Radiocarbon, Vol 61, Nr 5, 2019, p 1211-1220

Selected Papers from the 23rd International Radiocarbon Conference, Trondheim, Norway, 17–22 June, 2018 © 2019 by the Arizona Board of Regents on behalf of the University of Arizona

NEW RADIOCARBON DATING RESULTS FROM THE UPPER PALEOLITHIC–MESOLITHIC LEVELS IN GROTTA ROMANELLI (APULIA, SOUTHERN ITALY)

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ABSTRACT. In this paper, we present the results of the accelerator mass spectrometry radiocarbon (AMS¹⁴C) dating campaign performed on samples selected from different levels in Grotta Romanelli (Castro, Italy). Grotta Romanelli is one of the key sites for the chronology of Middle Pleistocene–Holocene in Mediterranean region. After the first excavation campaigns carried out in the first decades of the 1900s, the cave has been systematically re-excavated only since 2015. During the last excavation campaigns different faunal remains were selected and submitted for ¹⁴C dating in order to confirm the chronology of the cave with a higher resolution. Isotopic ratio mass spectrometry (IRMS) measurements were also carried out on faunal remains.

KEYWORDS: AMS, Bayesian OxCal model, Middle Pleistocene–Holocene, stable isotopes, Salento peninsula.

INTRODUCTION

"Grotta Romanelli" (Figure 1) (40°00'58.30"N, 18°26'00.01"E) is a natural cave located along the Ionian coast of the Salento Peninsula in southeastern Italy in the territory of Castro (Lecce, Apulia). The cave, regarded as the first Upper Paleolithic deposit in Italy, represents a key site for understanding of the Mediterranean Middle Pleistocene–Holocene regional context and establishing the framework of human presence in the Mediterranean region. The deposit was discovered in 1871, but it was only thanks to the systematic and pioneering excavations carried out at the beginning of 1900 that its great archaeological importance was assessed and the presence of fossil vertebrate assemblages recognized (Stasi and Regalia 1904; Blanc 1920, 1928; Spinapolice 2018).

The stratigraphy of the cave subdivided in two main complexes, as proposed by Blanc (1928; Figure 2), is regarded as a reference for the definition of the Late Pleistocene and Palaeolithic chronology in Italy. The infilling deposit is bounded by Cretaceous limestone (level L), at 7.4 m asl, that Blanc considered shaped during MIS5, constraining the age of the deposits to the Late Pleistocene. The lower complex includes the beach deposit (K), the bone breccia (I), the stalagmitic layer (H) and the *terre rosse* (red soils—level G) with large mammals and limestone artifacts referring to the Middle Paleolithic. The two complexes are divided by the sub-horizontal stalagmitic layer (level F). The upper complex, known as *terre brune* (brown

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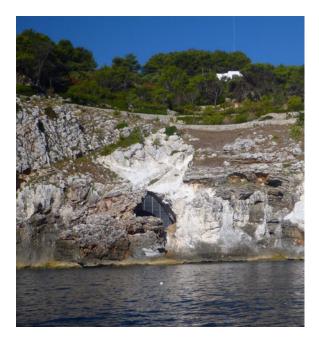


Figure 1 Grotta Romanelli opens into Cretaceous limestone (layer L in Figure 2) along the Adriatic coast of the Salento peninsula in southern Italy.

soils—levels E–A), is characterized by Upper Paleolithic artifacts and by a diversified fossil vertebrate assemblage including mammals and more than one hundred species of birds.

Blanc (1920) referred the rich lithic material recovered from the *terre brune* to the end of the Late Pleistocene. Based on the typological analysis of the artifacts recovered from the upper part of the infilling succession, the so called Romanellian facies was identified (Taschini and Bietti 1972; Mussi 2002; Bietti 2003; Palma di Cesnola 2003; Martini et al. 2017). The peculiarity of the lithic complex and the chronology of the *terre brune*, corroborated by radiometric age (Bella et al. 1958; Vogel and Waterbolk 1963; Alessio et al. 1964, 1965), set Grotta Romanelli as a key site for understanding the framework of human evolution in the Mediterranean area during the end of the Late Pleistocene–Early Holocene (Mussi 2002; Bietti 2003).

Albeit the scholars have long considered the site of international significance, many issues need to be investigated and clarified (see Sardella et al. 2018 for discussion).

A systematic excavation campaign started in 2015 under the scientific supervision of the University of Rome "La Sapienza" with several aims: (1) increase the resolution of the site stratigraphy; (2) revise and update the existing palaeontological and archaeological research; (3) develop conservation strategies of the infilling sequence (deeply excavated and eroded but still well preserved, at least in the lower part). The state of the art of the research about the cave, a brief report on the preservation conditions and first results obtained during the new excavation campaigns have been recently summarized (Sardella et al. 2018). New mineralogical and sedimentological analyses of the *terre brune* sequence confirmed the eolian origin of the sediments and allowed the reconstruction of

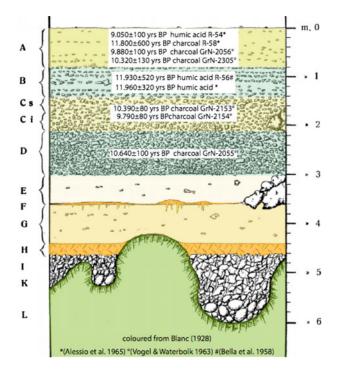


Figure 2 Schematic stratigraphy of the cave with the radiocarbon ages so far available for the different layers.

environmental variations related to climate changes, between the end of the last glaciation and the Early Holocene (Giustini et al. 2018).

Our research aims to define an absolute chronology of the different levels of the *terre brune* so far based on only nine ¹⁴C dates measurement, sometimes scarcely consistent and with large statistical uncertainties, performed on humic acid (sample R54 and R56), and charcoal samples from the uppermost layers in the 1960s (Figure 2).

MATERIALS AND METHODS

Fourteen animal bones (Table 1) were sampled during the 2015–2016 excavation campaigns from the different layers identified in the different stratigraphic sections opened in the cave (Figure 3).

Collagen was extracted from the samples by using the Longin protocol (Longin 1971) at the chemical laboratories of the Centre for Applied Physics, Dating and Diagnostics (CEDAD)-Department of Mathematics and Physics "Ennio de Giorgi"-University of Salento (D'Elia et al. 2004). A fraction of the extracted collagen was separated from nine of the samples for the determination of the C:N ratio and for carbon and nitrogen stable isotope analyses which were performed by using an elemental analyzer (EA-Mod. Flash 2000 HT by Thermo) and an isotopic ratio mass spectrometry (IRMS) system (Delta V Plus by Thermo; Braione et al. 2015; Maruccio et al. 2017). The fraction selected for ¹⁴C analysis was combusted to CO₂ in sealed quartz tubes with CuO and Silver wool and then reduced at 600°C to graphite with H₂ on Fe powder used as catalyst. AMS ¹⁴C measurement were

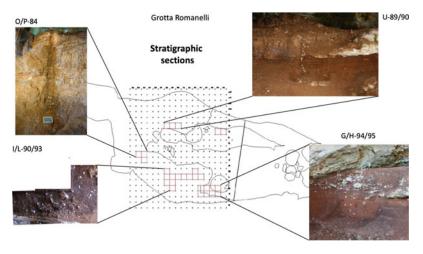


Figure 3 Planimetry and stratigrafic sections of the Grotta Romanelli Cave (Castro-Italy).

carried out with the 3 MV Tandetron at CEDAD-University of Salento (High Voltage Engineering Europa BV Mod. 4130HC; Calcagnile et al. 2004, 2005). The measured ${}^{14}C/{}^{12}C$ isotopic ratios were corrected for isotopic fractionation by using the $\delta^{13}C$ term measured on line with the accelerator, and for machine and chemical processing background. The conventional ${}^{14}C$ ages were then calculated according to Stuiver and Polach (1977) and calibrated by using the last internationally accepted IntCal13 calibration curve (Reimer et al. 2013) and the OxCal ver. 4.3 software (Bronk Ramsey and Lee 2013).

RESULTS AND DISCUSSION

The measured ¹⁴C ages are shown in Table 1 together with the determined C/N ratios. It can be noted that all the samples gave C/N ratios within the range considered indicative of a good preservation of collagen (2.9–3.6; van Klinken 1999). Only sample LTL17303A gave a C/N ratio significantly outside the optimal range. Nevertheless, no other indication of poorly preserved collagen has been highlighted (C and N stable isotopic ratios and collagen yield) so this result was not excluded a priori from the following analysis.

The analysis of the whole set of ¹⁴C ages was also performed by using the advanced Bayesian tools available in OxCal Ver. 4.3 (Bronk Ramsey 2009a). A *Model* was then generated in which all the samples were grouped in different *Phases* divided by *boundaries* and forming a *Sequence* named *Romanelli*. The identification of possible *outliers* was carried out by using the dedicated routines available in OxCal and by following Bronk Ramsey (2009b).

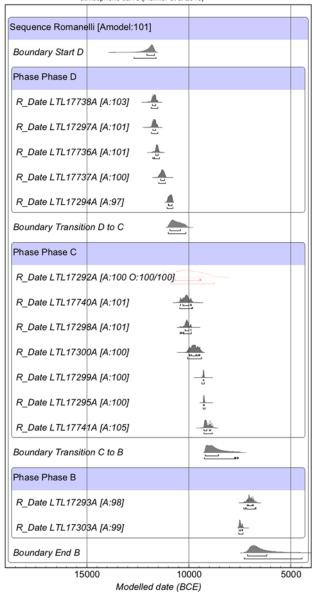
Indeed, one of the samples (LTL17292A) can be considered an outlier and was then removed from the following analysis whose results are given in Figure 4 and Table 1.

It can be seen that all the obtained results are consistent with the stratigraphic position of the samples. In particular the phase D is dated between 12,000–11,000 cal BC, the phase C between 11,000–9,000 cal BC and phase B to 8th–7th millennium BCE.

Sample	Pit	Lab code	Phase	C/N	¹⁴ C age	Modeled calibrated age (cal BC, 2σ)
GR2016 106		LTL17293A	В	3.5	8048 ± 75	7302–7223
						7190–6746
GR2016-162-US2	H95	LTL 17303A	В	3.9	8397 ± 45	7555-7350
US1	189	LTL 17741A	С	3.6	9657 ± 65	9265-8846
GR2016 -154-US988	P84	LTL 17295A	С		9774 ± 40	9300-9204
GR2016- 158-US983	P84	LTL 17299A	С	3.6	9822 ± 45	9362-9231
GR2016-105-US995	O84	LTL 17292A	С		$11,328 \pm 60$	
GR2016-159-US 998	P84	LTL 17300A	С	3.4	$10,100 \pm 80$	10,066–9386
GR2016-157-US998	M88	LTL 17298A	С		$10,277 \pm 45$	10,288–9870
US14	189	LTL 17740A	С	3.6	$10,295 \pm 75$	10,444–9852
GR2016-153	H95	LTL 17294A	D		$10,990 \pm 50$	11,058-10,784
GR2016-622-US1004	U89	LTL 17737A	D	3.5	$11,409 \pm 85$	11,473–11,144
GR2016-581-US1005	U 90	LTL 17736A	D	3.5	$11,685 \pm 65$	11,761–11,436
GR2016-156-US1005	Q84	LTL 17297A	D	3.3	$11,829 \pm 80$	11,821–11,521
GR2016-616	U89	LTL 17738A	D	3.6	$11,858 \pm 85$	11,860-11,516

Table 1 Analyzed samples, C/N ratios, and conventional ¹⁴C ages.

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OxCal v4.3.2 Bronk Ramsey (2017); r:5 IntCal13 atmospheric curve (Reimer et al 2013)

Figure 4 Calibrated ¹⁴C ages and Bayesian OxCal model.

The comparison with the previously available data is given in Figure 5 where the current ${}^{14}C$ ages are shown together previous ones (indicated as "old" in the figure). It can be immediately recognized that the current study allowed to significantly constraint the previous results which were highly scattered and only partially consistent with the stratigraphic provenance of the samples. In particular, phase D is significantly older than what previously obtained on the basis of the ${}^{14}C$ results on one single sample. The age obtained for the Phase C is essentially consistent with previous determinations. The results so far available for the

Phase D Phase LT L17738A R_Date(11858,85) LT L17297A R_Date(11829,80) LT L17297A R_Date(11829,80) LT L17736A R_Date(11829,80) LT L17737A R_Date(11409,85) LT L17294A R_Date(11409,85) LT L17295A R_Date(10990,50) Phase D-Old Phase GRN-2055 R_Date(10640,100) Phase C Phase LT L17740A R_Date(10295,75) LT L17298A R_Date(10277,45) LT L17300A R_Date(10100,80) LT L17295A R_Date(9822,45) LT L17295A R_Date(9774,40)	0		
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LTL17295A R_Date(9774,40)		0	
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LTL17741A R_Date(9657,65)			
Phase C-Old Phase			
GRN-2153 R_Date(10390,90)			
GRN-2154 R_Date(9790,80)			
Phase B Phase			
LTL17303A R_Date(8397,45)			
LTL17293A R_Date(8048,75)			
Phase B-Old Phase			
R-56# R_Date(11960,520)			
R_Date(11930,320)			
Phase A-Old Phase			
R-58* R_Date(11800,600)			
GrN-2305° R_Date(10320,130)			
GrN-2056° R_Date(9880,100)			
R-54* R_Date(9050,100)			
20000 18000 16000 140 Calibrated date			

OxCal v4.3.2 Bronk Ramsey (2017); r:5 IntCal13 atmospheric curve (Reimer et al 2013)

Figure 5 Comparison of the results obtained in the present study and previous ¹⁴C data from the same stratigraphic units. Dark and light gray indicate one and two standard deviations confidence levels intervals, respectively.

Phase B were non-consistent with the stratigraphy of the cave and characterized by large uncertainties on the single measurements. These discrepancies are now solved by the results obtained on the samples LTL17293A and LTL17303A which gave a ¹⁴C age consistent with the cave stratigraphy and significantly younger than previous determinations.

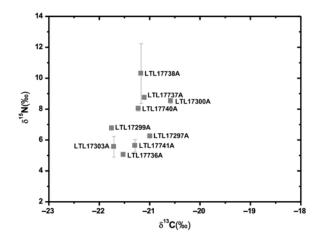


Figure 6 Carbon and nitrogen stable isotopic data obtained by IRMS on some of the analyzed samples. Error bars represent the scattering obtained on the replicate measurements performed on each sample, when no error bars are visible they are included in the size of the symbol. Data are expressed in the usual δ notation, nitrogen vs. AIR standard and carbon vs. V-PDB.

The results of IRMS measurement of stable carbon and nitrogen content are given in Figure 6. It can be seen that δ^{13} C values range from -21.75 to $-20.58\%_0$ and are then typical of animals feeding of C3 plants. The δ^{15} N values range from 5.05 to 10.30‰. Figure 6 shows that the data appear to fall into two groups, mainly because of the nitrogen isotope compositions; the first group is approximately centred around $+6\%_0$, whereas the second group around $+8.5\%_0$. The difference (about 2.5‰) supports the conclusion that the two groups are representative of two trophic levels, likely herbivores and carnivores (Post 2002). The identification of one sample (LTL17736A) as Cervidae could confirm the above distinction.

CONCLUSIONS

New systematic excavations carried out in 2015–2017 in Grotta Romanelli allowed us to assess the stratigraphic sequence identified at the beginning of 1900s. Animal bone samples were selected and AMS ¹⁴C dated. C/N ratio data indicate that the collagen is well preserved.

The results expand and refine the previous chronology. Phase D resulted much older, encompassing the Late Pleistocene-Holocene boundary. The new age of the uppermost part of the stratigraphy (level B) is quite different from the older data, extending the chronology to Mesolithic (Northgrippian, Middle Holocene). Considering the great importance of the site in the Late Paleolithic framework of the Mediterranean area, new typological analysis of the lithic complex recovered from the *terre brune* will be required.

Finally, further studies are already in progress for refining the chronology of the *terre brune*, with a particular focus on the level E (no ¹⁴C age available) and level A (if still present in the cave). New ¹⁴C dating enable to extending the data of the older stratigraphic units to provide a detailed chronology of human occupation in the cave.

ACKNOWLEDGMENTS

The undergoing excavation campaigns are being carried out with the permission of the "Soprintendenza archeologia, belle arti e paesaggio delle province di Brindisi, Lecce e Taranto" (former Soprintendenza archeologica) whose support is deeply acknowledged, and granted by Sapienza, Rome university (Progetto Grandi Scavi 2017, resp. Raffaele Sardella).

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