]_{\text{gw}} + ^{14}\text{C}_{\text{contami}} \times 1.0}{[\text{DIC}]_{\text{gw}} + 1.0} \tag{4}$$

Equation 3 is frequently used for estimation of groundwater age from ^{14}C (Clark and Fritz 1997; Mook and Plicht 1999). $^{14}\text{C age}_{\text{ex}}$ is the expected ^{14}C age after precipitation and $^{14}\text{C}_{\text{precip}}$ is the ^{14}C concentration after precipitation with 1 mg modern carbon contamination.

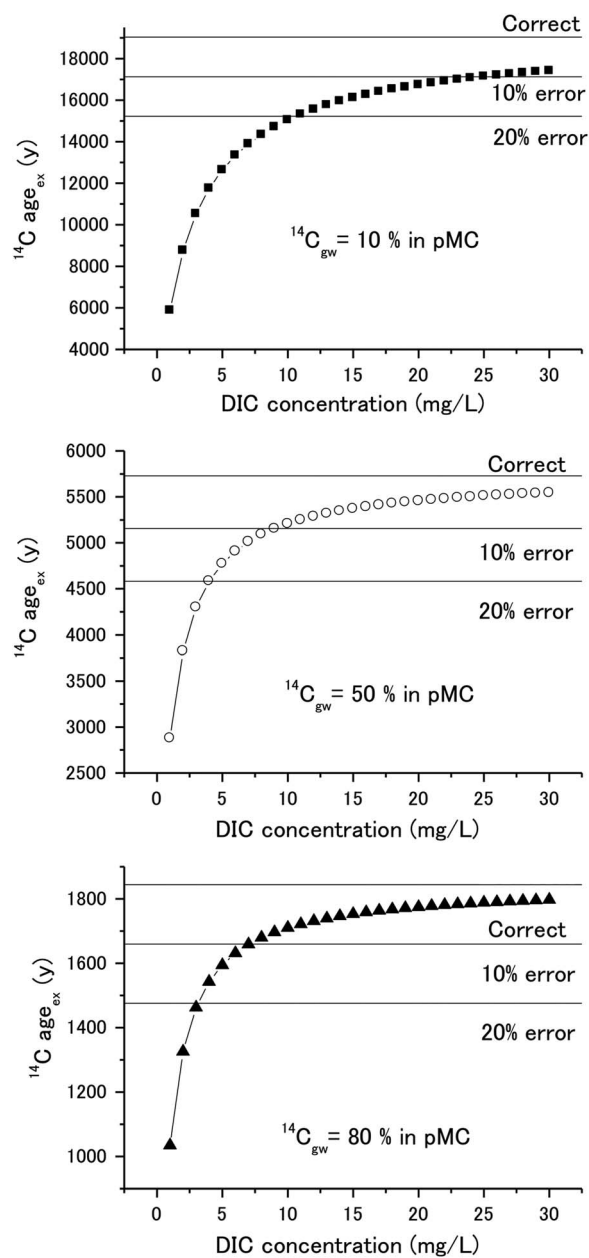


Figure 4 Relationship between concentration of DIC in groundwater and expected ^{14}C ages (^{14}C age_{ex}) for samples prepared by the precipitation method where groundwater ^{14}C ($^{14}\text{C}_{\text{gw}}$) is 10, 50, and 80 pMC. ^{14}C age_{ex} values were calculated using Equations 3 and 4.

$^{14}\text{C}_{\text{gw}}$ and $^{14}\text{C}_{\text{contami}}$ are the concentrations of DI^{14}C in the groundwater samples and modern carbon, respectively. The value of $^{14}\text{C}_{\text{contami}}$ is set to 100 pMC. $[\text{DIC}]_{\text{gw}}$ is the concentration of DIC in the groundwater samples.

The acceptable error depends on the purpose of the study. When we applied the precipitation method to estimate groundwater age, the acceptable error for the study and conditions that affect the ^{14}C age (DIC and ^{14}C concentrations in targeted groundwater) had to be considered. The information provided in Figures 3 and 4 is useful to determine if the precipitation method is applicable for a targeted groundwater. According to the results of ^{14}C values for NaHCO_3 solutions and groundwater samples treated using the precipitation method, we could assume that the amount of modern carbon contamination to be about 1 mg. However, this value might depend on the atmosphere of the laboratory and reagents used for the precipitation procedure. Thus, quantitative estimation of contamination by modern carbon in each laboratory during the precipitation procedure should be evaluated if the precipitation method is chosen.

CONCLUSION

In this study, the differences in ^{14}C values of samples prepared by the gas-strip and precipitation methods were compared in artificially prepared and natural groundwater samples. The ^{14}C values obtained by the gas-strip method were very similar to the theoretically predicted ^{14}C values in chemically distinct waters, indicating that this method can be applied to a variety of groundwaters. Conversely, the ^{14}C value obtained by the precipitation method showed higher than predicted ^{14}C values. In the case of natural groundwater samples, ^{14}C obtained by the precipitation method always showed higher values than those from the gas-strip method. It was assumed that an average of 1 mg/L of modern carbon contaminated the samples during the precipitation procedure. This contamination affects the estimation of groundwater ages, especially when DIC concentrations and ^{14}C contents are low. Thus, the applicability of the precipitation method should be considered according to the required accuracy and/or purpose of the study. The amount of contamination with modern carbon might differ among laboratories and should therefore be estimated in each laboratory to evaluate the applicability of the precipitation method.

For solutions with high salinity, the differences between ^{14}C measured from precipitation samples and predicted ^{14}C values increased. Thus, for groundwater with low DIC concentration and/or high salinity and old groundwater, the gas-strip method should be applied to estimate the groundwater age by DI^{14}C .

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REFERENCES

- Aggarwal PK, Araguas-Araguas L, Choudhry M, van Duren M, Froehlich K. 2014. Lower groundwater ^{14}C age by atmospheric CO_2 uptake during sampling and analysis. *Groundwater* 52(1):20–4.
- Atekwana EA, Krishnamurthy RV. 1998. Seasonal variations of dissolved inorganic carbon and $\delta^{13}\text{C}$ of surface waters: application of a modified gas evolution technique. *Journal of Hydrology* 205(3–4):265–78.
- Clark I, Fritz P. 1997. Age and dating old groundwater. Chapter 8. In: *Environmental Isotopes in Hydrogeology*. New York: Lewis Publishers. p 197–243.
- Corcho Alvarado JA, Paces T, Purtschert R. 2013. Dating groundwater in the Bohemian Cretaceous Basin: understanding tracer variations in the subsurface. *Applied Geochemistry* 29:189–98.
- Dorsett A, Cherrier J, Martin JB, Cable JE. 2011. Assessing hydrologic and biogeochemical controls

- on pore-water dissolved inorganic carbon cycling in a subterranean estuary: a ¹⁴C and ¹³C mass balance approach. *Marine Chemistry* 127(1–4): 76–89.
- Dulinski M, Rozanski K, Kuc T, Gorczyca Z, Kania J, Kapusta M. 2013. Evolution of radiocarbon in a sandy aquifer across large temporal and spatial scales: case study from southern Poland. *Radiocarbon* 55(2–3):905–19.
- Geyh MA. 2000. An overview of ¹⁴C analysis in the study of groundwater. *Radiocarbon* 42(1):99–114.
- Iwatsuki T, Xu S, Itoh S, Abe M, Watanabe M. 2000. Estimation of relative groundwater age in the granite at the Tono research site, central Japan. *Nuclear Instruments and Methods in Physics Research B* 172(1–4):524–9.
- Iwatsuki T, Furue R, Mie H, Ioka S, Mizuno T. 2005. Hydrochemical baseline condition of groundwater at the Mizunami underground research laboratory (MIU). *Applied Geochemistry* 20(12): 2283–302.
- Iwatsuki T, Hagiwara H, Ohmori K, Munemoto T, Onoe H. 2015. Hydrochemical disturbances measured in groundwater during the construction and operation of a large-scale underground facility in deep crystalline rock in Japan. *Environmental Earth Science* 74(4):3041–57.
- Kalin RM. 2000. Radiocarbon Dating of Groundwater Systems. Chapter 4. In: Cook PG, Herczeg AL, editors. *Environmental Tracers in Subsurface Hydrology*. Boston: Kluwer Academic Publishers. p 111–44.
- Levin I, Kromer B, Hammer S. 2013. Atmospheric ¹⁴CO₂ trend in western European background air from 2000 to 2012. *Tellus B* 65:20092.
- Minami M, Takahashi HA. 2015. Radiocarbon analysis of water samples by the precipitation method. *Proceedings of the Annual Meeting of the Geochemical Society Japan*. p 329. In Japanese.
- Mook WG, van der Plicht J. 1999. Reporting ¹⁴C activities and concentrations. *Radiocarbon* 41(3): 227–39.
- Stewart MK. 2012. A 40-year record of carbon-14 and tritium in the Christchurch groundwater system, New Zealand: dating of young samples with carbon-14. *Journal of Hydrology* 430–431:50–68.
- Sültenfuß J, Purtschert R, Führböter JF. 2011. Age structure and recharge conditions of a coastal aquifer (northern Germany) investigated with ³⁹Ar, ¹⁴C, ³H, He isotopes and Ne. *Hydrogeology Journal* 19(1):221–36.
- Taylor CB. 2004. Time-dependent factors inherent in the age equation for determining residence times of groundwater using ¹⁴C: a procedure to compensate for the past variability of ¹⁴C in atmospheric carbon dioxide, with application to the Wairau deep aquifer, Marlborough, New Zealand. *Radiocarbon* 46(2):501–15.