

Silurian tarphycerid *Discoceras* (Cephalopoda, Nautiloidea): systematics, embryonic development and paleoecology

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Abstract.—Tarphycerids were diverse and abundant in Ordovician marine faunas. Beginning at the Late Ordovician extinction, the diversity of tarphycerids declined throughout the Silurian, until their extinction in the latest Silurian. Two genera survived the Late Ordovician extinction: *Trocholites* Conrad, 1838 (from which *Ophioceras* Barrande, 1865 probably diverged) and *Discoceras* Barrande, 1867 (= *Graftonoceras* Foerste, 1925). *Discoceras graftonense* (Meek and Worthen, 1870), so far known from the US, China, and Australia, is recorded from the Silurian of Bohemia and Gotland. *Discoceras stridsbergi* n. sp., *D. lindstroemi* n. sp., and *D.* sp. indet. from the Wenlock of Gotland and *D. amissus* (Barrande, 1865) from the Llandovery of Bohemia are all endemic species probably derived from *D. graftonense*. The distribution of *D. graftonense* and the origin of four species of *Discoceras* in the latest Sheinwoodian and early Homeric represent the last diversification and dispersion of the Tarphycerida. No tarphycerid species originated after the mid-Homeric extinction (Mulde and Lundgreni events). Silurian *Discoceras* retained the morphology and habitats of their Ordovician ancestors. The hatching time and autecology of juveniles has remained unclear. Evidence from the material studied suggests that juveniles were planktonic in habit, possessing a minute curved shell with few phragmocone chambers. *Discoceras lindstroemi* n. sp. is exceptional owing to its heteromorphic planispiral shell with coiling that changed during ontogeny, resulting in a changing aperture orientation and decreased maneuverability.

Introduction

Order Tarphycerida (Early Ordovician–latest Silurian) includes nautiloids with a tightly coiled planispiral shell, a wide siphuncle that changed its position during ontogeny, layered and thick connecting rings, and an aperture possessing a deep, narrow hyponomic sinus (Furnish and Glenister, 1964; Dzik, 1984). Tarphycerids inhabited shallow tropical and subtropical epeiric seas (Flower, 1957; Westermann, 1998), but occasionally spread pole-ward, into open sea or deeper water settings (Babin and Gutiérrez-Marco, 1992; Manda, 2008a; Kröger et al., 2009). The planispiral shell and deep hyponomic sinus indicate that tarphycerids were capable of active swimming and were, like *Nautilus*, most probably demersal animals (Dzik, 1984; Westermann, 1998). In some species, the last whorl becomes uncoiled, which indicates a changed mode of life from nektonic towards a demersal habit, with limited swimming ability (Flower, 1955). The first tarphycerids appeared in the early Tremadocian, and their diversity suddenly increased, reaching a maximum in early Floian time. Their generic diversity fluctuated only slightly, until a slow diversity decline in the middle Katian and Hirnantian (Kröger and Zhang, 2009). The tarphycerida included the vast majority of Ordovician cephalopods possessing coiled shells.

The diversity of tarphycerids after the Late Ordovician extinction was very low throughout the Silurian. Although some species were locally abundant, only three genera/families are

known. *Ophioceras* Barrande, 1865 is the single known representative of the family Ophioceratidae Hyatt, 1894, and it is the only tarphycerid genus to have originated in the Silurian. It comprises only two long-ranging species, occurring in middle and upper Silurian rocks of the US, England, Bohemia, Poland, Sweden, Estonia (Stridsberg and Turek, 1997), and China (Chen et al., 1981). *Trocholites* Conrad, 1838 (Trocholitidae Chapman, 1857) includes more than 20 Ordovician species, but is represented in the early Silurian by only a single species known from the Anglo-Welsh Basin (Holland, 2010). *Discoceras* Barrande, 1867 (= *Graftonoceras* Foerste, 1925 and *Baterboceras* Zou, 1983; Discoceratidae Dzik, 1984) is a genus with more than 30 Middle and Upper Ordovician species. Including the four species described here and evaluating previously published data, *Discoceras* includes six Silurian species occurring in Llandovery strata in Bohemia, and Wenlock strata in North America (Foerste, 1925), Australia (Etheridge, 1904), Inner Mongolia (Zou, 1983), Sweden, and Bohemia.

We describe here three new Silurian species of *Discoceras* from Gotland (Sweden) and Central Bohemia (Czech Republic). Silurian *Discoceras* retained the tarphycerid shell morphology and habitat of its Ordovician ancestors, but one new species—*D. lindstroemi* n. sp.—exhibits an unusual coiled shell, resembling that of some heteromorph ammonoids.

Well-preserved specimens of Silurian *Discoceras* from Gotland have made possible the study of the early growth stages of the shell as well as the reconstruction of the early ontogeny of

the genus. This study supports the conclusion that early-hatched tarphycerids possessed minute curved shells, and were planktonic, rather than having the coiled shell and life mode of adults. A planktonic habit for Ordovician tarphycerid juveniles was suggested by Holland (1985). This has been documented by Turek and Manda (2016) in Silurian *Ophioceras*, an evolutionarily younger tarphycerid, displaying several evolutionary novelties unknown in other tarphycerids, even those from the Silurian. Thus, to expand the hypothesis of a planktonic habit of juveniles to other tarphycerids, more evidence from evolutionarily basal tarphycerids was needed.

Geological setting and taphonomy

Studied material comes from Silurian rocks of Central Bohemia (Bohemian Massif, Czech Republic) and Gotland (Baltic Shield, Sweden).

The Central Bohemian Paleozoic is a part of the Barrandian area consisting of Proterozoic and Paleozoic rocks (Teplá-Barrandian Unit, Bohemian Massive), which represents the peri-Gondwanan terrain, Perunica (see Torsvik and Cocks, 2013). Silurian rocks are preserved in the Prague Synform, a structure formed during Variscian Orogeny. Silurian rocks form a part of the marine sedimentary succession of the so-called Prague Basin (Ordovician–Middle Devonian). Five lithostratigraphic units were established in the Silurian strata (for summary see Kříž, 1998), of which two yielded *Discoceras*. The first unit, the Želkovice Formation, is of Rhuddanian and Aeronian age, consists of black laminated graptolitic shales, and has a maximum thickness of ~25 m (Kříž, 1998; Štorch, 2006). The shale was deposited under anoxic conditions in an offshore setting; there is no evidence of near-shore conditions (Štorch, 2006). Benthic fauna includes very rare rhynchonelliform and common linguliform brachiopods that were described from the lowermost and uppermost parts of the formation, respectively (Kříž, 1998). *Discoceras* occurs in black shales that lack benthic fauna. Graptolites indicate the *Demirastrites triangulatus*, *D. simulans*, and *D. convolutes* biozones (as defined by Štorch, 2006). These graptolitic shales contain rare *Discinocaris* Woodward, 1866 and *Peltocaris* Salter, 1863 (considered to be cephalopod opercula) (Turek, 1978). Despite over 150 years of intensive collecting throughout the formation, only two flattened shells and one fragment of *Discoceras* are available for study.

The second unit, the Motol Formation, contains an up to 300 m thick sedimentary succession ranging in age from upper Llandovery to uppermost Wenlock (Kříž, 1998). The formation consists of alternating volcano-sedimentary complex sets and shale facies. The single available flattened specimen of *Discoceras* comes from shale in the lower part of the formation, which has been referred to the *Cyrtograptus murchisoni* Biozone as defined by Štorch (1994). Oxygen deficient, offshore graptolite-rich shales dominate the lower part of the formation; the volcanic activity is limited to occasional basalt effusions (Kříž, 1998). Shales and platy muddy limestones with a benthic *Niorhinx* Community (Havlíček and Štorch, 1990) developed locally in areas of volcanic activity and indicate the presence of low and isolated bottom elevations swept by bottom currents (Kříž, 1998). The community includes

rhynchonelliform brachiopods, trilobites, crinoids, gastropods (Havlíček and Štorch, 1990), and cephalopods (straight shelled pelagic forms, the demersal orthocerid *Dawsonoceras annulatum* [Sowerby, 1818], the discosorid *Phragmoceras munthei* Hedström, 1917 [Manda, 2008b], and the herein described tarphycerid *Discoceras*).

The Silurian bedrock of Gotland is a remnant of an extensive low-latitude carbonate platform complex that evolved along the margins of the Baltic Basin and extended from the western parts of the present-day Baltic Sea across the East Baltic and farther southeast to Ukraine (for summary see Calner et al., 2002). The majority of the *Discoceras* specimens were collected by G. Lindström in the nineteenth century; some of his associated labels indicate his unrealized goal of describing four new species. Precise localities are not specified; rather, the locality is given as the name of the nearest village in the same land registry as the actual locality. Preservation states of the studied specimens, however, facilitate identification of their original horizons and locations (Laufeld, 1974; Stridsberg, 1985). All studied specimens were collected from several sites in the lower part of the Slite beds (units d and g), which correspond with the upper Sheinwoodian to lower Homerian, *Kockellella walliseri* to *Ozarkodina sagitta sagitta* conodont biozones and the *Monograptus belophorus* to *Cyrtograptus lundgreni* graptolite biozones, respectively (Jeppsson et al., 1994; Jeppsson and Calner, 2003). Studied specimens were collected from light-colored stromatoporoid boundstone, crinoidal grainstone, and dolomite—facies that form biohermal, biostromal, and shoal areas of a proximal platform characterized by stromatoporoid-coral reef complexes (for summary see Calner et al., 2002; Calner and Jeppsson, 2003). Most of the studied specimens of *Discoceras* were removed from their rock matrix, and consequently taphonomic information was lost. A few specimens of *D. lindstroemi* n. sp. were preserved in cephalopod coquina deposited in a reef cavern (Manten, 1971). The coquinas consist of small straight cephalopod shells and rounded fragments of stromatoporoids (Fig. 7.1, 7.4), or a diverse assemblage of cephalopods associated with crinoid, stromatoporoid, and coral fragments (Fig. 7.3, 7.5). It is uncertain whether the coquinas are the result of post-mortem accumulations or represent a natural faunal assemblage. The caverns may have served as hatching grounds (evinced by the accumulation of small shells), a refuge, or hunting grounds for large predators. Based on observations of the collections of the Naturhistoriska Riksmuseet in Stockholm, *Discoceras* was associated with a diverse cephalopod fauna of orthocerids, discosorids, oncocerids, nautilids, and actinocerids, of which only a small part has been described (for summary see Stridsberg, 1985, 1988a; Stridsberg and Turek 1997). These mostly demersal cephalopods were part of a highly diversified reef and peri-reef fauna including corals, stromatoporoids, bryozoans, crinoids, brachiopods, gastropods, trilobites, and other fauna (for overview see Manten, 1971; Laufeld, 1974; Calner et al., 2002). Several shells of *D. lindstroemi* n. sp. are infested with microconchid tubes of *Annuliconchus* sp. (Fig. 6), which are attached to the shell between undulated frills. Apertures of tubes exhibit various orientations, but are usually directed up or down relative to the life orientation of *Discoceras*. Microconchids attached to both organic and inorganic substrates (Taylor and Vinn, 2006). Their occurrence

on both sides of the *Discoceras* shell suggests that in this case microconchids colonized a living animal.

Materials and methods

Morphological characters and measured morphological features are explained in Figure 1. To detail rates of shell expansion, a whorl expansion rate (WER) and revolving index (RI) are used. Whorl expansion rate (WER) is a coiling parameter that is used as defined for gyroconic conchs by Korn and Klug (2003) as the square of the ratio between maximal shell diameter and maximal shell diameter minus aperture height. WER as defined in regularly coiled ammonoids is not used because the smaller shell diameter is not measurable because the impressed zone in *Discoceras* is very shallow, irregular, and usually not visible in the material studied. Nevertheless, the values of WER used here are very close to WER used in regularly coiled ammonoids. Whorl width index (WWI) is used as defined by Korn and Klug (2003) as the ratio of whorl width and whorl height. RI defined by Stridsberg and Turek (1997) is the proportion of whorl height at any point to whorl height of the previous whorl exactly one turn earlier; revolving index is used because it is also applicable to fragmentary shells and has been used in Silurian tarphycerids (Turek and Manda, 2016).

Repositories and institutional abbreviations.—Tarphycerids from Gotland are housed in the Naturhistoriska Riksmuseet, Stockholm, Sweden (prefix RM Mo). Silurian *Discoceras* from Bohemia are deposited in the National Museum, Prague, Czech Republic (prefix NM L).

Systematic paleontology

Abbreviations: chamber length (CL), hyponomic sinus width (HW), hyponomic sinus height (HH), internal whorl height (IWH), maximal shell diameter (SD), whorl height (WH), whorl width (WW); minimum (min.), maximum (max.), median (med.), number of measured specimens (N), standard deviation (σ). Three types of sculpture elements are differentiated: growth lines, ridges, and ribs. Ridges are taller than growth lines; contrary to ribs, ridges are never visible on internal molds, while ribs usually form raised zones on the mold. Lateral furrow and pseudumbilicus, both newly defined terms, are defined below.

Subclass Nautiloidea Agassiz, 1847
Order Tarphycerida Flower, 1950
Family Discoceratidae Dzik, 1984
Genus *Discoceras* Barrande, 1867

Type species.—*Clymenia antiquissima* Eichwald, 1842 from the “Lyckholm stage” (Late Ordovician) of Estonia, by original designation.

Diagnosis.—Tarphycerid with slightly to moderately depressed evolute shell with variously spaced frills; caecum subcentral, shifted toward ventral side, the siphuncle shifts in second and third chambers toward dorsal side, its final dorsal or subdorsal position is attained in third septum.

Remarks.—Previously described Silurian species of *Discoceras* were assigned to *Graftonoceras* Foerste, 1925, which differs from the former genus in having a marginodorsal siphuncle in the third and subsequent chambers. In *Discoceras*, the siphuncle becomes subdorsal when the shell reaches 1.5–2 whorls. Dzik (1984, p. 46) noted that the relatively slight shift of the siphuncle during late ontogeny is a species-specific feature, and synonymized both genera. Stridsberg and Turek (1997) reported differences in the location of the siphuncle in the postembryonic growth stage in two different specimens of a single tarphycerid species—*Ophioceras rudens* Barrande, 1865. *Bateroboceras* Zou, 1983, based on *B. obliquwhorlum* Zou, 1983, from the middle Silurian of Inner Mongolia, exhibits typical features of *Discoceras*; therefore, it too is synonymized with this genus.

Although the shell morphology of tarphycerids is well known, a new morphological feature—lateral furrow—has been observed in two species assigned to *Discoceras*. The furrow appears when the shell attains 0.75–1 whorl. It starts in the ventrolateral or mid-lateral position (Figs. 4.9, 5.1); in the latter case, the furrow shifted abruptly ventrolaterally. The furrow is developed symmetrically on both sides of the shell as a very shallow depression, or sometimes as a light-colored band on the shell (Fig. 5.2). Function of the furrow remains unknown. Pseudumbilicus is herein introduced for the

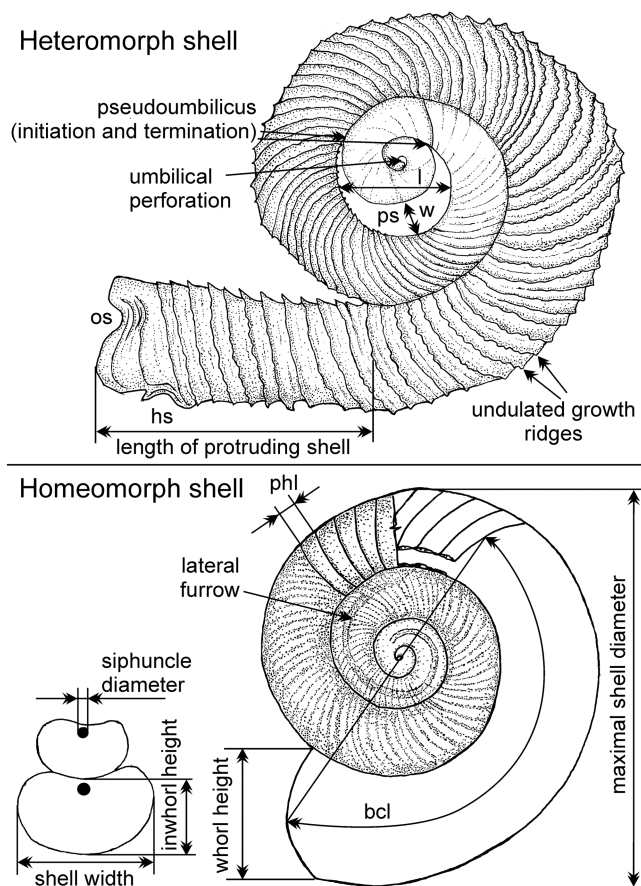


Figure 1. Schematic drawing explaining shell morphology of *Discoceras*. bcl: body chamber length; hs: hyponomic sinus; os: ocular sinus; phl: phragmocone chamber length; w/l: width and length of pseudumbilicus (ps).

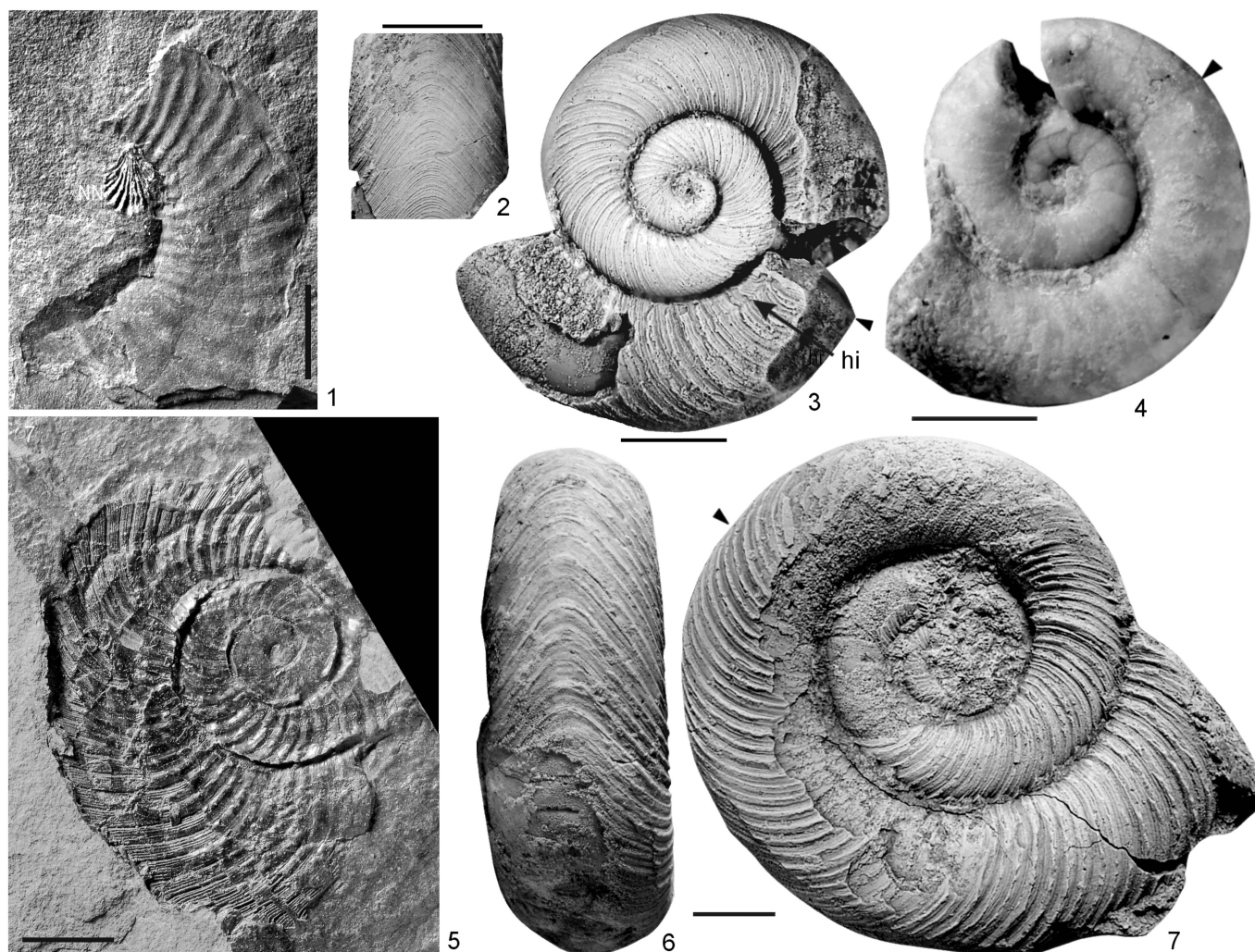


Figure 2. *Discoceras graftonense* (Meek and Worthen, 1870); (1) NM L 42239, lateral view, Malá Chuchle-Vyskočilka, Wenlock, early Sheinwoodian, *Cyrtograptus murchisoni* Biozone, NN indicates brachiopod *Nyorhinx nyobe* (Barrande, 1847) attached on shell surface; (2, 3) RM Mo 59701, ventral and lateral view, Färö, Sandvik, Slite beds, late Sheinwoodian, hi: healed injury; (4) RM Mo 59706, lateral view, Färö, Sandvik, Slite beds, late Sheinwoodian; (5) NM L 42237, lateral view, Koněprusy, Llandovery, Aeronian, *Demirastrites simulans* Biozone; (6, 7) RM Mo 59696, lateral and ventral view, Lärbro, Slite beds, late Sheinwoodian. Specimens in Fig. 2.2, 2.3, 2.6, and 2.7 were coated by ammonium chloride before photographing. Arrow indicates base of body chamber. Scale bars 10 mm.

sickle-shaped cavity between whorls originating due to a brief interval of detached coiling of a shell whorl from preceding whorl (Fig. 1). This cavity corresponds to the secondary umbilical window in early ammonoids (see De Baets et al., 2015).

Discoceras graftonense (Meek and Worthen, 1870)
Figures 2.1–2.7, 13.1, 13.2, 13.4, Table 1

- 1870 *Lituites Graftonensis* Meek and Worthen, p. 51.
1875 *Lituites Graftonensis*; Meek and Worthen, p. 507, pl. 25, fig. 1
1888 *Lituites Graftonensis*; Newell, p. 485.
1894 *Discoceras graftonense*; Hyatt, p. 501.
1904 *Cyclolituites bowningensis* Etheridge, p. 77, pl. 8, figs. 1–6.
1925 *Graftonoceras graftonense*; Foerste, p. 59, pl. 12, figs. 2, 3.
1944 *Graftonoceras graftonense*; Shimer and Shrock, p. 543, pl. 222, figs. 4, 5.

- 1964 *Graftonoceras graftonense*; Furnish and Glenister, p. K360, fig. 258/4.
1983 *Graftonoceras* cf. *griftonense*; Zou, p. 166, pl. 1, figs. 5–7.
1983 *Bateroboceras obliquwhorlum* Zou, p. 172, pl. 1, figs. 1, 2, 8–15.
1984 *Discoceras graftonense*; Dzik, p. 42, fig. 12.

Holotype.—By monotypy, a specimen from the middle Wenlock Port Byron Dolomites of Illinois (Meek and Worthen, 1870, p. 51; figured by Meek and Worthen, 1875, pl. 25, fig. 1).

Diagnosis.—*Discoceras* with marginodorsal siphuncle starting with the third phragmocone chamber; growth ridges separating a set of growth lines appeared at about one-fifth to one whorl.

Occurrence.—Sweden, Gotland, Silurian, Wenlock, latest Sheinwoodian–early Homerian; Färö, Lansa 1: Slite beds, unit d; Färö, Sandvik: Slite beds; Follingbo, Stora Vede 1: Slite beds, unit g; Larbro: upper Slite beds; Othem, Spillings; Slite

Table 1. Dimensions (in millimeters) in *D. graftonense* (Meek and Worthen). N: specimen RM Mo number; SD: maximal shell diameter; WH: whorl height; WW: whorl width; WWI: whorl width index.

N	SD	WW	WH	WWI	WER	RI
59467	28	12.5	7.5	1.67	1.87	—
59696	57.6	18	16.8	1.07	1.99	0.53
59698	55.4	18.8	18.5	1.02	2.25	0.50
59701	37.5	15.6	11	1.42	2.00	0.50
59706	24.2	11.1	8.8	1.26	2.47	0.50
59709	24.2	11.1	7.4	1.5	2.08	0.50
59712	62	—	20	—	2.18	0.48
59714	69	26	19.5	1.33	1.94	0.46
59718	45	—	15	—	2.25	0.53
59775	65.4	23	16.7	1.38	1.80	0.45
59782	39.4	—	11.2	—	1.95	0.50
59801	57.5	—	15	—	1.83	0.54
59803	42.3	16.5	12.3	1.34	1.99	0.41
59839a	60	20	16.9	1.18	1.94	0.47
59839b	66.7	22	20.7	1.06	2.10	0.69

beds, unit g; Othem, Samsuguns 1: Slite beds, unit g. Bohemia; Llandoverly, Aeronian, *Demirastrites triangulatus* Biozone, Solopisky locality, *Demirastrites simulans* Biozone, Koněprusy locality (exact site unknown); Wenlock, Sheinwoodian, *Cyrtograptus purchisoni* Biozone, Praha, Malá Chuchle-Vyskočilka. USA (Foerste, 1925); Australia (Etheridge, 1904; Furnish and Glenister, 1964); and Inner Mongolia (Zou, 1983).

Description.—Shell tightly coiled, exogastric, slightly expanding at most with six whorls; WER: min. = 1.8, med. = 1.99, max. = 2.48, N = 15, σ = 0.18 (Fig. 10). RI: min. = 0.41, med. = 0.5, max. = 0.69, N = 14, σ = 0.06. Umbilical perforation very small, drop shaped, with length ~0.6 mm. Cross section slightly depressed, in later growth stage subcircular or almost subquadrate (WWI: min. = 1.02, med. = 1.33, max. = 1.67, N = 11, σ = 0.19, Fig. 10), impressed zone shallow. Length of body chamber about two-thirds of whorl. Aperture open, or may be slightly contracted in fully grown specimens with uncoiled adapertural part of the shell. Hyponomic sinus V-shaped (HW/HH: min. = 1.54, med. = 1.79, max. = 2.5, N = 6, σ = 0.35; Fig. 10). Phragmocone chambers relatively long (WH/CL: min. = 1.7, med. = 2.71, max. = 4.3, N = 36, σ = 0.69; Fig. 10). First six chambers longer than those succeeding; first chamber cup-like (CL = 2.7 mm, WH = 2.9 mm), length of following chambers measured as abscises along the ventral side (Fig. 1): 3.7 mm (second), 3.6 mm (third to sixth), and 3.1 mm (seventh and following). Depth of septa during shell growth increases from about one-third of chamber length (first to fifth chamber) to about one-half of phragmocone chamber length, and later becomes slightly shorter. Suture oblique, with shallow lateral lobe, ventrally straight. Caecum central, with diameter 0.5 mm, thick-walled; in second phragmocone chamber siphuncle shifted dorsally and in third chamber reached dorsal side. Diameter of siphuncle in apical shell increased to reach constant relative thickness one-fifth of IWH in third chamber. Septal necks hemichoanitic. Connecting rings relatively thick, tubular or slightly vaulted. Sculpture consists of growth lines and transversal growth ridges. Early shell with fine regularly arranged growth lines, densely packed and appear when shell reached one-fifth to one-third of whorl. In following growth stages, growth lines run obliquely to shell axis and form deepening V-shaped hyponomic sinus; middle part of hyponomic sinus rounded. Growth lines in later

shell growth stages irregularly arranged with one or two more distinct growth lines between ridges. Growth ridges distinct laterally as well as ventrally, parallel with growth lines; internal mold usually smooth. Crests of growth ridges sharp. Lateral furrow, observed in a single specimen, is very shallow, ~2 mm wide and accentuated by distinct lighter-colored band on shell. Maximum shell thickness 2 mm.

Three flattened specimens of *D. graftonense* preserved in shale are known from Bohemia. Specimen NM L 42237 (Fig. 2.5) is a damaged, moderately expanding shell (SD = 56 mm, three and one-quarter whorls); adapertural part of the shell missing. Shell tightly coiled, impressed zone very shallow, umbilical perforation small. Sculpture consists of prominent growth ridges and fine growth lines. Growth ridges, discernible from the beginning of the second whorl, appear as curved narrow elevations running obliquely to the shell axis. They are densely spaced, moderately vaulted to the aperture laterally, indicating a deep hyponomic sinus ventrally. Growth lines are prominent, regularly arranged, and parallel with growth ridges; 7–10 lines between crests of adjacent growth ridges. The second specimen of *D. graftonense* (NM L 4223) is a flattened internal mold of a body chamber (Fig. 2.1). The third specimen (NM L 46556; not figured) is a flattened fragment of the ventral side showing ribs forming a deep hyponomic sinus.

Materials.—17 specimens from Gotland: RM Mo 59467, 59505, 59696, 59698, 59701, 59706, 59709, 59712, 59714, 59718, 59775, 59782, 59801, 59803, 59806, 59839, 155975; three specimens from Bohemia: NM L 42237, L 42239, L 46556.

Remarks.—Specimens from Gotland and Bohemia are very similar in the shape of the shell, hyponomic sinus, course of growth lines, position and thickness of siphuncle, sutures, and length of phragmocone chambers to *D. graftonense* from Ohio and Illinois. Two type specimens figured by Foerste (1925, pl. 12, figs. 2, 3) fall within the variability of WER and WWI described herein (Fig. 10). Four endemic species diverged from widely distributed *D. graftonense* (shell morphology of first whorl in all these species and principal conch parameter [WER] are very similar with *D. graftonense*): *D. ortonii* (Meek, 1873) from Ohio differs in having a shallow hyponomic sinus (HW/HH = 2.7–3, which is a higher value than those in *D. graftonense* from Gotland); *D. amissus* (Barrande, 1865); *D. stridsbergi* n. sp., which has distinct ribs; and *D. lindstroemi* n. sp., which has undulated frills, a less-depressed cross section, and a heteromorphic shell.

Discoceras stridsbergi new species
Figure 3, Table 2

Type specimens.—Holotype, an almost complete shell RM 59777 (Fig. 3.8), from Stora Vede (Gotland), Slite Beds, early Homeric, Wenlock, Silurian. Paratypes: RM 59452 (Fig. 3.7), 59463 (Fig. 3.6), 59457 (Fig. 3.5), 59719 (Fig. 3.1), 59786 (Fig. 3.4), 59836 (Fig. 3.3), and 59948 (Fig. 3.2).

Diagnosis.—*Discoceras* with moderately expanding shell and recurrent sharp ribs around the shell, running obliquely to shell

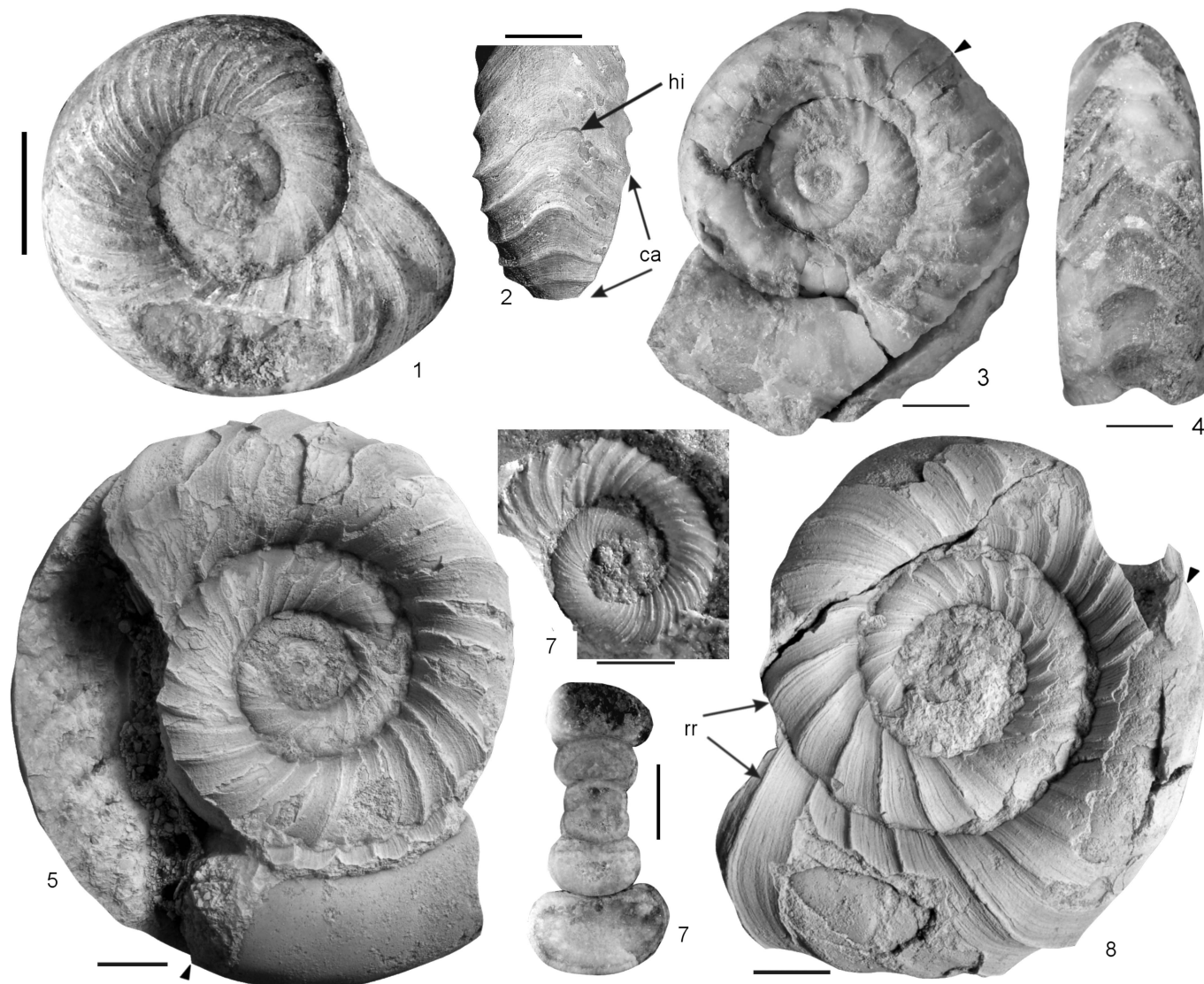


Figure 3. *Discoceras stridsbergi* n. sp. from Slite Beds of Gotland. (1) RM Mo 59719, lateral view, Färö, Sandvik, late Sheinwoodian; (2) RM Mo 59948, ventral view to adapertural region, Storugns, early Homerian (hi: healed injury, ca: contracted aperture); (3) RM Mo 59836, lateral view, Follingbo, Stora Vede 1, early Homerian; (4) RM Mo 59786, ventral view, Follingbo, Stora Vede 1, early Homerian; (5) RM Mo 59457, lateral view, Follingbo, Stora Vede 1, early Homerian; (6) RM Mo 59463, lateral view, Färö, Sandvik, late Sheinwoodian; (7) RM Mo 59452, cross section Follingbo, Stora Vede 1, early Homerian (gr: growth ridges); (8) RM Mo 59777, lateral view, Follingbo, Stora Vede 1, early Homerian. Specimens in Figure 3.5 and 3.8 were coated with ammonium chloride before photographing. Arrows in Figure 3.3, 3.5, and 3.8 indicate base of body chamber. Scale bars 10 mm.

axis laterally and forming prominent and deep hyponomic sinus ventrally.

Occurrence.—Stora Vede 1, Larbo, Storugns, and Follingbo, Slite Beds, early Homerian, Wenlock.

Description.—Evolute, exogastric shell with at most four whorls (WER: min. = 1.7, med. = 2.02, max. = 2.66, N = 24, σ = 0.20; Fig. 10). RI: min. = 0.43, med. = 0.48, max. = 0.73, N = 16, σ = 0.08. Cross section of early whorls strongly depressed, later moderately depressed to subcircular (WWI: min. = 1.03, med. = 1.37, max. = 1.68, N = 26, σ = 0.16; Fig. 10). Length of body chamber slightly less than one-half of whorl. Aperture open or slightly contracted. Hyponomic sinus V-shaped (HW/HH: min. = 1.79, med. = 1.9, max. = 2.4, N = 9, σ = 0.21; Fig. 10).

Phragmocone chambers relatively long (WH/CL: min. = 2.11, med. = 2.73, max. = 3.78, N = 37, σ = 0.44; Fig. 10). Septa slightly vaulted, suture oblique with shallow lateral lobes. Siphuncle relatively thin (one-ninth of IWH). Distinct growth lines running obliquely to shell axis are laterally straight or slightly undulate. Ventral lobe (= hyponomic sinus) deep, V-shaped, with rounded apex. Slightly indicated ribs in form of narrow ridges appear at beginning of second whorl. Their course is parallel with growth lines, and they are well developed at the end of the second whorl. Height of ribs increases with shell growth, but only on the last whorl do the ribs become visible on the internal mold as a faintly developed raised zone. Distance between ribs equal to phragmocone chamber length; their density is 6–8 per quarter whorl. Ribs asymmetric in cross section, adapertural part of ribs much steeper than adapertural part. Maximum shell thickness 1.2 mm.

Table 2. Dimensions (in millimeters) in *D. stridsbergi* n. sp. For abbreviations see Table 1.

N	SD	WW	WH	WWI	WER	RI
59447	61.8	24.7	16.8	1.47	1.89	—
59448	51.4	21.3	16.3	1.31	2.14	—
59452	—	20.6	17.8	1.16	—	0.73
59453	66.5	26.8	19.2	1.4	1.98	—
59456	49.6	21.3	15.6	1.37	2.13	—
59457	65.4	22.2	18.2	1.22	1.92	0.6
59458	53.7	21.8	15.2	1.43	1.95	—
59459	62.9	22.5	14.6	1.54	1.70	—
59699	32	15.3	9.8	1.56	2.08	—
59719	37	24	14.3	1.68	2.66	0.44
59764	67.2	23.8	18	1.32	1.87	0.6
59768	61.5	22.9	16.3	1.41	1.86	—
59771	58	19	16.9	1.12	1.99	0.5
59772	50.6	19.2	13.2	1.46	1.83	0.47
59773	33.5	15.9	10.7	1.49	2.16	0.44
59777	67.5	22.1	18.8	1.18	1.92	0.52
59780	—	21.8	17.8	1.22	—	0.44
59781	46.2	17.7	13.9	1.27	2.05	0.46
59786	57	22	18	1.22	2.14	0.58
59806	42.5	22.5	15	1.5	2.39	0.48
59836	54	24	16.3	1.47	2.05	0.46
59840	37.3	15.5	10.6	1.46	1.95	—
59905	44.5	15.2	14.2	1.07	2.16	—
59906	48	16	15.6	1.03	2.20	—
59948	41.6	16.6	13.4	1.24	2.18	0.43
157058	42.3	16.5	12.1	1.36	1.96	—

Etymology.—Specific name *stridsbergi* after Sven Stridsberg (Lund University, Sweden), who has made significant contributions to our knowledge of Silurian cephalopods.

Materials.—28 specimens, including types (Table 2).

Remarks.—*Discoceras stridsbergi* n. sp. most probably diverged from *D. graftonense* because both species have a similar early shell morphology. Shell shape in both species is similar, but *D. stridsbergi* n. sp. differs in having a shorter body chamber, sharp ribs around the shell, and a wider hyponomic sinus. Ribs evolved from growth ridges in early whorls, ridges here are much lower and symmetrical in cross section.

Discoceras lindstroemi new species
Figures 4–7, 12, 13.3, 13.5, Table 3

Type specimens.—Holotype, a complete shell RM Mo 59511 (Fig. 4.7) from Othem, Kvarnberget 1, Gotland, Slite Beds, late Sheinwoodian, Wenlock, Silurian. Paratypes: RM Mo 40265 (Figs. 4.4, 7.4), 59420 (Figs. 4.1, 12.3), 59432 (Fig. 4.3), 59427 (Fig. 4.6, 4.14), 59442 (Fig. 4.2), 59505 (Fig. 6.1), 157592 (Fig. 4.8), 59506 (Fig. 4.10), 59887 (Fig. 4.13), 59416 (Fig. 4.15), 59419 (Fig. 5.2), 59884 (Fig. 13.3), 59422 (Fig. 5.1), 162186 (Fig. 6.3), and 181926 (Fig. 2.11) from Slite Beds, late Sheinwoodian, Wenlock.

Diagnosis.—*Discoceras* with heteromorph shell, pseudoumbilicus and undulate growth frills.

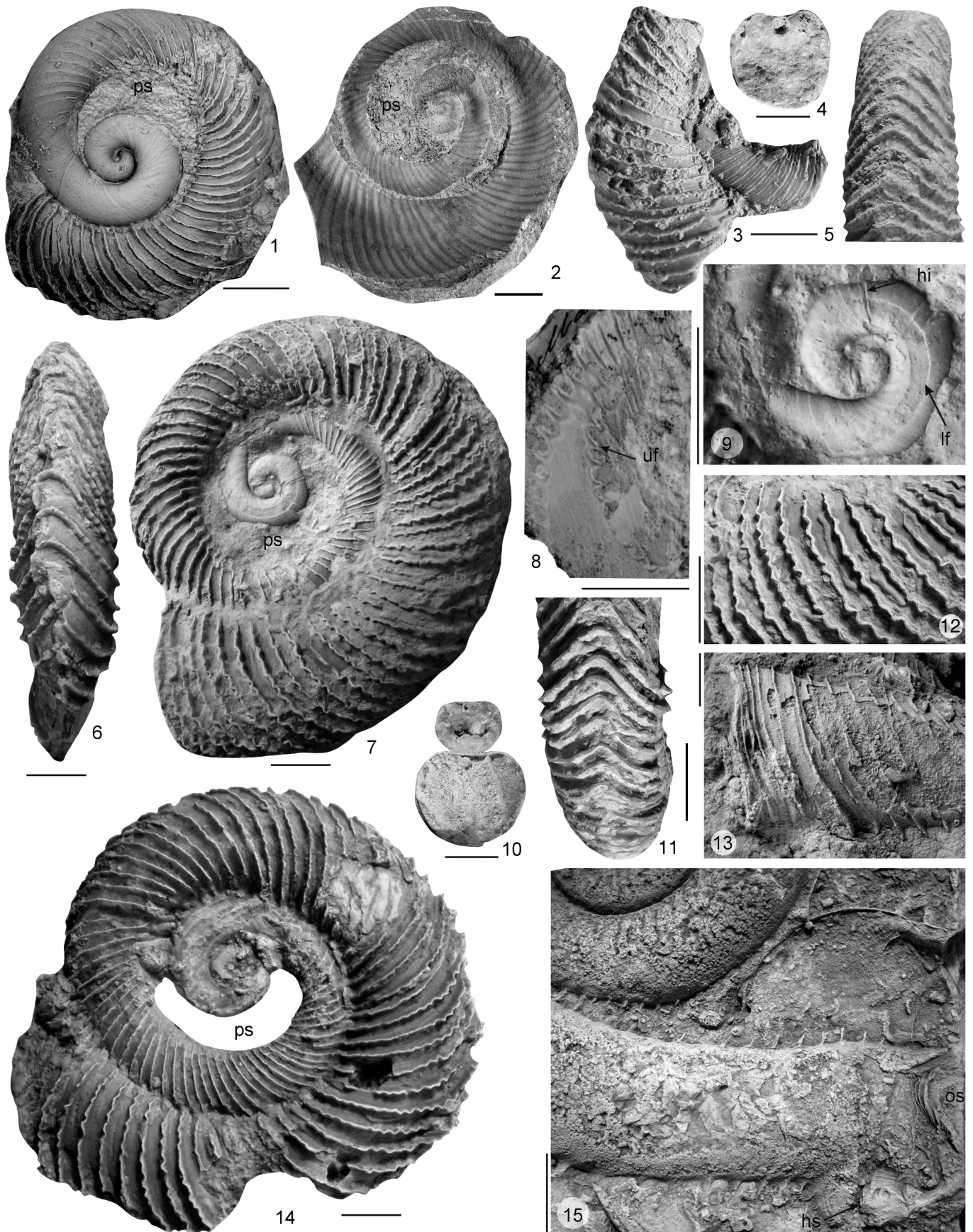
Occurrence.—Silurian, Wenlock, latest Sheinwoodian–early Homerian; Slite beds; Gotland; Färö, Länaberget 1; Färö, Lanså 1; Föllingbo, Stora Vede 1; Othem, Spillings; Othem, Samsuguns 1; Slite, Kvarnberget 1.

Description.—Evolutive, exogastric, slightly depressed shell with a maximum of three whorls (WER: min. = 1.83, med. = 1.98, max. = 2.46, N = 19, σ = 0.20; Fig. 10). RI: min. = 0.33, med. = 0.47, max. = 0.6, N = 16, σ = 0.07. Shell heteromorphic: early shell tightly coiled (0.5–1.5 whorl), following part of the shell (about one-half of whorl, Table 3) coiled in loose spire, followed by shell tightly coiled; adapertural part of shell in adult specimens straight. Indication of an additional enrolling, observed in a single specimen (Fig. 13.3), appeared at end of second whorl. The shell before the end of the second whorl is here tightly coiled, in the following one-quarter whorl, a very narrow window between the external shell walls of two neighboring whorls appears. Only distal ends of growth ridges oriented adapically are here in contact with the last whorl (body chamber). Umbilical perforation very small. Cross section slightly depressed, impressed zone shallow (WWI: min. = 0.73, med. = 0.96, max. = 1.18, N = 17, σ = 0.10; Fig. 10). Length of body chamber equal to two-thirds of whorl. Aperture open, in fully grown shells with protruding adapertural part of last whorl, aperture sometimes slightly contracted.

Hyponomic sinus V-shaped (HW/HH: min. = 1.6, med. = 2.14, max. = 3.17, N = 11, σ = 0.5; Fig. 10). Ocular sinus well developed, situated close to dorsal side (Fig. 4.15). Phragmocone chambers relatively long (WH/CL: min. = 1.7, med. = 2.71, max. = 4.43, N = 36, σ = 0.69; Fig. 10). In early shell (Fig. 12.2), first five chambers are longer than succeeding chambers. Initial chamber cup-like, with maximum length 1 mm, dorsoventral distance in first septum is 2.9 mm, length of following chambers measured on ventral side ~2.9 mm (second to fourth), following chambers are shorter, length ~1.5 mm (Fig. 13.5). Depth of septa during shell growth increases from about one-third of chamber length (second to fourth chamber) to about one-half of phragmocone chamber length in following chambers. Suture forms oblique angle with shallow lateral lobe, but is straight ventrally.

Caecum subcentral, slightly expanded within first chamber (diameter 0.5 mm). In second and third chamber, siphuncle shifted toward dorsal side and diameter markedly increased. In the third chamber, siphuncle reaches dorsal side and diameter is about one-sixth of whorl height. Septal necks hemichoanitic; connecting rings relatively thick, tubular or slightly vaulted. Episeptal cameral deposits slightly developed close to siphuncle. First half of whorl smooth or with fine irregular growth lines; growth ridges, accompanied by small constriction of shell, are developed in some specimens at beginning of pseudoumbilicus (WH = 5.2–7.2 mm; Fig. 12.3); irregularly situated growth ridges appeared between WH 4–7.4 mm, they

Figure 4. *Discoceras lindstroemi* n. sp. from Slite beds, late Sheinwoodian, Gotland; ps: pseudoumbilicus. (1) RM Mo 59420, lateral view, Othem, Spillings; Slite beds; (2) RM Mo 59442, lateral view, counterpart, Färö; (3, 5), RM Mo 59432, lateral and ventral view, Länaberget; (4) RM Mo 590265, cross section, Othem, Samsuguns 1; (6, 14) RM Mo 59427, ventral and lateral view, ps: pseudoumbilicus, Othem, Samsuguns 1; (7, 9, 12) RM Mo 59511, (7) lateral view, (9) detail of early shell showing lateral furrow (lf) and a healed injury (hi), (12) detail of sculpture, Othem, Samsugun; (8) RM Mo 157592, lateral view, polished specimen with undulated frills (uf), Martebo; (10) RM Mo 509506, cross section, Othem, Spillings; (11) RM 181926 ventral view to adapertural region, Färö; (13) RM Mo 59887, lateral view to adapertural region, ocular sinus slightly indicated, Othem, Samsuguns 1; (15) RM Mo 59416, lateral view, protruding last whorl with contracted aperture, Othem, Samsuguns 1, os: ocular sinus, hs: hyponomic sinus. All specimens except Figure 4.8 were coated with ammonium chloride before photographing. Scale bars 10 mm.



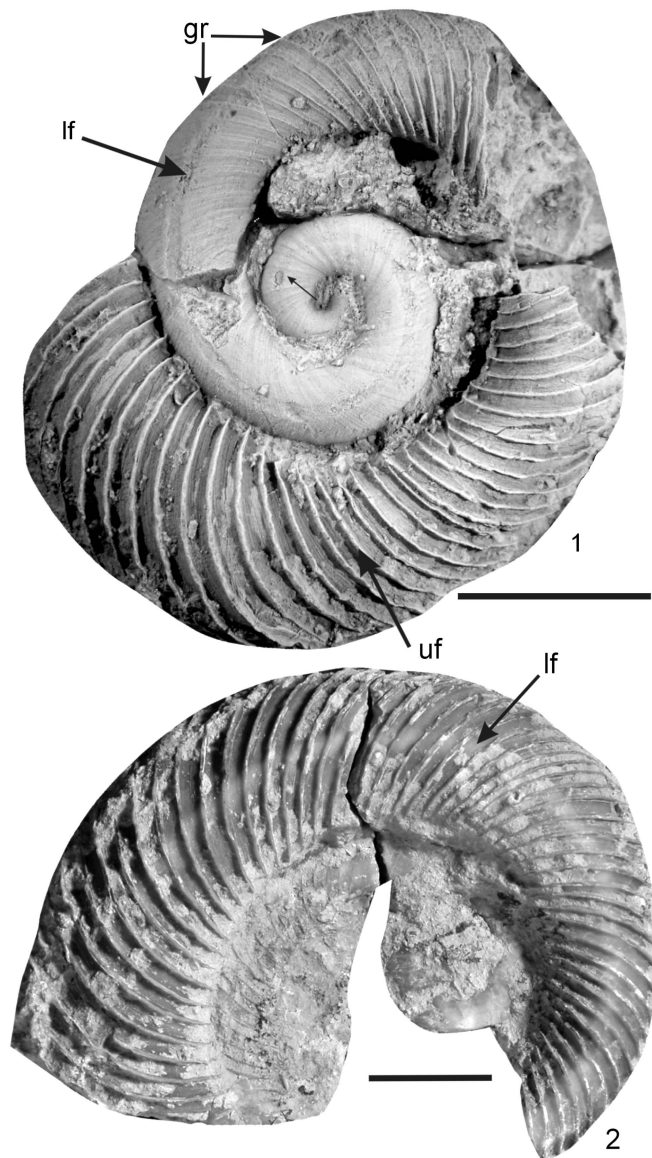


Figure 5. Lateral furrow (lf) in *Discoceras lindstroemi* n. sp., Othem, Klints, Slite beds, late Sheinwoodian. (1) RM Mo 59422, lateral view (gr: growth ridges, uf: undulated frill); (2) RM Mo 59419, lateral view. Specimen in Figure 5.1 was coated with ammonium chloride before photographing. Scale bars 10 mm.

pass rapidly into regularly distributed growth ridges and finally into regular undulated frills near end of pseudoumbilicus (Figs. 4, 5.1). Height of frills continually increases up to 3 mm with shell growth (Figs. 4.8, 4.12, 7.2). Distance of frills increases from 1 mm to 2.5 mm. In fully grown specimens, the frills at the aperture are densely packed, and a shallow ocular sinus appeared (Fig. 4.13). Frills are intercalated with irregularly distributed growth lines. Transverse sculptural elements are oblique to the shell axis, with shallow lateral saddles. Maximum shell thickness is 3 mm.

In four specimens, lateral furrows are developed symmetrically in both sides of the shell. The furrows form shallow depressions or light-colored bands on shell; width of furrows equals one-quarter of whorl height. Furrows appear on shell reaching three-quarters of a whorl (Figs. 4.9, 5). They start in the

mid-lateral part of the shell, and rapidly shift toward the venter to attain a ventrolateral position. The furrows fade away with the appearance of distinct growth ridges or continue as a light-colored band.

The sculpture in a single incomplete counterpart from Färö (Figs. 4.2, 12.1) differs from the sculpture of specimens from Othem. Distinct widely spaced growth ridges with one or two inserted fine growth lines appear in one-half of the first whorl, where the ventral side of the shell lost contact with the following whorl. Subsequently, growth ridges become laterally slightly vaulted to the aperture, and subordinate growth lines are more distinct. A prominent repair is located dorsolaterally at the beginning of the second whorl.

Etymology.—Specific name, *lindstroemi*, after Gustav Lindström, who has made significant contributions to the understanding of the Silurian fauna of Gotland.

Materials.—In addition to the types, these nine specimens: RM Mo 59418, 59425, 59426, 59428, 59443, 59499, 59504, 158752, 162186.

Remarks.—*Discoceras lindstroemi* n. sp. exhibits considerable variability in the shape of the hyponomic sinus. In comparison with coeval species, the cross section of the shell is less depressed (Fig. 10). In addition, the protruding final whorls in *D. lindstroemi* n. sp. have a contracted aperture with well-developed ocular sinuses. Similar mature modification is known in the contemporaneous tarphycerid *Ophioceras* Barrande, 1865 and Ordovician lituitids (Furnish and Glenister, 1964).

Discoceras sp. indet.

Figure 8

Description.—Incomplete specimen preserved as a weathered internal mold in gray dolomite. Specimen comprises at least two and one-half whorls, adapertural part of the shell strongly ventrally damaged. Shell evolute, gradually expanding, exogastric, RI is 0.68. Complete reconstructed shell reached four whorls (SD = 90 mm). First half of second whorl coiled in loose spire, followed by tightly coiled whorls. Cross section subcircular, impressed zone not developed. Prominent regularly arranged ribs, no substantial changes in density of ribs during growth (10 ribs per quarter whorl). Hyponomic sinus deep and wide.

Occurrence.—Middle Wenlock of Gotland, Sweden.

Material.—A single specimen RM Mo 59972.

Remarks.—In mode of coiling, *Discoceras* sp. indet. resembles *D. lindstroemi* n. sp.; nevertheless, straightening of the final whorl is not documented. Ribbing in *Discoceras* sp. indet. is similar to that in *D. stridsbergi* n. sp. In comparison with later species, the shell is more gradually expanding (WER is not measurable due to preservation, but RI is higher than in *D. lindstroemi* n. sp.), ribs on internal molds are more pronounced and more densely spaced, and cross section is sub-circular. Because only one specimen of *Discoceras* sp. indet. was available for study, it is left in open nomenclature.

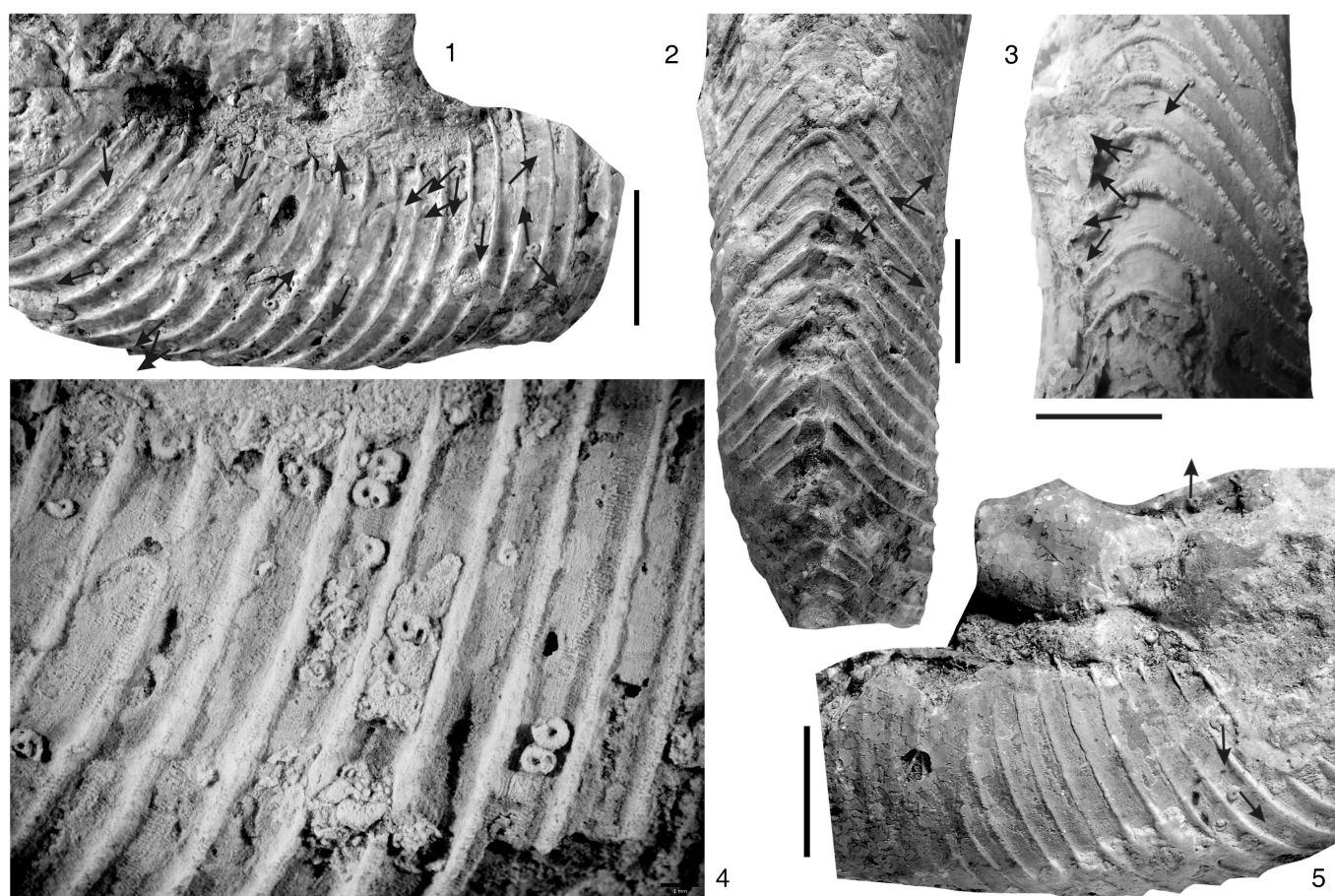


Figure 6. Epibionts in *Discoceras lindstroemi* n. sp., Othem, Klinti, Slite beds, late Sheinwoodian. (1, 2, 4, 5) RM Mo 59505, right and left lateral (1, 5), ventral (2) views and detail of microconchid tubes in lateral shell side (4); (3) RM Mo 162186, ventral view to adapertural region. Arrows indicated aperture orientation of tubes. All specimens were coated by ammonium chloride before photographing. Scale bars 5 mm (1–3, 5) and 1 mm (4).

Discoceras amissus (Barrande, 1865)

Figure 9

1865 *Lituites* (*Ophioceras*) *amissus* Barrande, pl. 45, figs. 26, 27.

1867 *Ophioceras* (*Lituites*) *amissus*; Barrande, p. 182.

Holotype.—Holotype by monotypy, NM L 9107, illustrated by Barrande (1865) on pl. 45, figs. 26, 27 from Králův Dvůr, Ovčinec, Bohemia; Landoverý, Aeronian, *Demirastrites convolutus* Biozone.

Diagnosis.—*Discoceras* with combination of distinct growth lines and prominent ribs, symmetrical in cross section.

Description.—Fragment of shell consists of one and three-quarter whorls, SD = 30 mm, WER = 2.07. Adapertural part of shell missing. Shell evolute, gradually expanding, exogastric, cross section unknown, due to flattening in shale. First half of first whorl smooth, almost equally developed growth lines appeared on second half of first whorl, forming deep and wide hyponomic sinus. Near beginning of second whorl, narrow crests appear, passing gradually into prominent wide ribs running obliquely to shell axis. Laterally, ribs slightly vaulted adaperturally; ventrally, they pass into the deep hyponomic sinus. Top parts of ribs rounded. Prominent growth lines

inserted between ribs. On last quarter of shell, nine ribs are visible. Growth lines regularly arranged, parallel with ribs, 5–11 in number between crests of adjacent ribs.

Material.—Holotype only.

Remarks.—*Discoceras amissus* resembles *D. stridsbergi* n. sp., differing in having fewer more distinct ribs that fade out towards the adapertural end of the shell; ribs are symmetrically rounded in cross section.

Distribution of *Discoceras*

More than 30 species of the widely distributed genus *Discoceras* have been described from Middle and Upper Ordovician strata. *Discoceras* has been reported from Australia (Teichert and Glenister, 1954); Canada (Whiteaves, 1897); Denmark (Rasmussen and Surlyk, 2012); Estonia (Balashov, 1953; Stumbur 1962); Kazakhstan (Barskov, 1972); northwest China (Lai and Wang, 1986; Lai, 1987); Norway (Strand, 1934; Sweet, 1958); Sweden (Kröger et al., 2011; Kröger, 2013); Tibet (Lai, 1982); and Wisconsin (Whitfield, 1882; Hyatt, 1894). *Discoceras* has also been found in erratic boulders of Baltoscandian origin (Lossen, 1860; Roemer, 1861; Remelé, 1890; Hyatt, 1894; Dzik, 1984). *Discoceras* is globally dispersed

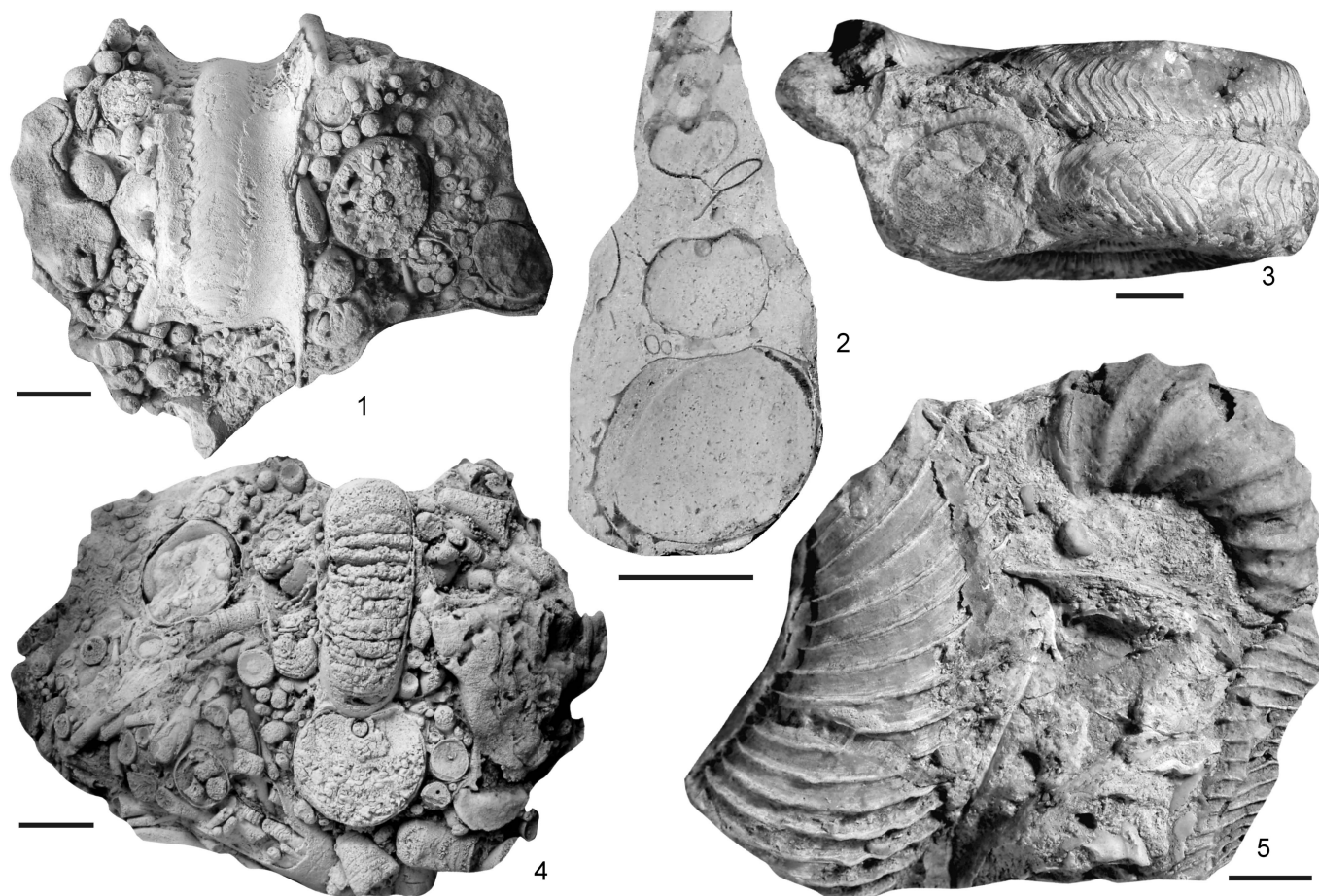


Figure 7. *Discoceras lindstroemi* n. sp. preserved in reef cephalopod coquinas, late Sheinwoodian, Slite Beds, Gotland; (1, 4) RM Mo 40265, Othem, Spillings; (2) Polished slab with cross section of *D. lindstroemi*, RM Mo 59882, Othem, Spillings; (3) Two aggregated specimens, one from anterior and one from posterior view, RM Mo 59891, Othem, Samsuguns 1; (5) *Lechritrochoceras* sp. and two fragments of *D. lindstroemi*, RM Mo 59891, Othem, Samsuguns 1. Specimens in Figure 7.1 and 7.3–7.5 were coated by ammonium chloride before photographing. Scale bars 10 mm.

Table 3. Dimensions (in millimeters) and aperture characteristic of selected specimens of *D. lindstroemi* n. sp. SD1: shell diameter at beginning pseudoubilicus; SD2: shell diameter at termination of pseudoubilicus; PL: pseudoubilicus length; PW: pseudoubilicus width; LPS: length of protruding shell (ND: not developed); A: Aperture (O: open, C: contracted); for other abbreviations, see Table 1.

N	SD	WW	WH	WWI	WER	RI	SD1	SD2	PL	PW	LPS	A
59416	64.7	19	18.7	1.02	1.98	0.54	14	35.2	17	4	46.7	C
59419	39	13	13.5	0.96	2.34	—	14.7	24.3	—	—	—	O
59420	39	—	13	—	2.25	0.53	15	24	14.4	6	ND	O
59422	37.3	—	10	—	1.87	0.44	16	22.2	16.1	4.2	ND	O
59425	65.4	19.2	22	0.87	2.27	0.5	14.8	29.7	—	4.8	—	—
59426	69.3	18	18.8	0.96	1.88	—	16.9	—	—	—	—	—
59427	60	14	16.7	0.84	1.92	0.6	17.5	41.6	20.8	5.6	ND	O
59428	52.9	16	18	0.89	2.30	0.46	—	45.4	20	6.2	ND	O
59442	51	—	14.5	—	1.95	0.36	13.2	46	26.4	8.5	10	O
59443	67.7	17.8	18.2	0.98	1.87	0.47	—	32.2	17.4	4.7	ND	O
59499	53.8	18	15.2	1.18	1.94	—	13.7	37.3	—	—	—	—
59504	56.7	16.5	16.3	1.01	1.97	0.48	—	—	—	—	—	—
59511	69.6	17.7	22.7	0.78	2.2	0.33	15.8	46.19	22.8	5.2	ND	O
59877	—	21.8	24.4	0.89	—	0.44	—	—	—	—	29	O
59884	59	16	18.6	0.86	2.13	—	—	24	—	—	—	—
59891a	64	21.3	21	1.01	2.22	0.53	—	—	—	—	—	O
59891b	67	20.3	24.3	0.84	2.46	0.44	—	—	—	—	—	C
155975	54.1	18.3	19	0.96	2.38	0.37	11.9	37.2	—	—	—	—
158752	49.5	12.4	12.9	0.96	1.83	0.4	18.8	36.4	22.5	5.5	—	—
181926	72.5	14	19.1	0.73	1.84	0.55	19.3	—	27.5	—	ND	O

in a belt between 20°N and 50°S when the Ordovician reconstruction of Torsvik and Cocks (2013) is used. It is known from the Baltica, North China, NE Gondwana, Kazakhstania, Laurentia, and Tibet paleocontinents. *Discoceras* inhabited

carbonate platforms, in shallows as well as in their deeper parts up to a depth of ~100 m. Some species are also known from offshore shales deposited under dysoxic/anoxic conditions (Kröger et al., 2009, 2011; Kröger and Ebbestad, 2014).

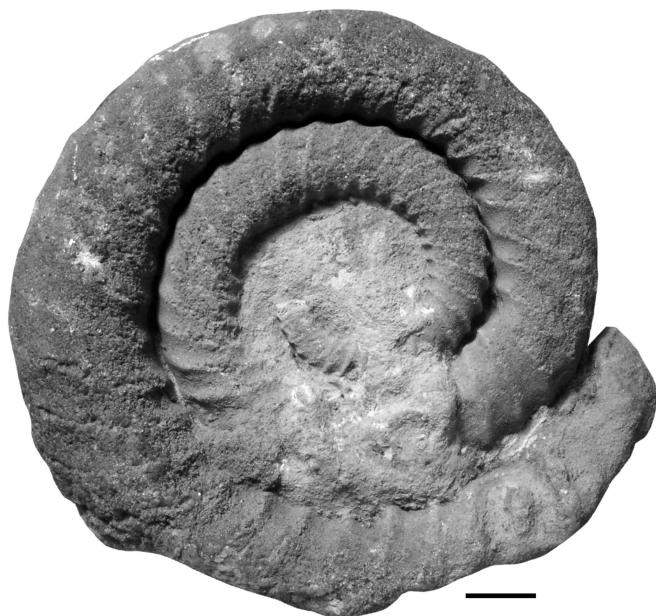


Figure 8. *Discoceras* sp. indet., RM Mo 59972, lateral view, Gotland, locality unknown. Scale bar 10 mm.

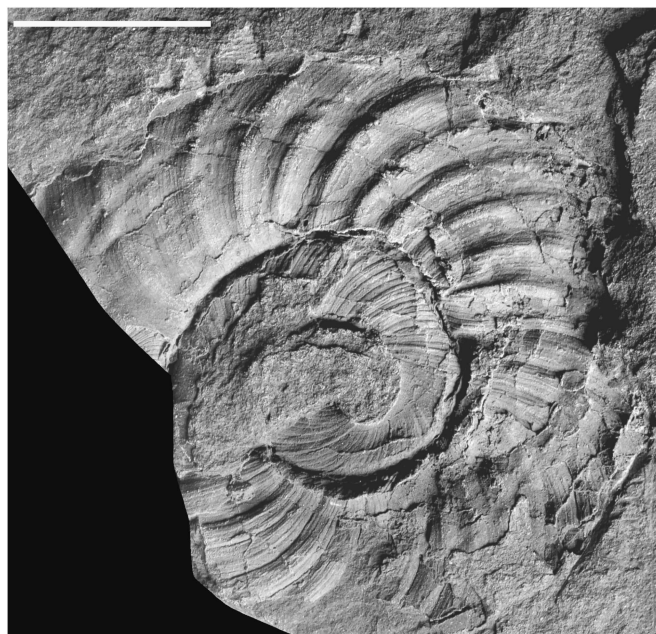


Figure 9. *Discoceras amissus* (Barrande, 1865), NM L 9107, holotype, lateral view, specimen flattened in shale, Bohemia, Beroun-Králův Dvůr, Ověčinec, Llandovery, Aeronian, *Demirastrites convolutus* Biozone. Scale bar 10 mm.

Contrary to the worldwide dispersion of tarphycerids during the Ordovician, *Discoceras* and *Trocholites* (the latter represented by a single species in the Llandovery of Great Britain; Holland, 2010) are the only tarphycerid genera known to have survived the late-Ordovician extinction event. This event was accompanied by a marked decrease in nautiloid diversity (Frey et al., 2004; Kröger and Zhang, 2009). *Discoceras graftonense* is a widely distributed species, ranging from the Aeronian (Llandovery) to early Homerian (Wenlock). Other Silurian species of *Discoceras* were endemic, with very limited distribution even in one basin (Table 4). The widespread geographic distribution of *D. graftonense* and the origin of four species of *Discoceras* in the latest Sheinwoodian and early Homerian represented the last weak diversification and dispersion event of tarphycerids, terminated by the mid-Homerian extinction (Mulde and Lundgreni events; Calner, 2008). Only a single tarphycerid, *Ophioceras* Barrande, 1865, survived this extinction, including two long-ranging species; the youngest of these became extinct just below Silurian-Devonian boundary (Stridsberg and Turek, 1997; Turek and Manda, 2016). No tarphycerid species originated following the middle Homerian; there is no evidence of a splitting lineage or a reason to split a lineage into a series of successive chronospecies. *Ophioceras* co-occurs with *Discoceras* in Gotland, North America, and North China. While *Discoceras* retained shell morphology similar to Ordovician tarphycerids, the shell of *Ophioceras* acquired a thin siphuncle and a higher expansion rate of the shell (Stridsberg and Turek, 1997: revolving index 1.3–1.7; cf. RI of Silurian *Discoceras* 0.33–0.73).

The low diversity and extinction of tarphycerids in the Silurian, together with their relatively thick marginal siphuncle, is consistent with a macroevolutionary trend toward reduction of siphuncle diameter and thinning of connecting rings in Paleozoic cephalopods (Kröger, 2003). It is also consistent with an exceptionally low adaptive pressure toward planispirally coiled shells during the Silurian and the tendency towards a coiled shell with a central siphuncle (Kröger, 2005).

In the Wenlock of the Baltic Basin, *Discoceras* inhabited a shallow water environment close to reefs and its shells also occur in reef caverns. *Discoceras* also inhabited shallow water platforms in proximity to reefs in the Midwestern craton in Laurentia (Shaver, 1991) and the North China plate (Li et al., 1983). A single occurrence was documented from deeper water shales with a pioneer benthic community in the Prague Basin (Turek, 1983). The geographic distribution of *Discoceras* in the Silurian, in contrast to the wide distribution of

Table 4. Distribution of *Discoceras* species in the Silurian. PG indicates peri-Gondwana.

Species	Paleogeography	Age	Environment
<i>D. graftonense</i> (Meek and Worthen, 1870)	Northeast China Plate Prague Basin, Perunica, PG Baltica	Late Sheinwoodian–early Homerian, Wenlock Aeronian, Llandovery; earlier Sheinwoodian, Wenlock Late Sheinwoodian–early Homerian, Wenlock	Carbonate platform, proximal Off-shore black shale Carbonate platform
<i>D. amissus</i> (Barrande, 1865)	Northeast Gondwana Prague Basin, Perunica, PG	Late Sheinwoodian–early Homerian, Wenlock Aeronian	An limestone Off-shore black shale
<i>D. ortoni</i> (Meek, 1873)	Laurentia, Midwestern craton	Late Sheinwoodian–early Homerian, Wenlock	Carbonate platform, proximal
<i>D. stridsbergi</i> n. sp.	Baltica	Early Homerian, Wenlock	Carbonate platform, proximal
<i>D. lindstromi</i> n. sp.	Baltica	Late Sheinwoodian–early Homerian, Wenlock	Carbonate platform, proximal
<i>D. sp. indet.</i> n. sp.	Baltica	Late Sheinwoodian–early Homerian, Wenlock	Carbonate platform, proximal

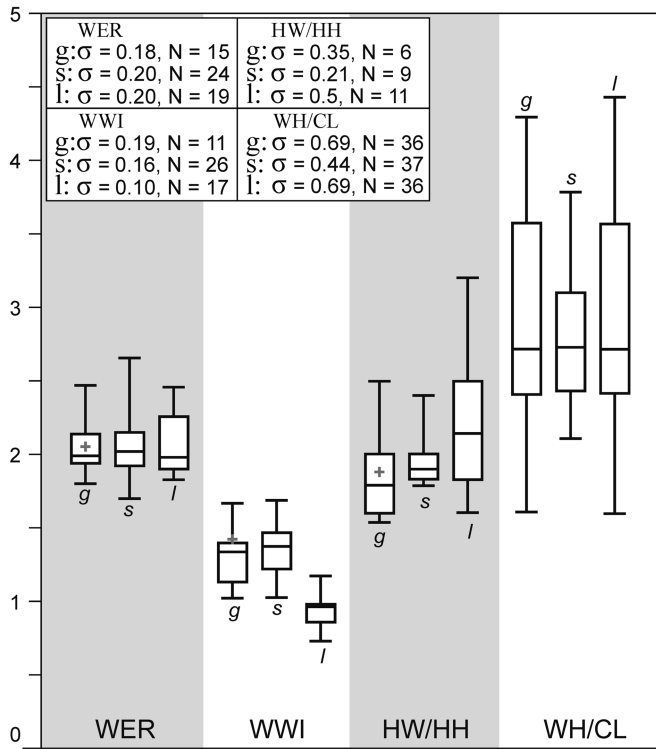


Figure 10. Boxes (minimum, maximum, median, and first and second quartile) showing WER, WWI, HW/HH ratio and WH/CL ratio in Silurian *Discoceras* from Gotland; g: *D. graftonense*, s: *D. stridsbergi*, l: *D. lindstroemi*. Gray cross indicates value in holotype of *G. graftonense* (data from Foerste, 1925).

Ordovician tarphycerid genera, is restricted to the proximal areas of low-latitude platforms and occasionally more distal areas of black shale sedimentation between 20°N and 30°S (Table 4, Fig. 11).

The occurrence of *Discoceras* in Llandovery off-shore black graptolitic shales of the Prague Basin (Perunica microplate, peri-Gondwana), which were deposited under anoxic near-bottom conditions (Štorch, 2006) well removed from a carbonate platform, resembles Ordovician occurrences in black graptolite shales of Baltoscandia (Kröger et al., 2009; Rasmussen and Surlyk, 2012). The earlier Silurian anoxic episode restricted almost all fauna except graptolites in peri-Gondwanan basins. Nautiloid immigration from low latitudes into peri-Gondwanan basins with isolated elevated submarine regions is linked with a decline in early Silurian anoxia, and activation of currents beginning in the middle Llandovery to Wenlock (Stridsberg, 1988b; Manda, 2008b; Histon, 2012; Evans et al., 2015; Fig. 11).

Pioneer nautiloid immigrants were forms with a coiled shell. Aeronian *Discoceras* is the earliest known stray immigrant into the Prague Basin; later, in the latest Llandovery, the discosorid *Phragmoceras* Broderip, 1839 in Murchison (1839) appeared. Permanent nautiloid populations and continuous faunal exchange in the Prague Basin appeared in the late Wenlock (Manda, 2008b). The occurrence of *Discoceras* in black off-shore shales in association with graptolites indicates its migration potential and swimming ability in surface currents across an open sea. Dispersion potential also could have been enhanced for possible planktotrophic juveniles transported by ocean currents. This also explains the dispersion of

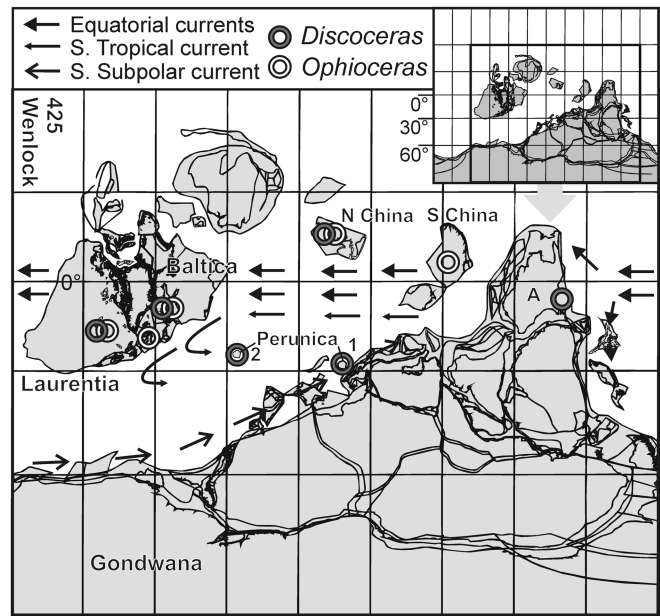


Figure 11. Paleogeographic map (after Torsvik and Cocks, 2013) showing distribution of tarphycerids and current system (after Wilde et al., 1991) in the late Wenlock. Alternative paleogeographic position of Central Bohemia (2), microcontinent Perunica (1), respectively, after Cocks and Torsvik (2002); this position corresponds better with the dispersion pattern of *Discoceras*, which never occurs in cool water areas influenced by subpolar current (Silurian tarphycerids are unknown in such areas of peri-Gondwana including Austria, France, Germany, Spain, etc.). A indicates Australia.

D. graftonense in distant continents (e.g., Laurentia and the Northeast China Plate). Empty shells could have been transported together with immigrants via current (Hamada et al., 1980). However, even the fine details of the shell sculpture are well preserved, suggesting that drift time and distance of empty shells were probably relatively short. The lack of bioerosion exclude their long-running drift between continents, which is consistent with rather limited post-mortem transport suggested for early Paleozoic nautiloid assemblages (Flower, 1957; Hewitt and Watkins, 1980; Frey, 1989).

Embryonic development and hatching time in Tarphycerida

The tarphycerid early shell is planispiral and tightly coiled with a very small umbilical perforation (Furnish and Glenister, 1964; Dzik, 1984). Its apex is blunt, the first chamber is curved and cup-like, and its length varies between 0.8–3.5 mm (Stumbar, 1959; Shimansky and Zhuravleva, 1961). The nepionic constriction characterizing the end of the embryonic phase in later nautilids is not present, but sculpture usually shows some change in growth line spacing. However, the sculpture development in embryonic and juvenile growth stages of tarphycerid shells is still poorly known. Consequently, the internal structures of the phragmocone have been used for determination of hatching time. This is a methodological approach inferred from extant *Nautilus* (Schindewolf, 1934; Stumbar, 1959).

A change in septal spacing between the seventh and eight chambers in *Nautilus* coincides with its emergence from an egg and formation of the nepionic constriction (Naef, 1921–1923;

Arnold et al., 1987; Ward, 1987; Tajika et al., 2015). Both reflect stress following the emergence of the hatched animal from the egg capsule (Arnold et al., 1987, 2010). Although the change in septal spacing is often linked with hatching in *Nautilus*, the intraspecific variability in the chamber length pattern is high (Stenzel, 1964). Except in *Ophioceras* (Turek and Manda, 2016), the variability of septal spacing in early ontogeny has not been studied in early Paleozoic nautiloids. A change in septal spacing in the first whorl in Ordovician tarphycerids was linked with hatching by Schindewolf (1934). Stumbur (1959, 1960) studied the early ontogeny in four genera representing the main evolutionary lineages of Ordovician tarphycerids. According to him, the hatching in tarphycerids is manifested by a sudden change in septal spacing—the first five to seven phragmocone chambers are longer than those following, in which length decreases gradually or abruptly. Following this hypothesis, the shell of early-hatched tarphycerids would have reached 1.25–1.5 whorls (WH = 4–8 mm, SD = 10–20 mm) and the phragmocone would consist of 5–7 chambers. Consequently, the hatching size in Tarphycerida would be close to that in post-Triassic nautilids (Wani et al., 2011). Stumbur (1959) supported his conclusion with two species of Ordovician *Discoceras* from Norway, illustrated by Sweet (1958). However, in these species, the change in septal spacing occurs between the eighth and ninth chambers. Sweet (1958) illustrated a specimen of *Estonioceras proteus demissa* Holm, 1885 in which the first two chambers are long and the third chamber markedly shortened; length of the following chambers gradually decreased, and after the ninth chamber increased again. Similarly, in *Eurystomites amplectens* Ruedemann, 1906 (pl. 18, fig. 5) and *Tarphyceras multicameratum* Ruedemann, 1906 (pl. 19, fig. 3), the first two chambers are longer than the following ones, which have almost equal lengths. In *Estonioceras imperfectum* (Quenstedt, 1845), the first chamber is long, the second chamber is shorter, the third to eighth chambers are almost the same length as the second; the ninth and following chambers are, again, markedly shorter (Stumbur, 1959; fig. 1a). Stumbur (1959) considered this development atypical, and assumed that the change in length between the seventh and eighth chambers indicates hatching. However, Shimansky and Zhuravleva (1961, p. 78) described another tarphycerid, *Trocholites* sp., with a high first chamber, a much shorter second chamber, followed by chambers with gradually increasing lengths.

Differing interpretations of the significance of changes in septal spacing during early development were addressed by Shimansky and Zhuravleva (1961, p. 84). They suggested two early ontogenetic paths in tarphycerids: (1) the shell in the egg capsule possessed one or two chambers and a body chamber (e.g., *Estonioceras imperfectum* and *Trocholites* sp.), in which case, the hatched animal differed from an adult in having a small curved shell; and (2) the embryonic shell consisted of five to six chambers and a body chamber (all other tarphycerids), and early-hatched animals possessing coiled shells that did not differ from adults. Nevertheless, such differing hatching size reflecting remarkable differences in the early development of hatchlings seems unlikely in such closely related taxa.

Turek and Manda (2016) demonstrated that septal spacing in Silurian *Ophioceras* is not coupled with hatching; hatching in this case undoubtedly preceded the change in septal spacing,

and juveniles had small curved shells with only one phragmocone chamber. However, *Ophioceras* is evolutionarily the youngest tarphycerid with some evolutionary novelties (e.g., a single ventral retractor muscle scar), and it is uncertain whether this type of embryonic development is applicable to older tarphycerid taxa.

Hatching indication in the Silurian *Discoceras*

The apex of *D. graftonense* is very similar in shape to that of Ordovician species of *Discoceras*. The earliest part of the shell is smooth; growth lines have been observed starting at about three-eighths of the first whorl and at whorl height 7–9 mm (SD = 16–22 mm) growth lines become raised. Length of the first chamber is ~2.7 mm, and the second to sixth phragmocone chambers are about one-third longer than following chambers (Fig. 13.4). Change in septal spacing occurs at a whorl height of 3 mm, when the shell attained one whorl (SD = 8.8 mm). Following Schindewolf's (1934) and Stumbur's (1959) approach to linking hatching with change in septal spacing, the shell of hatched *D. graftonense* would have reached one and one-half of a whorl and consisted of six phragmocone chambers plus a body chamber about half a whorl in length (WH = 5.2 mm, SD = 14 mm). A change in sculpture appeared immediately after the supposed indication of hatching. However, in specimen RM Mo 59803 (Fig. 13.1), the difference in chamber length is not obvious, because chamber lengths increase gradually.

The early shell of *D. lindstroemi* n. sp. is similar in shape and caecum position to *D. graftonense*, but differs in having much less pronounced growth lines, and a change in septal spacing occurs between the fifth and sixth chambers (WH = 2.6 mm, shell reached first whorl; Figs. 12.2, 13.5). If this change really indicates hatching time (Stumbur, 1959), then the whorl height would have reached 5 mm and shell diameter 12 mm.

A unique morphological feature—a deep groove situated at three-quarters of the first whorl (WH = 5 mm)—preceded the change in septal spacing (Fig. 4.9). The character of the groove and its adapical widening on the ventral side of the shell indicate serious mechanical damage caused by a predator. The injury originated after hatching and thus preceded the change in septal spacing (i.e., the shell of the hatched individual did not reach three-quarters of a whorl). Close to three-eighths of the first whorl, a change in shell coiling occurs, and growth lines are enhanced.

It is highly probable that the change in septal spacing did not coincide with hatching in *Discoceras*, as well as in Tarphycerida in general (see Turek and Manda, 2016). According to Stumbur (1959), the growth sculpture appeared in tarphycerids after hatching. However, in *Ophioceras*, fine growth lines have been observed on the first chamber. In the Ordovician species *Discoceras vasaemmense* Balashov, 1953, growth lines are discernible in about a half of the first whorl; they are regularly arranged, and in the adjacent juvenile part of the shell no striking change in their spacing has been observed. In Silurian *Discoceras*, growth lines appear at about three-eighths to one-half of the first whorl. Also at close to three-eighths of the first whorl, a slightly changed course of the shell spire and slight shell expansion has been observed. Taken together, there is as yet no

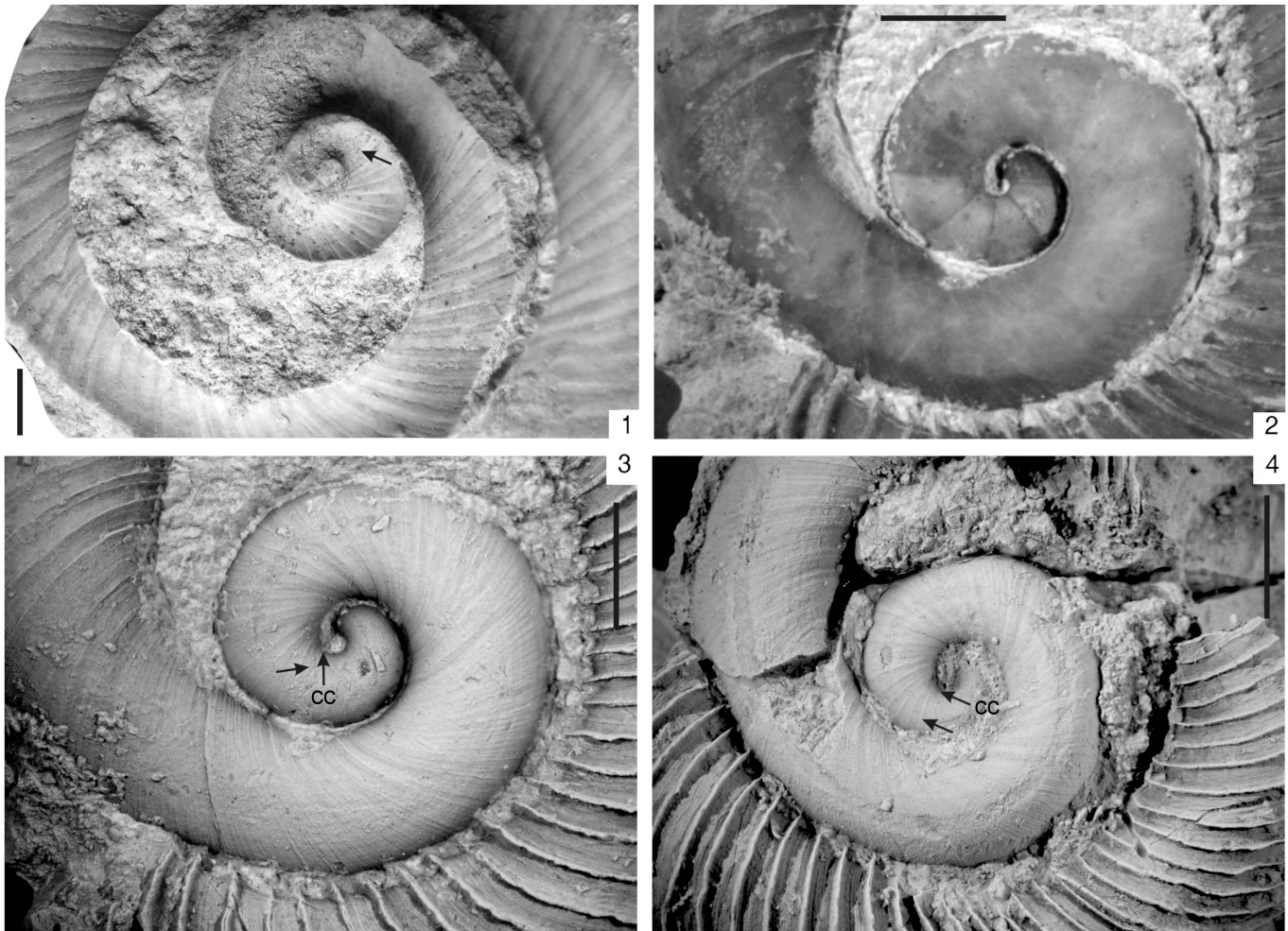


Figure 12. Early shell of *Discoceras lindstroemi* n. sp.; arrows indicate appearance of growth lines, cc indicates a change in shell coiling. (1) RM Mo 59442, lateral view, counterpart, Färö, Slite beds, late Sheinwoodian. Growth ridges regularly spaced, visible from first half of whorl. Minor shell damage seen dorsolaterally at one and one-half whorls; (2, 3) RM Mo 59420, lateral view, Othem, Samsugun, Slite beds, late Sheinwoodian; (4) RM Mo 59422, lateral view, specimen with well-developed lateral furrow, Othem, Samsugun, Slite beds, late Sheinwoodian. Specimens in Figure 12.1, 12.3, and 12.4 coated with ammonium chloride before photographing. Scale bars 5 mm.

clear evidence indicating hatching time in *Discoceras*. More likely, the hatchling possessed a curved shell (one-half or greater than three-quarters of a whorl) with two/three phragmocone chambers. Early hatched specimens thus differed in life from adults; the large volume of the first phragmocone chamber and the undeveloped hyponomic sinus suggest a macroplanktic habit. This is consistent with the early development of the evolutionarily youngest tarphycerid, *Ophioceras* (Turek and Manda, 2016). Longer chambers in the first whorl may be linked with accelerated growth before the shell reached one whorl, resulting in a stable shell orientation and higher rigidity of the shell.

Heteromorph shell in *Discoceras*

Silurian *Discoceras* shows shell morphology and habitat derived directly from Ordovician species. However, *D. lindstroemi* n. sp. (Figs. 4, 5) possesses a heteromorphic shell and elaborate sculpture. The early shell of *D. lindstroemi* n. sp. is tightly coiled, but subsequently, during growth, the second whorl becomes loosely coiled and then reverts again to being

tightly coiled (a slight indication of a second decoiling may once again appear; Fig. 13.3). In adult specimens, the final growth stage of the shell is straight. Changes in shell coiling resulted in a change in both the aperture orientation of the animal and the hydrodynamic properties of the shell (Chamberlain, 1976, 1981; Naglik et al., 2015). Growth stages with tightly coiled shells were nektonic, with the aperture oriented forward as in other tarphycerids (e.g., Flower, 1957; Westermann, 1998). In the growth stage with a loosely coiled shell, and in the adult stage possessing a protruding adapertural part of the shell, the aperture was oriented obliquely downward. Shells of the openly coiled stage exhibit longer phragmocone chambers, while whorl height slightly decreases. Apertures of fully grown specimens are constricted, with deepened ocular sinuses, suggesting well-developed eyes. Similarly, the deep hyponomic sinus suggests a well-developed hyponome.

A premature part of the shell, either loosely coiled or straight, is known in coiled Tarphycerida (Furnish and Glenister, 1964), Nautilida (Hyatt, 1894; Turek, 1975), and Oncocerida (Manda and Turek, 2011). A similar shell shape is unknown in post-Triassic nautilids, probably due to increasing

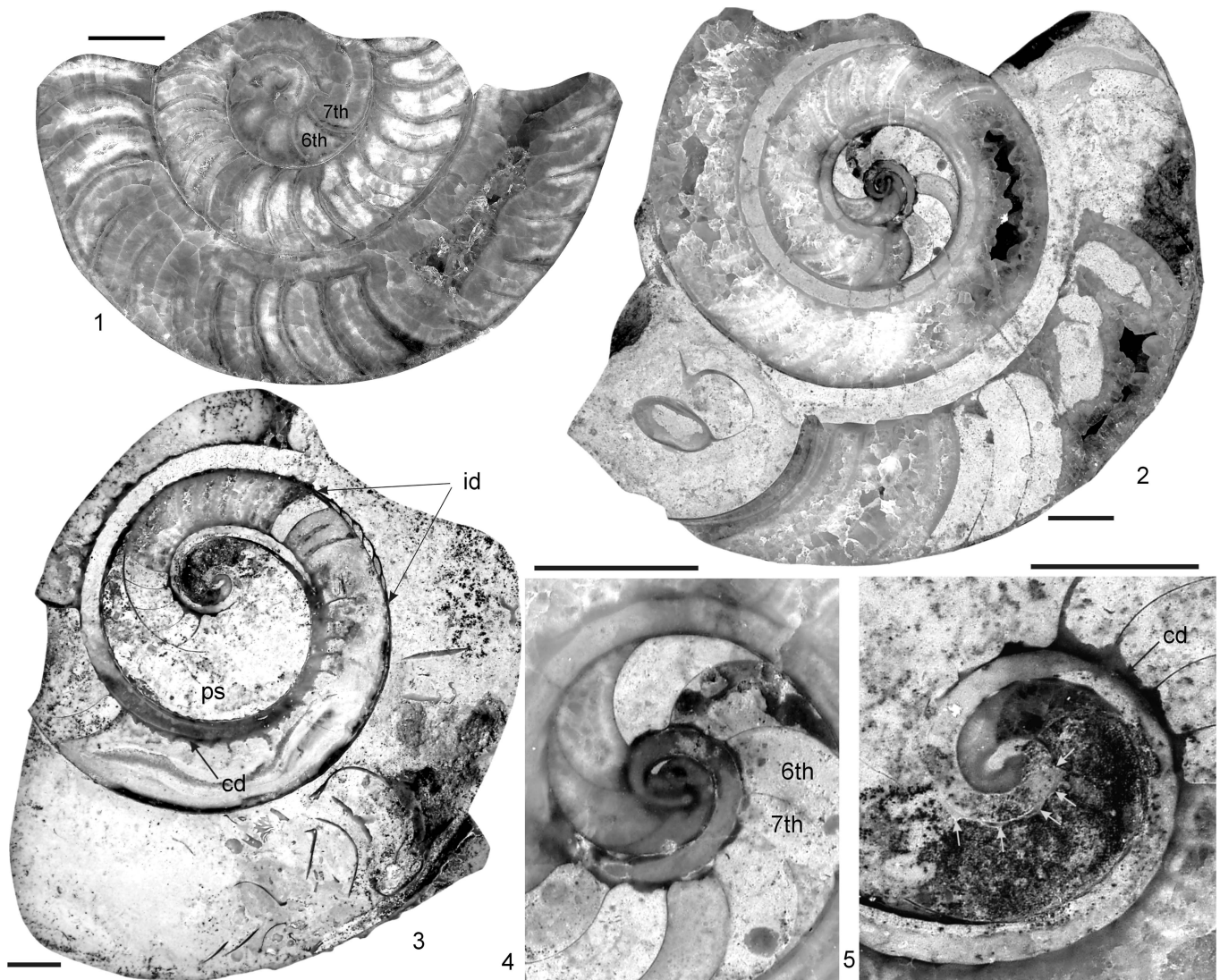


Figure 13. Inner structure in median section of shell in *Discoceras* from Silte beds (cd: cameral deposits, ps: pseudumbilicus, id: indication of the second decoiling). (1, 2, 4) *D. graffonense* (Meek and Worthen, 1870); (1) RM Mo 59803 Follingbo, Stora Vede 1, early Homerian; (2, 4) RM Mo 155975, Othem, Samsuguns 1, late Sheinwoodian, note rapidly increasing volume of siphuncle in third chamber and marked shortening of seventh chamber; (3, 5) *D. lindstroemi* n. sp. RM Mo 59884, white arrows indicate relicts of septa. Scale bars 5 mm.

competitive pressure among demersal predators (Klug et al., 2009). A final uncoiled growth stage is frequently coupled with the development of a contracted aperture; coiled nautiloids with an aperture oriented forward in life position always have an open aperture. The constricted aperture in nautiloids with an aperture oriented obliquely toward the seafloor probably had a protective function against predators (Flower, 1957), and provided an improved protection of the mantle margin against parasite infestation, which occurred rather frequently in the late Silurian *Ophioceras* (Turek and Manda, 2016). Growth of the uncoiled part of the shell coupled with the change in biological orientation could be beneficial to individuals on reaching a critical size. The length of uncoiled parts of the shells is highly variable; it is lower in more robust specimens and higher in specimens with slender shells. As in recent *Nautilus* (Collins and Ward, 1987), an increase in volume of the body chamber during adolescence probably corresponded to an increase in the size of the gonads (Stridsberg and Turek, 1997;

Manda and Turek, 2011; Turek and Manda, 2016). Changes in shell coiling occurred at least twice during the ontogeny of *D. lindstroemi* n. sp., but variability in the length and timing of the first decoiling remained high. A similar mode of coiling is indicated in Late Ordovician *Aphetoceras farnsworthi* (Billings, 1861) and *A. attenuatum* (Hyatt, 1894) (Ulrich et al., 1942, pl. 1, fig. 2, pl. 4, fig. 7). It is also seen in other estonioceratids (e.g., *Vasalemmoceras* Stumbr, 1962) and some lituitids (B. Kröger, personal communication, 2016). A poorly developed pseudumbilicus also occurs in the Silurian uranoceratid *Boionutilus sternbergi* (Barrande, 1865) (Turek, 2008, fig. 3d). Repeated significant decoiling of the shell resulted in a change in the living position of the animal, similar to that postulated for heteromorph ammonoids (Naglik et al., 2015; Kakabadze, 2016). The aperture in both cases reached a lower position in comparison with that in a tightly coiled shell (Fig. 14). Decoiling during the late juvenile and adolescent stages of *D. lindstroemi* n. sp. and other nautiloids might have enhanced

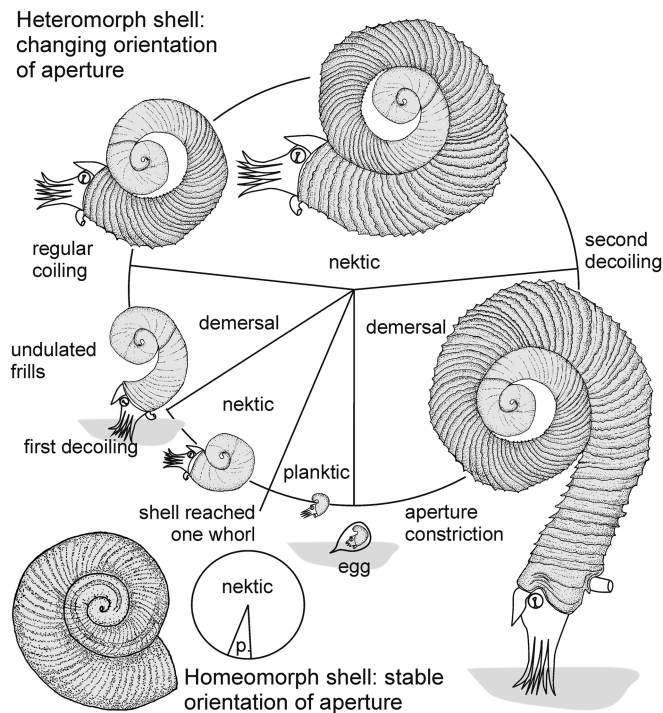


Figure 14. Schematic drawing of changing biological orientation in *D. lindstroemi* n. sp. throughout ontogeny.

the contact of the animal with the bottom (Stridsberg and Turek, 1997; Turek and Manda, 2016). Any advantage to this first decoiling is somewhat problematic. Decoiling of the shell could significantly alter the drag coefficient. Such a decoiled shell would have had a higher relative drag in comparison with a tightly coiled shell. The highly elaborate sculpture in *D. lindstroemi* n. sp. probably also negatively influenced the hydrodynamic properties of the shell (Chamberlain, 1976, 1981). The appearance of a heteromorph shell in *D. lindstroemi* n. sp. was probably an unsuccessful random evolutionary event among stratigraphically younger tarphycerids.

Conclusions

The first tarphycerids appeared in the early Tremadocian, then their diversity suddenly increased, reaching a maximum in the early Floian. Their generic diversity slowly declined in the late Upper Ordovician (Kröger and Zhang, 2009). Starting at the Late Ordovician extinction, diversity of tarphycerids was low throughout the Silurian, prior to their extinction just below Silurian-Devonian boundary. Three genera of tarphycerids are known from the Silurian, two of which survived the Late Ordovician extinction: *Discoceras* Barrande, 1867 and *Trocholites* Conrad, 1838, from which *Ophioceras* Barrande, 1865 probably diverged. Including the four species described here and evaluating previously published data, *Discoceras* comprises six Silurian species occurring in Llandovery (peri-Gondwanan Perunica) and Wenlock strata (Baltica, Laurentia, NE Gondwana, Northeast China Plate, peri-Gondwanan Perunica). The widespread geographic distribution of *D. graftonense* and the origin of four endemic species of *Discoceras* in the middle Wenlock represented the last weak

diversification and dispersion event of tarphycerids, terminated by the mid-Homerian extinction. A single tarphycerid genus, *Ophioceras* Barrande, 1865, which includes two long-ranging species, survived this extinction.

The geographic distribution of *Discoceras* in the Silurian is restricted to proximal areas of low-latitude platforms, and occasionally more distal areas of black shale sedimentation. The occurrence of *Discoceras* in off-shore shales and the dispersion of *D. graftonense* in distant continents indicate its migration potential and swimming ability in surface currents across an open sea.

Schindewolf (1934) was the first to link a change in septal spacing in the first whorl of Ordovician tarphycerids with their hatching phase. In this concept, hatching in tarphycerids is manifested by a sudden decrease in phragmocone chamber volume in phragmocone chambers six through eight. As a consequence, the shell of early-hatched tarphycerids would have reached slightly more than one whorl, and a shell diameter of 10–20 mm. Nevertheless, there is as yet no clear evidence indicating hatching time in *Discoceras*. A repaired injury in a Silurian *Discoceras* that took place after hatching indicates that the shell of the hatched individual did not reach three-quarters of a whorl, and thus hatching preceded the change in septal spacing. Moreover, this change in spacing is not present in every specimen studied. Growth lines appear in Silurian *Discoceras* at about three-eighths to one-half of the first whorl. Their appearance on the shell surface coincides with a slightly changed course of the shell spire, and a slight shell expansion. More likely, the hatchling possessed only a curved shell with two or three phragmocone chambers. Early hatched specimens thus differed in habit from adults (demersal swimmers). The large volume of the first phragmocone chamber and the undeveloped hyponomic sinus in juveniles suggest a macroplanktic habit, which is consistent with early development of the evolutionarily youngest tarphycerid *Ophioceras* (Turek and Manda, 2016).

Silurian *Discoceras* retained the morphology and habitats of its Ordovician ancestors. Nevertheless, the early shell of *D. lindstroemi* n. sp. is tightly coiled, but subsequently, during growth, the second whorl becomes loosely coiled and then reverts again to tight coiling; furthermore, a slight indication of a second decoiling may once again appear. A heteromorphic planispiral shell with coiling that changed during ontogeny resulted in changing aperture orientation and maneuverability in life. The appearance of a heteromorphic shell in *D. lindstroemi* n. sp. was probably an unsuccessful random evolutionary event among Silurian tarphycerids. Simultaneously, it was the stratigraphically first appearance of a well-elaborated heteromorph shell in externally shelled cephalopods in the fossil record.

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References

- Agassiz, L., 1847, An Introduction to the Study of Natural History, in a series of lectures delivered in the Hall of the College of Physicians and Surgeons: New York, Greeley and McElrath, 58 p.
- Arnold, J.M., Landman, N.H., and Mutvei, H., 1987, Development of the embryonic shell of *Nautilus*, in Saunders, W.B., and Landman, N.H., eds., *Nautilus: The Biology and Paleobiology of a Living Fossil*: New York, Plenum Press, p. 373–400.
- Arnold, J.M., Landman, N.H., and Mutvei, H., 2010, Development of the embryonic shell of *Nautilus*, in Saunders, W.B., and Landman, N.H.E., eds., *Nautilus The Biology and Paleobiology of a Living Fossil*. Reprint with additions: Dordrecht, Heidelberg, London, New York, Springer Verlag, p. 373–400.
- Babin, C., and Gutiérrez-Marco, J.C., 1992, Intérêt paléobiogéographique de la présence du genre *Trocholites* (Cephalopoda, Nautiloidea) dans le Dobrotivá (Llandeilo) inférieure d'Espagne: Neues Jahrbuch für Geologie und Paläontologie, Monatshefte, v. 9, p. 519–541.
- Balashov, Z.G., 1953, Coiled and semi-coiled Ordovician and Silurian nautiloids from Baltic platform, in Balashov, Z.G., and Stumbur X.A., eds., Stratigraphy and fauna of the Ordovician and Silurian of the western Russian platform: Trudy Vsesoyuznogo Neftyanogo Nauchno-issledovatel'skogo Geologo-razvedochnogo Instituta, Novaya Seriya, v. 73, p. 217–269. [in Russian]
- Barrande, J., 1847, Über die Brachiopoden der silurischen Schichten von Böhmen. II. Teil: Naturwissenschaftliche Abhandlungen (W. Haidinger, Wien), v. 2, 155–256.
- Barrande, J., 1865–1877, Système silurien du Centre de la Bohême, I. ère partie: Recherches Paléontologiques, vol. II, Classe de Mollusques, Ordre des Céphalopodes: Prague and Paris, privately published, 1865, ser. 6, pl. 1–107; 1866, ser. 7, pl. 108–244; 1867, ser. 1, 712 p.; 1868, ser. 8, pl. 245–350; 1870, ser. 2, 266 p., ser. 9, pl. 351–460; 1874, ser. 3, 804 p.; 1877, ser. 4, 742 p., ser. 5, 743 p.; Supplement 1, 297 p.; Supplement 2, pl. 461–544.
- Barskov, I.S., 1972, Late Ordovician and Silurian Cephalopod Mollusks of Kazakhstan and Middle Asia: Moscow, Nauka Press, 109 p. [in Russian]
- Billings, E., 1861, New species of fossils from Lower Silurian rocks of Canada: Canadian Naturalist and Geologist, Proceedings of the Society of Natural History Montreal, v. 5, p. 161–177.
- Calner, M., 2008, Silurian global events—at the tipping point of climate change, in Ashraf M.T., ed., Mass Extinctions: Berlin and Heidelberg, Springer Verlag, p. 21–58.
- Calner, M., and Jeppsson, L., 2003, Carbonate platform evolution and conodont stratigraphy during the middle Silurian Mulde Event, Gotland, Sweden: Geological Magazine, v. 140, p. 173–203. DOI: 10.1017/S0016756802007070.
- Calner, M., Jeppsson, L., and Munnecke, A., 2002, The Silurian of Gotland—Part I: Review of the stratigraphic framework, event stratigraphy, and stable carbon and oxygen isotope development: Erlanger geologische Abhandlungen, Sonderband, v. 5, p. 113–131.
- Chamberlain, J.A. Jr., 1976, Flow patterns and drag coefficients of cephalopod shells: Palaeontology, v. 19, p. 539–563.
- Chamberlain, J.A. Jr., 1981, Hydrodynamical design of fossil cephalopods, in House M.R., and Senior J.R., eds., The Ammonoidea. The Evolution, Classification, Mode of Life and Geological Usefulness of a Major Fossil Group: London, New York, Toronto, Sydney, San Francisco, Academic Press, p. 289–336.
- Chapman, E.J., 1857, On the occurrence of the genus *Cryptoceras* in Silurian rocks: Canadian Journal, New series, v. 2, p. 264–268.
- Chen, J., Liu, G., and Chen, T., 1981, Silurian nautiloid faunas of central and southwestern China: Memoirs of the Nanjing Institute of Geology and Palaeontology, v. 13, 104 p. [in Chinese with English summary]
- Cocks, L.R.M., and Torsvik, T.H., 2002, Earth Geography from 500 to 400 million years ago: a faunal and palaeomagnetic review: Journal Geological Society London, v. 159, p. 631–644.
- Collins, D., and Ward, P., 1987, Adolescent growth and maturity in *Nautilus*, in Saunders, W.B., and Landman, N.H., eds., *Nautilus: The Biology and Paleobiology of a Living Fossil*: New York, Plenum Press, p. 421–434.
- Conrad, T.A., 1838, Report on the Palaeontological Department of the Survey (New York): New York State Geological Survey, Annual Report, v. 2, p. 107–119.
- De Baets, K., Landman, N., and Tanabe, K., 2015, Ammonoid embryonic development, in Klug, C., Korn, D., De Baets, K., Kruta, I., and Mapes, R., eds., Ammonoid Paleobiology: From Anatomy to Ecology: Dordrecht, Heidelberg, New York, London, Springer, p. 113–205.
- Dzik, J., 1984, Phylogeny of the Nautiloidea: Paleontologia Polonica, v. 45, 155 p.
- Eichwald, E., von, 1842, Die Urwelt Russlands. 2. Heft. Beiträge zur Kenntnis des Russischen Reiches: St. Petersburg, K.E. von Baer, 184 p.
- Etheridge, R., 1904, On the occurrence of a litiutean in the Upper Silurian rocks of Bowring, New South Wales: Records of the Australian Museum, v. 5, p. 75–77.
- Evans, D.H., Ghobadipour, M., Popov, L.E., and Jahangir, H., 2015, An Early Silurian (Aeronian) cephalopod fauna from Kopet-Dagh, north-eastern Iran: including the earliest records of non-orthocerid cephalopods from the Silurian of Northern Gondwana: Bulletin of Geosciences, v. 90, p. 479–507.
- Flower, R.H., 1950, Order Tarphyceratida, Order Barrandoceratida, in Flower, R.H., and Kummel, B., A classification of the Nautiloidea: Journal of Paleontology, v. 24, p. 604–616.
- Flower, R.H., 1955, Saltations in nautiloid coiling: Evolution, v. 9, p. 244–260.
- Flower, R.H., 1957, Nautiloids of the Paleozoic: Memoirs of the Geological Society of America, v. 67, p. 829–852.
- Foerste, A.F., 1925, Notes on cephalopod genera; chiefly coiled Silurian forms: Bulletin Denison University, Journal of the Scientific Laboratories, v. 21, p. 1–69.
- Frey, R., 1989, Paleocology of a well-preserved nautiloid assemblage from a Late Ordovician shale unit, southwestern Ohio: Journal of Paleontology, v. 63, p. 604–620.
- Frey, R.C., Beresi, M.S., Evans, D.H., King, A.H., and Percival, I.G., 2004, Nautiloid cephalopods, in Webby, B.D., Paris, F., Droser, M.L., and Percival, I.G., eds., The Great Ordovician Biodiversification Event: New York, Columbia University Press, p. 209–213.
- Furnish, W.M., and Glenister, B.F., 1964, Nautiloidea—Tarphycerida, in Moore, R.C., ed., Treatise on Invertebrate Paleontology, part K, Mollusca 3: Boulder and Lawrence, Geological Society of America and The University of Kansas Press, p. 343–368.
- Hamada, T., Tanabe, K., and Hayasaka, S., 1980, The first capture of a living chambered *Nautilus* in Japan: Science Paper of the Collegiums General Education, University of Tokyo, v. 30, p. 63–66.
- Havlíček, V., and Storch, P., 1990, Silurian brachiopods and benthic communities in the Prague Basin (Czechoslovakia): Rozpravy Ústředního ústavu geologického, v. 48, p. 1–275.
- Hedström, H., 1917, Über die Gattung *Phragmoceras* in der Obersilurformation Gotlands: Sveriges Geologiska Undersökning, v. 15, p. 1–35.
- Hewitt, R.A., and Watkins, R., 1980, Cephalopod ecology across a late Silurian shelf tract: Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen: v. 160, p. 96–117.
- Histon, K., 2012, An Alpine immigrant: *Phragmoceras* Broderip, 1839 from the Silurian of the Carnic Alps (Austria): Geobios, v. 45, p. 41–48.
- Holland, C.H., 1985, Form and function in Silurian Cephalopoda: Special Papers in Palaeontology, v. 32, p. 151–164.
- Holland, C.H., 2010, Coiled nautiloid cephalopods from the British Silurian: Proceedings of the Geologists' Association, v. 121, p. 13–23.
- Holm, G., 1885, Über die innere Organisation einiger silurischer Cephalopoden: Palaeontologische Abhandlungen, v. 3, p. 1–27.
- Hyatt, A., 1894, Phylogeny of an acquired characteristic: Proceedings American Philosophical Society, v. 32, p. 349–647.
- Jeppsson, L., and Calner, M., 2003, The Silurian Mulde event and a scenario for secundo-secundo events: Transactions Royal Society of Edinburgh, Earth Sciences, v. 93, p. 135–154.
- Jeppsson, L., Viira, V., and Männik, P., 1994, Silurian conodont-based correlations between Gotland (Sweden) and Saaremaa (Estonia): Geological Magazine, v. 131, p. 201–218.
- Kakabadze, M.V., 2016, Speculation on the ethology of some heteromorph ammonites: Swiss Journal of Palaeontology, v. 135, p. 63–68.
- Klug, C., Kröger, B., Kiessling, W., Mullins, G.L., Servais, T., Frýda, J., Korn, D., and Turner, S., 2009, The Devonian nekton revolution: Lethaia, v. 43, p. 465–477.
- Korn, D., and Klug, C., 2003, Morphological pathways in the evolution of Early and Middle Devonian ammonoids: Paleobiology, v. 29, p. 329–348.
- Kříž, J., 1998, Silurian, in Chlupáč, I., Havlíček, V., Kříž, J., Kukul, Z., and Štorch, P., Paleozoic of the Barrandian (Cambrian to Devonian): Praha, Český Geologický Ústav, p. 79–101.
- Kröger, B., 2003, The size of siphuncle in cephalopod evolution: Senckenbergiana Lethaea, v. 83, p. 39–52.
- Kröger, B., 2005, Adaptive evolution in Paleozoic coiled cephalopods: Paleobiology, v. 31, p. 253–268.
- Kröger, B., 2013, The cephalopods of the Boda Limestone, Late Ordovician, of Dalarna, Sweden: European Journal of Taxonomy, v. 41, p. 1–110.

- Kröger, B., and Ebbestad, J.O.R., 2014, Palaeoecology and palaeogeography of Late Ordovician (Katian–Hirnantian) cephalopods of the Boda Limestone, Siljan district, Sweden: *Lethaia*, v. 47, p. 15–30.
- Kröger, B., and Zhang, Y.B., 2009, Pulsed cephalopod diversification during the Ordovician: Palaeogeography, Palaeoclimatology, Palaeoecology, v. 273, p. 174–183.
- Kröger, B., Servais, T., and Zhang, Y., 2009, The Origin and Initial Rise of Pelagic Cephalopods in the Ordovician: *PLoS ONE*, 4 (9), e7262.
- Kröger, B., Ebbestad, J.O.R., Höglström, H.E.S.A., and Frisk, M., 2011, Mass concentration of Hirnantian cephalopods from the Siljan District, Sweden: taxonomy, palaeoecology and palaeobiogeographic relationships: *Fossil Record*, v. 14, 35–53.
- Lai, C., 1982, Ordovician cephalopods from Xainza, Xizang (Tibet): *Acta Palaeontologica Sinica*, v. 21, p. 553–559. [in Chinese]
- Lai, C., 1987, Ordovician cephalopods from Mt. Querquertal (Mt. Charchag), Xinjiang: Professional Papers in Stratigraphy and Palaeontology, v. 17, 227–246. [in Chinese]
- Lai, C., and Wang, M., 1986, Ordovician cephalopods from Mt. Altun area N. W. China: *Acta Palaeontologica Sinica*, v. 25, 248–256. [In Chinese]
- Laufeld, S., 1974, Reference localities for palaeontology and geology in the Silurian of Gotland: *Sveriges Geologiska Undersökning C*, v. 705, 172 p.
- Li, W., Rong, J., Dong, D., Yang, D., Su, Y., and Wang, Y., 1983, Silurian and Devonian rocks of Bateaoabao area in Darhan Mumingan Joint Banner, Inner Mongolia, in Li, W., Rong, J., and Dong, D., ed., *Silurian and Devonian Rocks and Faunas of the Bateaoabao Area in Darhan Mumingan Joint Banner, Inner Mongolia*: Nanjing, The People's Publishing House, p. 1–25. [in Chinese]
- Lossen, C., 1860, Über einige Lituiten: *Zeitschrift der deutschen geologischen Gesellschaft*, v. 12, p. 15–28.
- Manda, Š., 2008a, *Trocholites* Conrad, 1838 (Nautiloidea, Tarphycerida) in the Middle Ordovician of the Prague Basin and its palaeobiogeographical significance: *Bulletin of Geosciences*, v. 83, p. 327–334.
- Manda, Š., 2008b, Palaeoecology and palaeogeographic relations of the Silurian phragmoceratids (Nautiloidea, Cephalopoda) of the Prague Basin (Bohemia): *Bulletin of Geosciences*, v. 83, p. 39–62.
- Manda, Š., and Turek, V., 2011, Late Emsian *Rutoceratoidea* (Nautiloidea) from the Prague Basin, Czech Republic: morphology, diversity and palaeoecology: *Palaeontology*, v. 54, p. 999–1024.
- Manten, A., 1971, Silurian reefs of Gotland: *Developments in Sedimentology*, v. 13, Berlin, Springer Verlag, 539 p.
- Meek, F.B., 1873, Descriptions of invertebrate fossils of the Silurian and Devonian systems. *Geological Survey of Ohio 1*, *Paleontology*: Columbus, Geological Survey of Ohio, 246 p.
- Meek, F.B., and Worthen, A.H., 1870, Descriptions of new species and genera of fossils from the Palaeozoic rocks of the western states: *Proceedings of the Academy of Natural Sciences*, Philadelphia, v. 22, p. 22–56.
- Meek, F.B., and Worthen, A.H., 1875, Descriptions of Invertebrates, in *Paleontology of Illinois*, Volume 6: Springfield, Geological Survey of Illinois, p. 491–532.
- Murchison, R.I., 1839, The Silurian System Founded on Geological Researches in the Counties of Salop, Hereford, Padnor, with Descriptions of the Coal Fields and Overlying Formations: London, John Murray, 768 p.
- Naglik, C., Tajika, A., Chamberlain, J., and Klug, C., 2015, Ammonoid locomotion, in Klug, C., Korn, D., De Baets, K., Kruta, I., and Mapes, R.H., eds., *The Ammonoid Paleobiology: From Anatomy to Ecology*, Topics in Geobiology, Heidelberg, New York, London, Springer, v. 43, p. 649–688.
- Naef, A., 1921–1923, Die Cephalopoden: Fauna e Flora del Golfo di Napoli: *Stazione Zoologica Napoli Monograph*, v. 35, 821 p.
- Newell, F.H., 1888, Niagara Cephalopods from Northern Indiana: *Proceedings of the Boston Society of Natural History*, v. 23, p. 466–486.
- Quenstedt, F.A., 1845, *Petrefactenkunde Deutschlands*, 1. Abteilung, 1. Band Cephalopoden: Tübingen, Fues Verlag, 580 p.
- Rasmussen, J.A., and Surlyk, F., 2012, Rare finds of the coiled cephalopod *Discoceras* from the Upper Ordovician of Bornholm, Denmark: *Bulletin of the Geological Society of Denmark*, v. 60, p. 15–22.
- Remelé, A., 1890, Untersuchungen über die versteinierungsführenden Diluvialgeschiebe des norddeutschen Flachlandes, mit besonderer Berücksichtigung der Mark Brandenburg: Berlin, Springer Verlag, 108 p.
- Roemer, C.F., 1861, Die fossile Fauna der silurischen Diluvial-Geschiebe von Sadewitz bei Oels in Niederschlesien—Eine palaeontologische Monographie: Breslau, Druck Robert Nischkowsky, 79 p.
- Ruedemann, R., 1906, Cephalopoda of the Beekmantown and Chazy formations of the Champlain Basin: *Bulletin New York State Museum*, v. 90, 605 p.
- Salter, J.W., 1863, On *Peltoicaris*, a new genus of Silurian Crustacea: *Quarterly Journal of the Geological Society of London*, v. 19, 87–92.
- Schindewolf, O.H., 1934, Zur Stammesgeschichte der Cephalopoden: *Jahrbuch der Preussischen Geologischen Landesanstalt*, v. 55, p. 258–283.
- Shaver, R.H., 1991, A history of study of Silurian reefs in the Michigan Basin environs: *Geological Society of America Special Paper*, v. 256, p. 101–138.
- Shimansky, V.N., and Zhuravleva, F.A., 1961, Fundamental questions of the systematics of the nautiloids and the relationship of this group: *Trudy Paleontologitscheskogo Instituta Akademii Nauk*, v. 90, 175 p. [in Russian]
- Shimer, H.W., and Shrock, R.R., 1944, *Index Fossils of North America*: New York, John Wiley and Sons, 837 p.
- Sowerby, J., 1812–18, *The mineral conchology of the Great Britain; or coloured figures and descriptions of those remains of testaceous animals or shells which have been preserved at various times and depths in the Earth*, v. Vol. 2: London, B. Meredith, 251 p.
- Stenzel, H.B., 1964, *Living Nautilus*, in Moore, R.C., ed., *Treatise on Invertebrate Paleontology*, part K, Mollusca 3: Boulder, Colorado and Lawrence, Kansas, Geological Society of America and The University of Kansas Press, p. K59–K93.
- Štorch, P., 1994, Graptolite biostratigraphy of the Lower Silurian (Llandovery and Wenlock) of Bohemia: *Geological Journal*, v. 29, p. 137–165.
- Štorch, P., 2006, Facies development, depositional settings and sequence stratigraphy across the Ordovician–Silurian boundary: a new perspective from Barrandian area of the Czech Republic: *Geological Journal*, v. 41, p. 163–192.
- Strand, T., 1934, *The Upper Ordovician Cephalopods of the Oslo Area*: *Norsk geologic Tidsskrift*, v. 14, p. 1–117.
- Stridsberg, S., 1985, Silurian oncocerid cephalopods from Gotland: *Fossils and Strata*, v. 18, 65 p.
- Stridsberg, S., 1988a, A Silurian cephalopod genus a reinforced frilled shell: *Palaeontology*, v. 31, p. 651–663.
- Stridsberg, S., 1988b, A stray cephalopod in the late Silurian of Sardinia: *Bollettino della Società Paleontologica Italiana*, v. 27, p. 83–85.
- Stridsberg, S., and Turek, V., 1997, A revision of the Silurian nautiloid genus *Ophioceras* Barrande: *GFF*, v. 19, p. 21–36.
- Stumbur, X.A., 1959, Embryonic shell of some Ordovician tarphyceratids: *Paleontologitscheskij Zhurnal*, v. 2, p. 25–29. [in Russian]
- Stumbur, X.A., 1960, Early shell of nautiloids: *Izvestia Akademii nauk Estonskoj SSR*, v. 4, p. 368–378. [in Russian]
- Stumbur, H.A., 1962, Nautiloid distribution in the Ordovician of Estonia (with description of some new genera): *Trudy Instituta Geologii Akademii Nauk Estonskoi SSR (ENSV Teaduste Akadeemia Geoloogia Instituudi Uurimused)*, v. 10, p. 131–147. [in Russian]
- Sweet, W.C., 1958, The Middle Ordovician of the Oslo region of Norway. 10. Nautiloid cephalopods: *Norsk geologisk Tidsskrift*, v. 3, 178 p.
- Tajika, A., Morimoto, N., Wani, R., Naglik, C., and Klug, C., 2015, Intraspecific variation of phragmocone chamber volumes throughout ontogeny in the modern nautilid *Nautilus* and the Jurassic ammonite *Normannites*: *PeerJ* 3: e1306.
- Taylor, P.D., and Vinn, O., 2006, Convergent morphology in small spiral worm tubes “*Spirorbis*” and its palaeoenvironmental implications: *Journal of the Geological Society*, v. 163, p. 225–228.
- Teichert, C., and Glenister, B.F., 1954, Early Ordovician cephalopod fauna from northwestern Australia: *Bulletins of American Paleontology*, v. 35, p. 7–112.
- Torsvik, T.H., and Cocks, L.R.M., 2013, New global palaeogeographical reconstructions for the Early Palaeozoic and their generation: *Geological Society, London, Memoirs*, v. 38, p. 5–24.
- Turek, V., 1975, Genus *Kosovoceras* gen. n. in the Silurian of Central Bohemia: *Sborník geologických věd, Paleontologie*, v. 17, p. 7–42.
- Turek, V., 1978, Biological and stratigraphical significance of the Silurian nautiloid *Aptychopsis*: *Lethaia*, v. 11, p. 127–138.
- Turek, V., 1983, Hydrodynamic conditions and the benthic community of Upper Wenlockian calcareous shale in western part of the Barrandian (Kosov quarry): *Časopis pro Mineralogii a Geologii*, v. 28, p. 245–260.
- Turek, V., 2008, *Boionautilus* gen. nov. from the Silurian of Europe and North Africa (Nautiloidea, Tarphycerida): *Bulletin of Geosciences*, v. 83, p. 141–152.
- Turek, V., and Manda, Š., 2016, Early ontogeny, anomalous growth and healed injuries in Silurian nautiloid *Ophioceras* Barrande—Implication for hatching in Tarphycerida (Nautiloidea): *Bulletin of Geosciences*, v. 91, 331–366.
- Ulrich, E.O., Foerste, A.F., Miller, A.K., and Furnish, W.M., 1942, *Ozarkian and Canadian Cephalopods Part I: Nautilicones*: *Geological Society of America Special Papers*, v. 37, 157 p.
- Wani, R., Kurihara, K., and Ayyasami, K., 2011, Large hatchling size in Cretaceous nautiloids persists across the end-Cretaceous mass extinction: new data of Hecroglossidae hatchlings: *Cretaceous Research*, v. 32, p. 618–622.
- Ward, P.D., 1987, *Natural History of Nautilus*: Boston, Allen and Unwin, 263 p.
- Westermann, G.E.G., 1998, Life habits of nautiloids, in Savazzi, E., ed., *Functional Morphology of the Invertebrate Skeleton*: London, Wiley, p. 263–298.
- Whiteaves, J.F., 1897, The fossils of the Galena-Trenton and Black River Formations of Lake Winnipeg and its vicinity: *Palaeozoic fossils*, v. 3: Montreal, Geological Survey of Canada, Dawson Brothers, p. 192–242.

- Whitfield, R.P., 1882, *Palaeontology*, Part 3: Geology of Wisconsin, v. 4, p. 163–363.
- Wilde, P., Berry, W.B.N., and Quinby, H.M.S., 1991, Silurian oceanic and atmospheric circulation and chemistry, in Bassett, M.G., Lane, P.D., and Edwards, D., eds., *The Murchison Symposium; Proceedings of an International Conference on the Silurian System: Special Papers in Palaeontology*, v. 44, p. 123–143.
- Woodward, H., 1866, On a new genus phyllopodous Crustacea from the Moffat Shales (Llandeilo Flags) Dumfriesshire: *Quarterly Journal of the Geological Society of London*, v. 22, 503–505.
- Zou, X.P., 1983, Silurian Nautiloids from Bateaobao, Darhan Muminglan Joint Banner, Inner Mongolia, in Li, W., Rong, J., and Dong, D., eds., *Silurian and Devonian Rocks and Faunas of the Bateaobao Area in Darhan Muminglan Joint Banner, Inner Mongolia: Nanjing, The People's Publishing House*, p. 165–173. [in Chinese]

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