# Leószilárdite, the first Na,Mg-containing uranyl carbonate from the Markey Mine, San Juan County, Utah, USA

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# ABSTRACT

Leószilárdite (IMA2015-128), Na<sub>6</sub>Mg(UO<sub>2</sub>)<sub>2</sub>(CO<sub>3</sub>)<sub>6</sub>·6H<sub>2</sub>O, was found in the Markey Mine, Red Canyon, White Canvon District, San Juan County, Utah, USA, in areas with abundant andersonite, natrozippeite, gypsum, anhydrite, and probable hydromagnesite along with other secondary uranium minerals bayleyite, čejkaite and johannite. The new mineral occurs as aggregates of pale vellow bladed crystals flattened on {001} and elongated along [010], individually reaching up to 0.2 mm long. More commonly it occurs as pale yellow pearlescent masses to 2 mm consisting of very small plates. Leószilárdite fluoresces green under both longwave and shortwave ultraviolet light, and is translucent with a white streak, hardness of 2 (Mohs), and brittle tenacity with uneven fracture. The new mineral is readily soluble in room temperature  $H_2O$ . Crystals have perfect cleavage along  $\{001\}$ , and exhibit the forms  $\{110\}$ ,  $\{001\}$ ,  $\{100\}$ ,  $\{101\}$  and  $\{\overline{1}01\}$ . Optically, leószilárdite is biaxial (-),  $\alpha = 1.504(1)$ ,  $\beta = 1.597(1)$ ,  $\gamma = 1.628(1)$  (white light); 2V (meas.) = 57(1)°, 2V  $(calc.) = 57.1^{\circ}$ ; dispersion r > v, slight. Pleochroism: X = colourless, Y and Z = light yellow;  $X < Y \approx Z$ . The average of six wavelength dispersive spectroscopic analyses provided Na<sub>2</sub>O 14.54, MgO 3.05, UO<sub>3</sub> 47.95, CO<sub>2</sub> 22.13, H<sub>2</sub>O 9.51, total 97.18 wt.%. The empirical formula is Na<sub>5.60</sub>Mg<sub>0.90</sub>U<sub>2</sub>O<sub>28</sub>C<sub>6</sub>H<sub>12.60</sub>, based on 28 O apfu. Leószilárdite is monoclinic, C2/m, a = 11.6093(21), b = 6.7843(13), c = 15.1058(28) Å,  $\beta =$ 91.378(3)°, V = 1189.4(4) Å<sup>3</sup> and Z = 2. The crystal structure ( $R_1 = 0.0387$  for 1394 reflections with  $I_{obs} >$  $4\sigma I$ , consists of uranyl tricarbonate anion clusters  $[(UO_2)(CO_3)_3]^4$  held together in part by irregular chains of  $NaO_{5}(H_{2}O)$  polyhedra sub parallel to [010]. Individual uranyl tricarbonate clusters are also linked together by three-octahedron units consisting of two Na-centred octahedra that share the opposite faces of a Mg-centred octahedron at the centre (Na–Mg–Na), and have the composition  $Na_2MgO_{12}(H_2O)_4$ . The name of the new mineral honours the Hungarian-American physicist, inventor and biologist Dr. Leó Szilárd (1898-1964).

Keywords: leószilárdite, new mineral, uranium, uranyl carbonate, crystal structure, Markey mine.

#### Introduction

\*E-mail: tolds@nd.edu https://doi.org/10.1180/minmag.2016.080.149 TYPICALLY, uranyl carbonates are soluble ephemeral solids and are the dominant phases that precipitate

from neutral to alkaline waters containing uranium (Finch and Ewing, 1992; Finch and Murakami, 1999). Uranyl carbonate complexes are exceptionally stable and dictate the mobility of uranium in carbonated groundwater, where the uranyl ion  $(UO_2)^{2+}$  is readily solubilized (Langmuir 1978; Clark et al., 1995; Gorman-Lewis et al., 2008). With the acceptance of leószilárdite, the family of uranyl carbonate minerals now totals 34. The majority of these contain the uranyl tricarbonate complex (henceforth referred to as UTC)  $[(UO_2)]$  $(CO_2)_2$ <sup>4-</sup>, which is commonly present in solution coordinated with other cations  $(K^+, Na^+, Ca^{2+}, etc.)$ and the phase that ultimately precipitates is influenced by the local groundwater chemistry, pH, pCO<sub>2</sub> and evaporation rates (Garrels and Christ, 1959; Hostetler and Garrels, 1962). Leószilárdite differs chemically from all others as the first Na- and Mg-containing uranyl carbonate a somewhat surprising occurrence considering the cations involved are commonly found together.

Leószilárdite is named in honour of the Hungarian-American physicist, inventor and biologist Dr. Leó Szilárd (1898-1964). Szilárd held many patents for his inventions, including those for the linear accelerator, the cyclotron, the Szilárd-Einstein refrigerator and the nuclear reactor with Enrico Fermi in 1933 - however their reactor was not based upon fission as the phenomenon had not vet been discovered. He is perhaps most wellknown for envisioning the concept of, and patenting, the neutron-induced nuclear chain reaction. Likewise, he played a compelling role in the development of the United States nuclear program, by penning a letter to Franklin D. Roosevelt warning that Germany may develop nuclear weapons that harnessed the fission reaction of uranium. The famous letter, signed by Albert Einstein, jumpstarted the research and development of the Manhattan project.

The new mineral and its name were approved by the Commission on New Minerals, Nomenclature and Classification of the International Mineralogical Association (IMA2015-128). The holotype specimen is deposited in the collections of the Natural History Museum of Los Angeles County, 900 Exposition Boulevard, Los Angeles, CA 90007, USA, catalogue number 65645.

# Occurrence

Leószilárdite was found underground at the Markey mine, Red Canyon, White Canyon District, San

Juan County, Utah, USA (UTM coordinates; 12 S 4155830 mN, 0561730 mE, NAD27). The workings lie 1 km southwest of the Blue Lizard mine, on the east-facing side of the canyon roughly 73 km west of Blanding, Utah and 35 km south of Hite on Lake Powell. Leószilárdite was later found occurring with abundant andersonite at the Pickett Corral mine, Montrose County, Colorado, USA, but this is not considered a type locality; the description is based only on material from the Markey mine. The geology of the Markey mine is quite similar to that of the nearby Blue Lizard mine (Chenoweth, 1993) although the mineralogy of the Markey mine is notably richer in carbonate phases. As an interesting aside, this difference in mineral chemistry may be reflected by the current mine air chemistry. Preliminary underground gas measurements collected in 2016 using a hand-held Crowcon Gasman CO<sub>2</sub> monitor showed consistently elevated CO<sub>2</sub> levels at the Markey mine, averaging  $\sim 1000 \text{ ppm CO}_2$  with a maximum recorded value of 1600 ppm CO<sub>2</sub>, levels considerably higher than at the nearby Blue Lizard (and Posey) mines where carbonate mineral species are less abundant.

This geological and historical information following is taken from a report by Strand (1954) contained within an application to the Defense Minerals Exploration Administration (docket 4260) for government aid to continue uranium exploration of the area surrounding the Markey mine. All large uranium deposits in the area of Red Canyon are characterized as channel-type sediments from streams that deposited the Triassic Shinarump conglomerate, and commonly cut into the underlying reddish-brown silty sandstones and mudstones of the Moenkopi Formation. Today, the characteristic red rocks of the Moenkopi Formation form much of the lower floor and walls of the canyon, assumed to be the origin of the name Red Canyon. Location and exploration of the claims encompassing the Markey mine area began in 1951 by Edward H. Eakland, Jr. and others, and the first batch of copper-uranium ore was mined and shipped to the Atomic Energy Commission mill in Monticello. Utah by the Anaconda Copper Mining Company in 1953. In total 1,214 tons of uranium ore were removed averaging 20% U<sub>3</sub>O<sub>8</sub>. The mine now lies abandoned and post-mining oxidation of primary minerals in the humid underground environment is responsible for producing interesting secondary uranium minerals, including abundant andersonite and natrozippeite. Leószilárdite is a relatively rare mineral in the



FIG. 1. Bladed pale yellow leószilárdite on andersonite. Horizontal field of view is ~1.2 mm. Travis Olds specimen and photo.

secondary carbonate assemblage at the Markey mine, and is found in areas with abundant andersonite and gypsum, along with the other secondary minerals bayleyite, čejkaite, johannite, natrozippeite and chalcanthite coating veins of uraninite in sulfide-laden carbonaceous wood fragments.

# Physical and optical properties

Leószilárdite forms as aggregates of long thin pale yellow blades individually reaching 0.2 mm long, and pale yellow pearlescent masses to 2 mm composed of thin plates (Figs 1, 2 and 3). Crystals are flattened on {001} and elongated on [010], and exhibit the forms {110}, {001}, {100}, {101} and { $\overline{101}$ } (Fig. 4). No twinning was



FIG. 2. Pearlescent masses of leószilárdite coat andersonite, with yellow natrozippeite and probable white spherical hydromagnesite. Horizontal field of view is  $\sim$ 1.9 mm. Travis Olds specimen and photo.



FIG. 3. Secondary electron images of leószilárdite microcrystals. Images by Owen Mills.

observed. Crystals are transparent with a vitreous lustre, while crystalline masses exhibit a pearly lustre. The mineral has a white streak, is weakly fluorescent green under both longwave and shortwave ultraviolet light. The Mohs hardness is  $\sim 2$ , estimated by the behaviour of crystals when broken. Crystals of leószilárdite are brittle with perfect {001} cleavage and even fracture. The density could not be measured due to the limited availability of material, and solubility in aqueous



FIG. 4. Crystal drawing of leószilárdite; clinographic projection in non-standard orientation (*b* vertical).

heavy liquids. The calculated density is  $3.256 \text{ g cm}^{-3}$  based on the empirical formula.

Optically, leószilárdite is biaxial (–), with  $\alpha = 1.504(1)$ ,  $\beta = 1.597(1)$ ,  $\gamma = 1.628(1)$  (measured in white light). The measured 2V is 57(1)°, based on extinction data collected on a spindle stage and analysed using *EXCALIBRW*; the calculated 2V is 57.1°. Dispersion is slight, r > v. The mineral is pleochroic with X = colourless, Y and Z = light

yellow;  $X \le Y \approx Z$ . The optical orientation is X =**b**,  $Y \approx$ **a**,  $Z \approx$ **c**.

The Gladstone-Dale compatibility,  $1 - (K_p/K_c)$ , is -0.011 (superior) for the ideal formula, and -0.053 (good) for the empirical formula (Mandarino, 2007). The value for  $k(UO_3) = 0.118$ , given by Mandarino (1976) was used for the calculation as it yielded the best compatibilities.

#### Raman spectroscopy

The Raman spectrum of leószilárdite was recorded using a Bruker Instruments Sentinel-785 laser head mounted on a Nikon Optiphot-2 microscope with Peltier-cooled integrated 785 nm diode laser, operated at 200 mW, with a spot size of 100  $\mu$ m and resolution of ~5 cm<sup>-1</sup>. The spectrometer was calibrated with software-controlled procedures (Opus software) using neon emission lines (wavelength calibration) and Tylenol® Raman bands (frequency calibration). A background correction was applied using the Opus software.

The Raman spectrum of leószilárdite is given in Fig. 5, and the following assignments are based on those given by various authors (Čejka, 1999; Jolivet *et al.*, 1980; Urbanec and Čejka, 1979). The v<sub>3</sub> (CO<sub>3</sub>)<sup>2–</sup> antisymmetric stretching vibrations are observed at 1535, 1396 and 1328 cm<sup>-1</sup>. No bands related to the v<sub>2</sub> ( $\delta$ )-bending vibration of



FIG. 5. The Raman spectrum of leószilárdite collected with a 785 nm laser.

1042

Constituent	Mean wt.%	Range	Standard. deviation	Probe standard			
Na <sub>2</sub> O	14.54	14.15-15.06	0.31	Albite			
MgO	3.05	2.71-3.35	0.27	Olivine			
UÕ3	47.95	47.35-48.74	0.63	$UO_2$ (syn.)			
CO <sub>2</sub> *	22.13			2 ( ) /			
H <sub>2</sub> Ô**	9.51						
Total	97.18						

TABLE 1. Chemical	l data fo	or leószil	lárdite.
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\*Calculated based on the structure model for 6 C per formula unit. \*\*Calculated to provide charge balance.

water were observed. The split  $v_1 (CO_3)^{2-}$  symmetric stretching vibrations are observed at 1078, 1062 and 1052 cm<sup>-1</sup>. These three bands verify the

presence of three structurally non-equivalent carbonate units seen in the X-ray data. A very strong Raman band at  $824 \text{ cm}^{-1}$  is assigned as the v<sub>1</sub>

TABLE 2. Powder X-ray data for leószilárdite (d values in Å).

$I_{\rm obs}$	$d_{\rm obs}$	$d_{\text{calc}}$	Icalc	h k l	$I_{\rm obs}$	$d_{\rm obs}$	$d_{\rm calc}$	Icalc	h k l	$I_{\rm obs}$	$d_{\rm obs}$	$d_{\rm calc}$	Icalc	h k l
		15.1016	4	001			(2.5448	2	$\overline{2} 2 3$	11	1 0 1 7 0	∫1.8204	6	027
36	7.59	7.5508	45	002	15	2 524	2.5403	4	403	11	1.81/8	1.8128	2	333
12	5 02	∫ 5.8570	7	110	15	2.334	2.5366	7	314			(1.8031	2	118
15	5.82	l 5.8028	7	200			2.5233	2	024	14	1 7020	1.7985	2	317
		5.4832 )	25	111			2.5169 )	2	006	14	1./920	1.7958	4	515
100	5.46	{ 5.4608	15	$\overline{2} 0 1$	8	2.492	{ 2.4880	7	403			1.7847	3	135
		5.4385	60	111			2.4845	3	314			1.7655	2	425
		5.3736	4	201	8	2 3 1 7	{2.3277	2	404			(1.7471	2	227
28	4 64	∫ 4.6554	24	112	0	2.317	l 2.3227	6	116	9	1.7444	{1.7428	3	334
20	7.07	l 4.5485	9	<u>2</u> 02			2.2891	3	206			1.7385	2	604
		( 3.8486	6	<u>2</u> 03			2.2557	2	025	8	1.6953	1.6961	6	040
30	3.82	3.8407	11	113			(2.2234	4	<u>3</u> 15			1.6803)	2	620
00	0.01	3.7949	14	113			2.1975	4	131	9	1.6735	1.6780	2	009
		3.7754	4	004	22	2 1 9 5	2.1962	2	510		110700	1.6737	3	426
		3.7580	3	203	22	2.195	2.1946	5	131			1.6686	2	136
33	3.383	{ 3.3922	21	020			2.1804	3	511			1.6548	2	042
00	01000	l 3.3606	11	310			2.1761	4	421			1.6364	2	516
		3.2659	4	$\frac{3}{11}$	8	2.158	2.1574	6	$\frac{0}{100}$			1.6197	2	241
6	3.209	3.1910	7	114	10		2.1322	2	132			1.5892	2	242
25	3.105	3.1304	9	204	10	2.103	2.0960	6	$\frac{5}{2}$ 1 2	6	1.5771	1.5783	4	$\frac{5}{2}32$
		(3.0943	18	022			(2.0351	4	$\frac{316}{100}$			[1.5518	3	336
		3.0467	2	312			2.0333	8	$\frac{423}{512}$	11	1.5482	{1.5495	2	$\frac{5}{2}33$
		2.9285	3	$\frac{220}{220}$	28	2.0283	2.0299	4	513			(1.5472	2	624
21	0.04	2.8815	9	$\frac{2}{4} \frac{2}{1} \frac{1}{1}$			2.02/6	3	133	5	1.4882	1.4913	2	244
31	2.864	{ 2.8621	9	401			2.0165	5	11/			(1.4/96)	2	$\frac{5}{4}$
		(2.830)	9	401	5	1.0464	2.0062	2	423		1.4587	1.4591	2	441
		2.8131	2	023	3	1.9404	1.9525	2	$\frac{330}{601}$	10	1 4270	1 4275	3	$\frac{441}{525}$
		2.7084	3 7	213	9	1.9174	1.9244	2 5	$\frac{001}{124}$	10	1.4370	1.43/3	4	555
20	2 704	$\begin{bmatrix} 2.7192 \\ 2.7050 \end{bmatrix}$	/	$\frac{2}{2}$ 0 5			(1.91/3	2	$\frac{134}{602}$					
20	2.704	2.7039	10	205 115	12	1.8830	1.8798	2 7	5 1 4					

The strongest lines are given in bold.

the  $(UO_2)^{2+}$  symmetric stretching vibration. Bartlett and Cooney (1989) provide an empirical relation to derive the approximate U–O<sub>y1</sub> bond length from the band position assigned to the v<sub>1</sub>  $(UO_2)^{2+}$  stretching vibration, which gives 1.79 Å. This value is in accordance with the U–O<sub>y1</sub> bond lengths given by Burns *et al.* (1997) for the uranyl cation in hexagonal bipyramidal coordination, and the U–O<sub>y1</sub> bond length of 1.79 Å derived from the X-ray data.

Raman bands at 742, 729, 705 and 695 cm<sup>-1</sup> are attributed to the split doubly degenerate  $v_4$  ( $\delta$ ) (CO<sub>3</sub>)<sup>2-</sup> in-plane bending vibrations. A weak band at 345 cm<sup>-1</sup> is assigned as a v (U–O<sub>*ligand*</sub>) stretch, and the doubly degenerate  $v_2$  ( $\delta$ ) (UO<sub>2</sub>)<sup>2+</sup> bending mode is found at 290 and 254 cm<sup>-1</sup>, though band assignment in this region is difficult due to overlap with U–O<sub>*ligand*</sub> stretches. Remaining bands at 193,

TABLE 3. Data collection and structure-refinement details for leószilárdite.

Diffractometer	Bruker Quazar II with Apex II detector
X-ray radiation/power	MoKα ( $\lambda$ = 0.71075 Å)/50 kV, 60 mA
Temperature	298(2) K
Structural formula	$Na_{c}Mg(UO_{a})_{a}(CO_{a})_{c}$ · 6H <sub>a</sub> O
Space group	C2/m
Unit-cell dimensions	a = 11.6093(21) Å
	b = 6.7843(13) Å
	c = 15.1058(28) Å
	$\beta = 91.378(3)^{\circ}$
V	1189.4(4) Å <sup>3</sup>
Ζ	2
Density (for above formula)	$3.269 \text{ g cm}^{-3}$
Absorption coefficient	$13.862 \text{ mm}^{-1}$
F(000)	1044.0
Crystal size (µm)	$45 \times 30 \times 6$
θ range	1.35 to 29.26°
Index ranges	$-15 \le h \le 15, -8 \le k \le 8, -20 \le l \le 19$
Reflections collected/ unique	$7517/1634; R_{\rm int} = 0.0719$
Reflections with $I_{obs} > 4\sigma(I)$	1394
Completeness to $\theta = 25.24^{\circ}$	100%
Refinement method	Full-matrix least-squares on $F^2$
Parameters refined	122
Goof (obs/all)	1.006/0.966
R (obs), $wR$ (obs)	0.0387, 0.0847
R (all), $wR$ (all)	0.0504, 0.0892
Largest diff. peak/hole	+2.34/-1.64 e

172, 161, 144 and 125 cm<sup>-1</sup> are assigned to either  $v_2 (\delta) (UO_2)^{2+}$  bending modes, U–O<sub>ligand</sub> stretches, or external lattice vibration modes.

# **Chemical composition**

Several crystals of leószilárdite were analysed using a Cameca SX50 electron microprobe (University of Notre Dame), operating with an accelerating voltage of 15 kV, a beam current of 10 nA and beam diameter of 25 µm. The following standards were used: NaK $\alpha$  (albite), MgK $\alpha$  (olivine), UM $\alpha$ (synthetic  $UO_2$ ). Other elements were sought but not detected. The counting time for each peak was 8 s, as was the counting time for each background. Matrix effects were accounted for using the PAP correction routine (Pouchou and Pichoir, 1985). The mineral contains a high proportion of volatile sodium, and the fragility of crystals under the beam led to slightly low totals (~97%) and required the analysis of multiple crystals using a wide beam. The CO<sub>2</sub> and H<sub>2</sub>O contents could not be measured due to the limited amount of material available; instead they are calculated based on the structural formula with 6 C, and 28 O atoms per formula unit, and with charge balance considerations satisfied by H. Analytical data averaged for six analyses from multiple crystals are given in Table 1. The empirical formula is  $Na_{5.60}Mg_{0.90}U_2O_{28}C_6H_{12.60}$ . The ideal formula is  $Na_6Mg(UO_2)_2(CO_3)_6 \cdot 6H_2O$ , which requires Na<sub>2</sub>O 15.88, MgO 3.45, UO<sub>3</sub> 48.88, CO<sub>2</sub> 22.55 and H<sub>2</sub>O 9.24, total 100 wt.%.

#### **Powder X-ray diffraction**

Room temperature powder diffraction data (Table 2) were recorded using a Rigaku R-Axis Rapid II curved imaging plate microdiffractometer with monochromated MoK $\alpha$  radiation. A Gandolfilike motion on the  $\varphi$  and  $\omega$  axes was used to randomize diffraction from the sample. Observed *d*-values and intensities were derived by profile fitting using *JADE 2010* software (Materials Data, Inc.). Unit-cell parameters refined from the powder data are as follows: a = 11.588(6) Å, b = 6.767(6) Å, c = 15.081(6) Å,  $\beta = 91.453(9)^\circ$ , V = 1182.2(13) Å<sup>3</sup>.

### Single-crystal X-ray diffraction

Single crystal data were collected from an optically homogeneous bladed crystal using Mo $K\alpha$  X-rays from a microfocus source and an Apex II CCDbased detector mounted to a Bruker Apex II Quazar

Atoms	x	У	Ζ	$U_{\rm eq/iso}$	$U^{11}$	$U^{22}$	$U^{33}$	$U^{23}$	$U^{13}$	$U^{12}$
U1	0.312623(3)	0.5	0.72748(2)	0.01418(13)	0.0113(2)	0.0176(2)	0.0137(2)	0	0.00213(13)	0
Na1	0.3146(4)	0	0.83965(29)	0.0220(6)	0.0188(14)	0.0249(17)	0.0223(18)	0	0.0011(13)	0
Na2	0.4274(2)	0.2611(5)	0.3783(2)	0.0274(7)	0.0298(18)	0.0259(17)	0.0267(17)	-0.0019(12)	0.0042(13)	-0.0032(13)
Mg1	0.5	0	1	0.0163(9)	0.0128(19)	0.0216(25)	0.0146(17)	0	0.0050(14)	0
CĨ	0.6045(10)	0	0.8216(7)	0.0199(20)	0.0189(19)	0.032(6)	0.0093(23)	0	0.0015(19)	0
C2	0.3150(9)	0.5	0.5369(7)	0.0202(14)	0.016(3)	0.029(4)	0.0165(14)	0	-0.0008(18)	0
C3	0.0336(10)	0	0.8234(7)	0.0201(20)	0.0177(18)	0.028(6)	0.014(3)	0	0.0003(20)	0
01	0.6033(6)	0	0.7349(4)	0.0142(14)	0.017(4)	0.017(4)	0.008(2)	0	0.002(2)	0
O2	0.0303(6)	0	0.7359(5)	0.0149(14)	0.019(4)	0.013(3)	0.0129	0	-0.002(2)	0
O3	-0.0659(6)	0	0.8611(5)	0.0201(15)	0.016(2)	0.027(4)	0.017(3)	0	0.001(2)	0
O4	0.4069(6)	0.5	0.5869(5)	0.0225(12)	0.014(3)	0.038(4)	0.0158(19)	0	0.0022(16)	0.0225(12)
O5	0.2193(6)	0.5	0.5825(5)	0.0205(12)	0.014(2)	0.031(3)	0.017(2)	0	-0.001(2)	0
06	0.3153(7)	0.5	0.4549(5)	0.0243(12)	0.019(4)	0.038(4)	0.0155(14)	0	0.0002(20)	0
O7	0.3151(5)	0.2359(9)	0.7269(3)	0.0237(10)	0.030(3)	0.0208(16)	0.0196(21)	-0.001(1)	-0.0036(17)	-0.001(2)
08	0.5152(7)	0	0.8644(5)	0.0221(14)	0.0183(16)	0.036(4)	0.0124(20)	0	0.0029(16)	0
09	0.7062(7)	0	0.8605(5)	0.0345(22)	0.0179(20)	0.075(7)	0.010(3)	0	0.001(2)	0
O10	0.1271(7)	0	0.8665(5)	0.0401(22)	0.0173(16)	0.089(7)	0.014(3)	0	-0.0004(17)	0
Ow1	0.5647(7)	0	0.4168(6)	0.0293(16)	0.020(4)	0.037(4)	0.030(5)	0	0.002(3)	0
Ow2	0.6259(6)	0.2133(12)	1.0210(4)	0.0493(21)	0.041(3)	0.078(5)	0.029(3)	-0.025(3)	0.0176(23)	-0.042(3)

TABLE 4. Atomic coordinates and displacement parameters (Å <sup>2</sup> ) for leószilárdite.	
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U1-O7 (×2)	1.792(6)	Na2-O1	2.482(6)
U1-01	2.477(7)	Na2-O2	2.426(6)
U1-O2	2.486(7)	Na2-O4	2.560(7)
U1-O3	2.410(7)	Na2-O5	2.537(6)
U1-04	2.392(7)	Na2-O6	2.394(7)
U1-O5	2.438(7)	Na2-Ow1	2.444(7)
U1-09	2.407(8)	<na2-o></na2-o>	2.47
<u1-0<sub>Ur&gt;</u1-0<sub>	1.792		
$< U1 - O_{Eq}^{O} >$	2.429		
$Mg1 - O\overline{8}^{1}(\times 2)$	2.061(7)	C1-O1	1.309(12)
Mg1-Ow2 (×4)	2.075(6)	C1-O8	1.235(13)
<mg1-o></mg1-o>	2.070	C1-O9	1.307(14)
		<c1-o></c1-o>	1.28
Na1-O10	2.223(9)	C2-O4	1.293(13)
Na1-O7 (×2)	2.337(6)	C2-O5	1.322(13)
Na1-O8	2.350(9)	C2-O6	1.239(13)
Na1-Ow2 (×2)	2.633(9)	<c2–o></c2–o>	1.28
<na1-o></na1-o>	2.32		
		C3-O2	1.321(12)
		C3–O3	1.300(13)
		C3-O10	1.253(13)
		<c3–o></c3–o>	1.29

TABLE 5. Selected bond distances in leószilárdite.

three-circle diffractometer. Reflections were integrated and corrected for Lorentz, polarization and background effects using the Bruker program SAINT. A multi-scan semi-empirical absorption correction was applied using equivalent reflections in SADABS-2012. An initial structure model was obtained by the charge-flipping method using SHELXT (Sheldrick, 2015) in the space group C2/m. Refinement proceeded by full-matrix leastsquares on  $F^2$  using SHELX-2013 (Sheldrick, 2008), and refined to an  $R_1$  of 3.87% for 1394 reflections with  $I_{obs} > 4\sigma(I)$ . Rigid bond restraints were applied to the anharmonic displacement parameters of all atoms using the DELU and RIGU instructions in SHELXL. Hydrogen atom positions were sought but were not located in difference-Fourier maps, as a result the nature of H-bonding interactions is not discussed. Further details regarding the data collection and refinement are given in Table 3. Final atom coordinates and displacement parameters are listed in Table 4, selected interatomic distances in Table 5, and the bond-valence sums are given in Table 6. The original crystallographic information file (cif) has been deposited with the Principal Editor of Mineralogical Magazine and is available from http://www.minersoc.org/pages/e\_journals/dep\_mat mm.html

### Features of the crystal structure

In the structure of leószilárdite there is one independent U, three C, one Mg, two Na and 12 O sites (Z=2). It contains hexagonal uranyl bipyramids chelated by three CO<sub>2</sub> groups, and has a cluster topology based upon uranyl tricarbonate anions (UTC) of the formula  $[(UO_2)(CO_2)_2]^{4-1}$ (Burns, 2005). Leószilárdite is a member of the uranyl tricarbonates, Strunz class 05.ED, with UO<sub>2</sub>:  $CO_2 = 1:3$ . It bears most resemblance to čejkaite (Plášil et al., 2013), synthetic  $Na_4(UO_2)(CO_2)_2$ (Li et al., 2001; Císařová et al., 2001) and ježekite (Plášil et al., 2015) in that the structures of all of these phases contain chains of face-sharing sodium polyhedra linked to the uranyl tricarbonate clusters. However, the structure of leószilárdite is unique amongst the uranyl carbonates with respect to the arrangement of Na<sup>+</sup> and Mg<sup>2+</sup> cations, and represents the first natural uranyl carbonate containing both Na and Mg as essential constituents.

The uranyl hexagonal bipyramid shares two edges with face-sharing chains of  $Na(H_2O)O_5$  trigonal prisms (Na2) that extend along [010] (Fig. 6). The structure also contains a distinctive three-octahedron unit of the formula  $Na_2MgO_{12}(H_2O)_4$  consisting of two Na-centred octahedra that share the opposite faces of a Mg-

\*Multiplicities indicated by  $\times 2 \rightarrow$ ;  $\times 2 \rightarrow$ ;  $\operatorname{Na}^+$ -O from Wood and Palenik (1999);  $\operatorname{Ma}^{2+}$ -O and  $\operatorname{Cd}^+$ -O bond strengths taken from Brown and Altermatt (1985);

Ow2		0.09 × 2→			0.36	× 44				0.46
Ow1			0.16	× 2↓						0.35
O10		0.28							1.45	1.77
60	0.50						1.25			1.75
08		0.20			0.37	$\times 2 \downarrow$	1.52			2.12
07	$1.65 \times 2 \rightarrow$	$0.21 \times 2 \rightarrow$								1.88
06			0.18	× 2↓				1.50		1.91
05	0.47		0.12	$\times 2\downarrow$				1.20		2.03
04	0.52		0.11	$\times 2\downarrow$				1.30		2.08
03	0.50								1.28	1.77
02	0.43		0.16	× 2↓					1.20	2.01
01	0.44		0.14	× 2↓			1.25			2.00
Atom	UI	Na1	Na2		Mg1		C1	C2	C3	$\Sigma_{ m an}$

2.16

0.87

4.02 4.01 3.93

6.16 1.08

 $\Sigma_{\rm cat}$ 

centred octahedron at the centre (Na1–Mg1–Na1). These three-octahedron units, aligned along [101], further connect the UTC units (Fig 7).

Chains of Na-centred polyhedra found in čejkaite coordinate to UTC units by sharing only apices with carbonate triangles from six individual uranyl polyhedra. In contrast, each Na-centred polyhedron in the chains of ježekite shares one edge with a uranyl polyhedron and one apex with the carbonate triangle of a different UTC unit. Individual Na-centred octahedra in the chains of leószilárdite coordinate to three UTC units, sharing two edges of uranyl polyhedra and one apex of a CO<sub>3</sub> triangle, each (Fig. 8). The structural formula obtained from the X-ray data is Na<sub>6</sub>Mg(UO<sub>2</sub>)<sub>2</sub>(CO<sub>3</sub>)<sub>6</sub>·6H<sub>2</sub>O, Z=2, with  $D_{calc}=3.269$  g cm<sup>3</sup>.

Hazen et al. (2016) predicted 16 yet undiscovered UTC minerals on the basis of isomorphous substitutions of some elements common to natural groundwater into known structures. They also predict 11 other UTC phases, bearing similar compositions or structures to known synthetic phases. Leószilárdite was not amongst the predicted phases, but highlights an unfortunate downside to estimations like these; in that our inability to predict novel arrangements of cation polyhedra in uranyl carbonate minerals prevents us from fully understanding these systems. Many of our recent descriptions of uranyl carbonates from Red Canyon, Utah, USA and Jáchymov, Czech Republic bear little to no resemblance to the sensible chemical substitutions expected in natural minerals, nor to any synthetic phases not yet found to occur naturally. Rather, nearly all recent new uranyl carbonate minerals constitute novel topological arrangements; in addition to a unique class of hybrid uranyl carbonate cluster minerals that contain bridging uranyl pentagonal bipyramid clusters. We are in the process of describing two such minerals; the first includes a combination of uranyl tricarbonate (UTC) and dicarbonate units, which share apices with trimeric  $[(UO_2)_3O(OH)_3]^+$  cluster units in ewingite (IMA2016-012). The second is an unnamed Ca, UTC mineral found by one of the authors (J.M.) in the Markey Mine, and includes UTC-bridged monomeric UO<sub>7</sub> units. Several uranyl carbonate minerals containing uranyl pentagonal bipyramids arranged into sheets are known; including fontanite (Hughes and Burns, 2003), wyartite (Burns and Finch, 1999), bijvoetite-Y (Li and Burns, 2001) and roubaultite (Ginderow and Cesbron, 1985). However, the structures of ewingite and the unnamed Markey mine phase are unique as the

 $J^{6+}-O$  bond strength from Burns *et al.* (1997).

TABLE 6. Bond-valence sums for leószilárdite.\*

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FIG. 6. A polyhedral representation of the structure of leószilárdite down b. The unit cell is shown as dashed lines.



FIG. 7. Irregular chains of Na-polyhedra display varying degrees of coordination to UTC units in the related minerals čejkaite, leószilárdite and ježekite. Uranium (yellow), sodium (blue), carbon (black). Various cations have been omitted to highlight the chains of Na-polyhedra in each structure.



FIG. 8. The coordination environment about three-octahedron Na1–Mg1–Na1 units in the structure of leószilárdite, and their connectivity to uranyl and carbonate oxygen atoms.

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first cluster, and open sheet topologies to contain this combination, respectively.

The rarity of leószilárdite is due mostly to the ephemeral nature of the mineral, and it may be a more common constituent in uranyl carbonate assemblages where the rate of dissolution is slow, and Na<sup>+</sup> and Mg<sup>2+</sup> are available. Newly described uranyl carbonate mineral structures motivate our work while supplementing our observations of the chemistry of uranium in carbonated groundwater. The search for rare uranyl carbonate structures continues, as it aids in delineating the complicated series of dissolution and precipitation reactions that these minerals are involved in.

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